

CAIMAN CF. LATIROSTRIS (ALLIGATORIDAE, CAIMANINAE) IN THE LATE MIOCENE PALO PINTADO FORMATION, SALTA PROVINCE, ARGENTINA: PALEOGEOGRAPHIC AND PALEOENVIRONMENTAL CONSIDERATIONS



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Abstract. The three recognized species of *Caiman* –*C. latirostris*, *C. yacare* and *C. crocodilus*– currently live in northern and central South America. Except for the fragmentary dentary of a putative *Caiman* from Oligocene rocks in Brazil, the genus has been reliably recorded in rocks of ages spanning the Neogene, when species of *Caiman* were a constant component of the South American crocodyliofauna. The major taxonomical diversification of *Caiman* occurred during the late Miocene, which is well documented in the area of Paraná (northeastern Argentina). Fossil crocodylians in Paraná are represented by one gavialid and caimanines, with at least five species of *Caiman* (including *C. latirostris*). This assemblage represents the southernmost record of Crocodylia living in "Amazonia" during the Miocene. In this work we confirm the record of Miocene caimans outside the Paraná and we prove the presence of *Caiman* cf. *latirostris* in present-day northwestern Argentina during the late Miocene. The taxonomic identification is based on a fragment of a left mandible with the same ornamentation, outline and dentition as *Caiman*, and with a symphyseal morphology similar to that of *Caiman latirostris*. The material comes from the upper part of the Palo Pintado Formation in the southern region of Valle Calchaquí (Salta Province). This unit was deposited in a sand-gravel fluvial system with associated ponds between 10.29 ±0.11 Ma (K/Ar) and 5.27 ±0.28 Ma (²⁰⁶Pb/²³⁸U).

Key words. Crocodylia. *Caiman*. Neogene. Salta Province. Argentina.

Resumen. CAIMAN CF. LATIROSTRIS (ALLIGATORIDAE, CAIMANINAE) EN LA FORMACIÓN PALO PINTADO DEL MIOCENO TARDÍO, PROVINCIA DE SALTA, ARGENTINA: CONSIDERACIONES PALEOGEOGRÁFICAS Y PALEOAMBIENTALES. El género *Caiman* está actualmente representado por tres especies, *C. latirostris*, *C. yacare* y *C. crocodilus*, distribuidas en el norte-centro de América del Sur. Exceptuando un resto fragmentario de dentario de un probable *Caiman* proveniente de rocas oligocenas de Brasil, este género se registra con seguridad desde el Mioceno y sus especies constituyeron un componente constante de la crocodylofauna sudamericana durante el Neógeno. La mayor diversificación taxonómica de *Caiman* ocurrió durante el Mioceno tardío y se encuentra bien documentada en el área de Paraná (noreste de Argentina). Los cocodrilos fósiles allí registrados son una especie de gavial y caimaninos, representados por al menos cinco especies de *Caiman* (incluyendo *C. latirostris*). Esta asociación constituye el registro más austral de Crocodylia que habitaron la "Amazonia" durante el Mioceno. En este trabajo se confirma el registro de caimanes fuera del área de Paraná y se demuestra la presencia de *Caiman* cf. *latirostris* en el Mioceno tardío en el noroeste de Argentina. La determinación taxonómica se basa en un fragmento izquierdo de mandíbula, con la misma ornamentación, contorno y dentición que en *Caiman*, y con una morfología de la región sinfisal similar a la que presenta *C. latirostris*. El material proviene de niveles superiores de la Formación Palo Pintado en el área sur de los Valles Calchaquíes (Provincia de Salta). Esta unidad fue depositada en un ambiente fluvial areno-gravoso con lagunas asociadas y pantanos entre los 10.29 ±0.11 Ma (K/Ar) y los 5.27 ±0.28 Ma (²⁰⁶Pb/²³⁸U).

Palabras clave. Crocodylia. *Caiman*. Neógeno. Provincia de Salta. Argentina.

THE so called "modern caimans" include the living representatives of Caimaninae, which is a monophyletic group of mainly South American Alligatoridae (Brochu, 1999, 2011). Six species of caimanines are currently recognized: *Caiman crocodilus* (Linnaeus, 1758), *Caiman yacare* (Daudin, 1802) (occasionally considered a subspecies of *C. crocodilus*), *Caiman latirostris* (Daudin, 1802), *Melanosuchus niger* (Spix, 1825),

Paleosuchus palpebrosus (Cuvier, 1807) and *Paleosuchus trigonatus* (Cuvier, 1807). Although the monophyly of this group is widely supported by morphological and molecular evidence (Densmore and Owen, 1989; Norell, 1988; Brochu, 1999, 2003; Gatesy et al., 2003; Aguilera et al., 2006; Hrbek et al., 2007; Willis et al., 2007; Willis, 2009), the systematics of the genus *Caiman* Spix, 1825, is controversial (Norell, 1988;

Poe, 1997; Brochu, 1999). The most accepted hypothesis divides the lineage in three species (*C. latirostris*, *C. yacare* and *C. crocodilus*; Busack and Pandya, 2001, see also Crocodile Specialist Group's web site <http://www.iucncsg.org/pages/Publications.html> and cites herein), mainly distributed in central-eastern South America. Except for *Caiman tremembensis* Chiappe, 1988—a putative member of the genus described on the basis of material from Oligocene rocks in southeastern Brazil (Tremembé Formation; São Paulo State) and which needs to be revised in the light of modern knowledge—*Caiman* has been recorded only since the early Miocene (Langston, 1965) in South America. Its species have been consistent components of the South American crocodyliofauna during the Neogene.

During the Miocene, the evolutionary history of Crocodylia (*sensu* Benton and Clark, 1988) in South America was characterized by large morphological and taxic diversification of lineages such as Alligatoridae (Brochu, 2003; Riff *et al.*, 2010; Bona *et al.*, in press). Alligatorid genera show a wide geographic distribution (*Mourasuchus* Price, 1946; *Purussaurus* Barbosa-Rodrigues, 1892; *Caiman*) and local species-level endemism (Bona *et al.*, in press). This situation is particularly documented and studied in the north of the continent (*e.g.*, La Venta, Colombia; Urumaco, Venezuela; Acre, Brazil). However, in southern areas of Amazonia (*sensu* Hoorn *et al.*, 2010), the most abundant and best studied record of crocodylians is that of northeastern Argentina (NEA; Bravard, 1858; Burmeister, 1883; Ambrosetti, 1887; Scalabrini, 1887; Rovereto, 1912; Rusconi, 1933, 1935; Langston, 1965; Gasparini, 1973; Langston and Gasparini, 1997; Cione *et al.*, 2000, Bona *et al.*, 2013, in press). The taxonomic diversity recorded in this area includes one gavialoid (*Gryposuchus neogaeus* Burmeister) and at least five *Caiman* species (including *C. latirostris*; Bona *et al.*, in press).

Conversely, the Miocene crocodylian record in other Argentine localities is sparse and always fragmentary. Starck and Vergani (1996) and Starck and Anzótegui (2001) reported a mandibular fragment tentatively referred to *Caiman* sp. found in the Palo Pintado Formation in the Calchaquí Valley, Salta Province (Fig. 1). Such a generic determination, although correct, was not supported by a comparative morphological study. This record from outside the Paraná area broadens the southern geographic distribution of late Miocene caimanines. Knowledge of the caimanine fossil record during the Neogene

at these latitudes is essential not only to understand the evolutionary and biogeographic history of *Caiman*, but also to provide an empirical basis for the interpretation of the evolution of natural environments in northern Argentina. This enhances the significance of this material—collected with precise stratigraphic control—enough to justify its study.

This work is based not only on a detailed anatomical comparative analysis of this fossil material of *Caiman* (Fig. 2), but also on sedimentological and paleontological observations of the exposures of the Palo Pintado Formation in the Quebrada de Salta (Fig. 3). Hence, one of the goals of this study is a paleoenvironmental interpretation based on the architectural-element level and the bounding-surface hierarchy and its relationship with the fossil crocodile specimen which, subsequently, suggest a possible scenario in which this caimanine may have lived.

GEOLOGICAL AND CHRONOSTRATIGRAPHIC SETTING

The studied material was collected from upper beds of the Palo Pintado Formation (Mauri, 1948; Díaz and Malizzia, 1983) exposed in the Quebrada de Salta in the southern Calchaquí Valley (between 25°41'34"S–66°03'31"W and 25°37'17"S–66°03'26"W; Fig. 1). This unit is part of the Payogastilla Group (Díaz and Malizzia, 1983), which comprises (from base to top) the Los Colorados, Angastaco, Palo Pintado and San Felipe formations. The measured thickness of this group exceeds 6 km in the area and was deposited in a Cenozoic (middle Eocene–Pliocene) foreland basin related to the Andean Orogeny (Starck and Vergani, 1996). Part of this foreland basin developed over the previous Salta Group Rift Basin (Cretaceous–Paleogene).

The Palo Pintado Formation comprises fluvial system deposits that are well exposed in the outcrops at Quebrada de Salta. Lithofacies boundaries and the characteristics and geometries of the architectural elements are documented in the stratigraphic column (Fig. 1) and in the photo mosaic of the outcrop (Fig. 3). The principles on bounding surface hierarchies of Miall (1985, 1994, 1995 and 2006) were followed for architectural analysis, with minor modifications to adapt these principles to the studied deposits. Lithofacies descriptions and their interpretation are summarized in Appendix 1 (Supplementary Online Information).

The Palo Pintado Formation was deposited between 10

Ma (K/Ar) (Galli *et al.*, 2008) and 5.2 Ma (Coutand *et al.*, 2006; Bywater Reyes *et al.*, 2010) and is mainly composed of upward fining and thinning cycles. Each cycle involves matrix

supported conglomerates, clean quartzitic and sublithic sandstones and is topped by green shale levels. These cycles have been interpreted as the record of rivers intermediate between

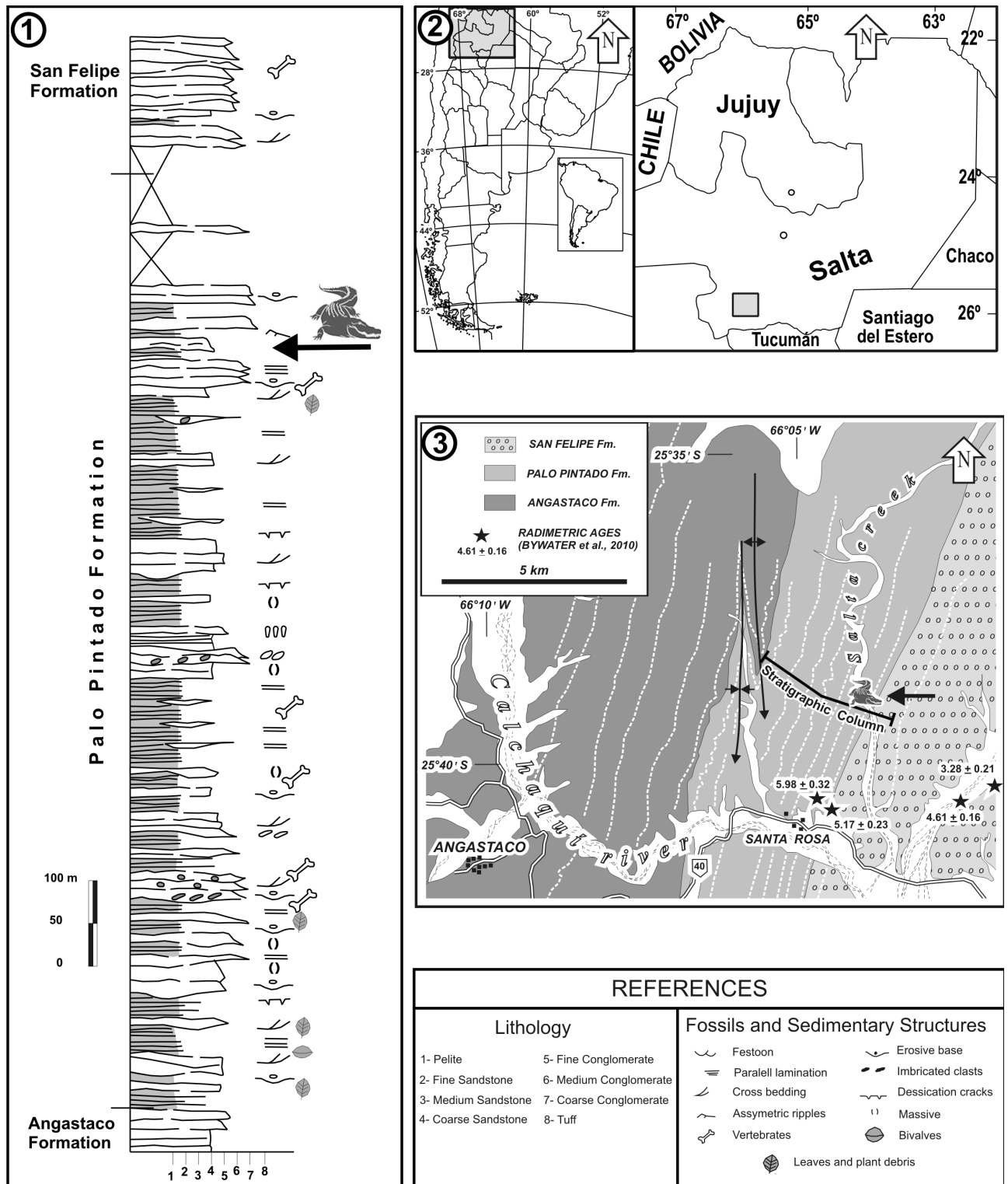


Figure 1. 1. Palo Pintado stratigraphic column at the Quebrada Salta section (location in 1.3) with the fossil stratigraphic location. 2. Geographic location of the study area (southern Calchaquí valley). 3. Geological map of the study area with the stratigraphic column and fossil location. Radiometric ages (Tuff zircons U-Pb) in Ma.

low sinuosity, multichannel rivers and others with highly sinuous and simple channels in a sandy-gravel wandering fluvial system (Galli *et al.*, 2011a).

The Palo Pintado Formation measured 1100 m at the surveyed section. The unit overlies transitionally the coarser Angastaco Formation and grades upward into the San Felipe Formation (fluvial and alluvial). The Palo Pintado and San Felipe formations form a continuous coarsening- and thickening-upward megasequence several kilometers thick (Starck and

Vergani, 1996).

The lower and middle sections of the Palo Pintado Formation contain a highly diverse fossil assemblage: well-preserved leaves (Herbst *et al.*, 1987; Anzótegui, 1998, 2006), fructifications (Anzótegui *et al.*, 2007), logs (Lutz and Martínez, 2007), palynomorphs (Anzótegui and Cuadrado, 1996; Acevedo *et al.*, 1997; Mautino, 2007), bivalves (Morton, 1992; Herbst *et al.*, 2000) and vertebrates (Díaz *et al.*, 1987, 1989; Starck and Vergani, 1996; Starck and Anzótegui, 2001).

MATERIALS AND METHODS

For the comparative study of specimen MLP 89-XII-5-1, a sample of 149 modern alligatorid specimens from herpetological collections was studied. The material is housed in the Museo de La Plata, La Plata, Argentina (MLP); Museo Argentino de Ciencias Naturales “Bernardino Rivadavia”, Buenos Aires, Argentina (MACN) and Zoologische Staatssammlung, Munich, Germany (ZSM). The sample included 25 specimens of *C. latirostris*, 74 of *C. yacare*, three of *P. trigonatus* (Schneider), five of *P. palpebrosus* Cuvier, 30 of *Melanosuchus niger*, one of *Alligator sinensis* Fauvel and one of *A. mississippiensis*. Additional specimens of Miocene caimans used as comparative material were those figured by Langston (1965), Brochu (1999), Medina (1976), Sánchez Villagra and Aguilera (2006), Gasparini (1981) and also specimens under study (Bona *et al.*, 2013, in press) housed in the MLP and MACN.

SYSTEMATIC PALEONTOLOGY

CROCODYLIA Gmelin, 1789 *sensu* Benton and Clark, 1988

ALLIGATORIDAE Cuvier, 1807 *sensu* Norell *et al.*, 1994

CAIMANINAE Brochu, 2003 (according to Norell, 1988)

Caiman Spix, 1825

Type species. *Caiman fissipes* Spix, 1825; original designation, central and southern South America, Recent.

Caiman cf. latirostris (Daudin, 1802)

Figure 2

Additional material. MLP 89-XII-5-1, fragment of left mandibular ramus, with most of the dentary, teeth 4th and 12th complete and implanted, and fragments of the 6th through 9th teeth (Fig. 2.1, 3, 4).

Geographic and stratigraphic provenance. Quebrada Salta, southern area of the Calchaquí Valley: 25°39'04"S–66°04'00"W, Salta Province. Upper levels of the Palo Pintado Formation (Díaz and Malizzia, 1983); late Miocene (Fig. 1).

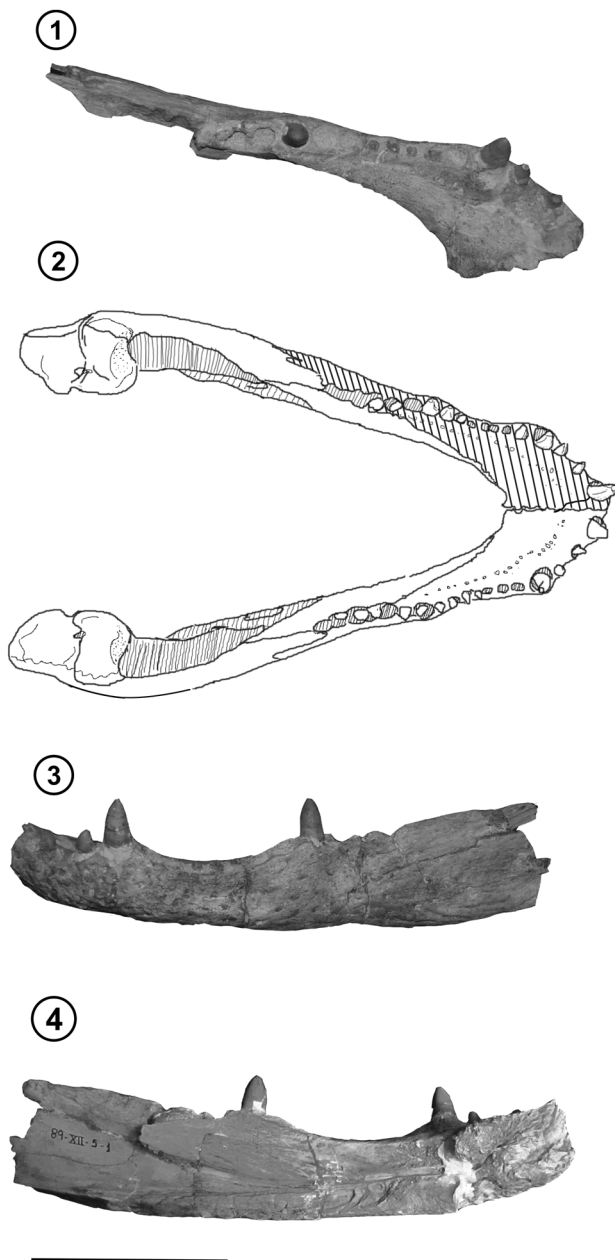


Figure 2. 1, 3, 4, *Caiman cf. latirostris*, MLP 89-XII-5-1; 1, dorsal view; 3, lateral view; 4, medial view. 2, *C. latirostris*, extant specimen, mandible in dorsal view. Scale bar = 10 cm.

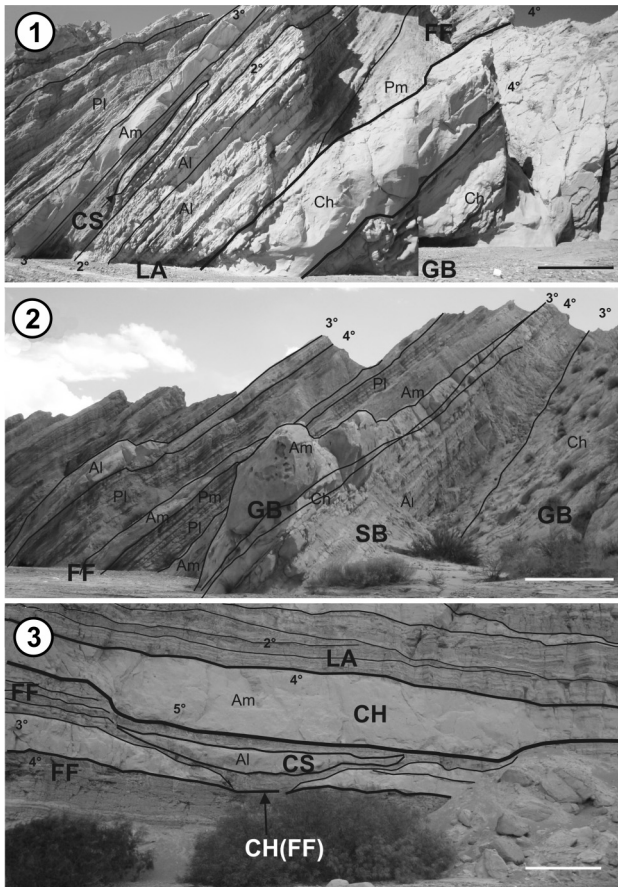


Figure 3. 1, Channel-fill deposits are dominated by lithofacies Ch with sand lithofacies Am forming the upper part of the bar deposits, with fining-upward successions (GB) and the architectural-element with lateral-accretion (LA), scale bar= 2m. 2, Facies showing a sequence of channel facies (GB) and floodplain fines with crevasse channel (FF), scale bar= 2m. 3, abandoned channel fills (element FF-CH-), scale bar= 4m.

Description

Dentary. Except for a small posterior fragment, the dentary is complete. It is similar in morphology to other specimens of *Caiman* with well marked curves of the dorsal margin (in lateral view) between the alveoli 1 and 4–12 (Fig. 2.3,4). Hence, as in the other adult caimans (e.g., *C. latirostris*, *C. yacare*, *C. crocodilus*), the dimensional heterodonty is marked: the largest dentary teeth are the 1st, 4th, 11th and 12th. These are implanted at the highest points of this bone. Alveoli 1 and 4 are high, at the same level, and as in other caimanines the anterior part of the dentary (at the level of alveolus 4) is as high as the middle part (at the level of alveolus 12). As in other caimans, the ornamentation of the lateral surface is given by small holes grouped on the symphyseal and latero-dorsal region, and continues with thin dichotomized furrows extended mainly over the ventral surface of the bone.

In dorsal view (Fig. 2.1) the mandibular symphysis extends posteriorly at the level of the posterior margin of alveolus 5 and forms a straight angle with the medial margin of the dentary. The latter curves posterolaterally at alveolus 6 forming a 135°–157° angle with the symphysis.

This adult specimen is similar in size and robustness to large specimens of extant *C. latirostris* (i.e., MACN V 1420, Bona and Desojo, 2001; Fig. 2). The dentary is 22 cm long from the anterior end of the bone to the posterior margin of the last alveolus. The proportions of the mandibular symphysis area are 6.7 cm length and width (measured on the sagittal axis of the specimen and at a 90° angle with the mandibular symphysis).

Splenial. This bone is not preserved, but the scar of its contact with the dentary is exposed in medial view (Fig. 2.4). It was excluded from the mandibular symphysis and it extended anteriorly through a projection dorsal to the Meckelian furrow. This condition is common to all Caimaninae and a synapomorphy of the clade (Brochu, 1999).

Dentition. Eighteen dentary alveoli are preserved in MLP 89-XII-5-1. The last four are partially preserved because of the lack of splenial, which in complete specimens bounds the medial part of the last alveoli. However, it can be inferred that these converge as in other caimans. The alveolar pattern of labiolingual diameter and space between alveoli is similar to that of *Caiman*: the width of the first alveolus, partially preserved, is similar to that of alveolus 4; alveoli 2 and 3 are small and well spaced, from 5 to 9 they decrease in size and then they increase from 10 to 12.

The best preserved teeth (d2, d3, d4, d12) are similar in morphology to those of other caimans. They are conical, with smooth vertical striae and an anteroposterior carina dividing the tooth into two sub-equal surfaces, “labial” or lateral, and lingual or medial.

Comparisons

In MLP 89-XII-5-1 the anterior projection of the splenial is dorsal to the Meckelian furrow as in members of Caimaninae. Unlike *Eocaiman* Simpson, and the large caimanines such as *Mourasuchus* Price and *Purussaurus* Barbosa Rodriguez, it has two marked wavelets in lateral view; hence, alveoli 1, 4 and 12 are almost at the same height. Unlike *Paleosuchus* Gray, the alveoli are not compressed laterally. In MLP 89-XII-5-1 the ornamentation of the lateral surface of the dentary and the

alveolar pattern are similar to those of *Melanosuchus* Spix and *Caiman*. As in these genera, the symphysis extends up to the posterior margin of alveolus 5 (in *C. latirostris* the length of the symphysis varies among species reaching up to the 5 or 5–6 alveoli; Bona and Desojo, 2011). Both the angle between the symphysis and the medial margin of the dentary, and the outline of the lower mandible show that the dentary is from an animal with a broad snout, similar to *C. latirostris* (Fig. 2.1–2). As mentioned above, in MLP-89-XII-5-1 the medial margin (inner) of the dentary in dorsal view is straight in relation to the symphysis and then it curves postero-laterally forming an angle of approximately 157°. This condition is different from that in other species of living caimans such as *Caiman yacare* and *Melanosuchus niger*, in which the internal margin of the dentary is postero-laterally oriented without curving, forming an angle of approximately 170° with the mandibular symphysis.

DISCUSSION

Caiman has not been adequately defined from an osteological point of view. According to phylogenetic hypotheses based on morphological data (Brochu, 1999, 2011; Bona et al., 2013, in press), *Caiman* and *Melanosuchus* form a clade supported by the presence of prominent rostral crests on the snout, by the morphology of the fronto-parietal suture, and by the arrangement of scales on the ventral side of the neck (Brochu, 1999). Nevertheless, molecular analyses generally support monophyly of *Caiman* exclusive of *Melanosuchus*.

In a phylogenetic analysis of the taxa represented in the Miocene of Paraná, Bona et al. (2013, in press) proposed that the relative size of the temporal fenestrae and the orbits, and the pattern of the alveoli and premaxillary-maxillary diastema (and possibly its counterpart in the dentary), could be potentially diagnostic characters of *Caiman*. In this context, these authors proposed the synonymy of *Melanosuchus* and *Caiman* (Norell, 1988; Poe, 1997). *Caiman* would have differentiated at least in the middle Miocene of South America (Bona and Paulina Carabajal, 2013), being the late Miocene the time of greatest morphological and taxonomical diversification. This diversification includes the differentiation of the modern species of this genus (e.g., *Caiman latirostris* and probably *C. yacare*, Bona et al., 2013, in press), the oldest record of which is documented in the area of Paraná, in levels assigned to the late Miocene (“Conglomerado osífero”; Cione

et al., 2000; Brandoni and Scillato-Yané, 2007; Brandoni, 2011).

Detailed morphological study of specimen MLP 89-XII-5-1 allows its assignment to the Caimaninae *Caiman*, and the similar outline of the symphyseal region allows considering it as *Caiman* cf. *latirostris* (Fig. 2.1, 2). The record of *Caiman* cf. *latirostris*, now verified for the Miocene of northwestern Argentina, demonstrates a wider distribution of this taxon during this times interval.

As mentioned above, specimen MLP 89-XII-5-1 was collected from the upper Palo Pintado Formation, close to its transition into the overlying San Felipe Formation. Both units yielded U-Pb zircon ages from their abundant intercalated piroclastic levels (Coutand et al., 2006; Bywater Reyes et al., 2010). One of these available ages (5.98 ± 0.32 Ma; Bywater Reyes et al., 2010) came from a level slightly lower than the *Caiman* bearing bed; thus the specimen should be slightly younger than 6 ma.

Paleoenvironmental and paleogeographic considerations

According to the architectural element characteristics of the Palo Pintado Formation and its associated lithofacies, this unit represents a fluvial system with deposits developed in both channels and overbanks. The intrachannel record is composed of gravel bars (GB) and bedform deposits related to downstream migration and vertical accretion of transverse bars and sand waves (SB). On the other side, the overbank deposits are represented by lateral accretion bars (LA) on the floodplain, small channels carved in the main channel edges during flooding stages interpreted as crevasse channels (CS), and a floodplain (FF) (Fig. 3).

The overall architectural-element characteristics are akin to a sinuous gravel fluvial system with intermediate to high sinuosity (Gravel-bed wandering river; Miall, 1985).

Compared to the Miall's model (Miall, 1985), the Palo Pintado Formation deposits show a dominance of the SB element. Based on this fact, a sandy-gravel wandering fluvial system denomination has been proposed for them (Galli et al., 2011a).

Sedimentologic analysis and fossil content suggest relatively humid climate during deposition of the Palo Pintado Formation. Preliminary X-ray diffraction data from floodplain clay minerals revealed the presence of illite, montmorillonite, magnesium-rich smectite and kaolinite generated by hydroly-

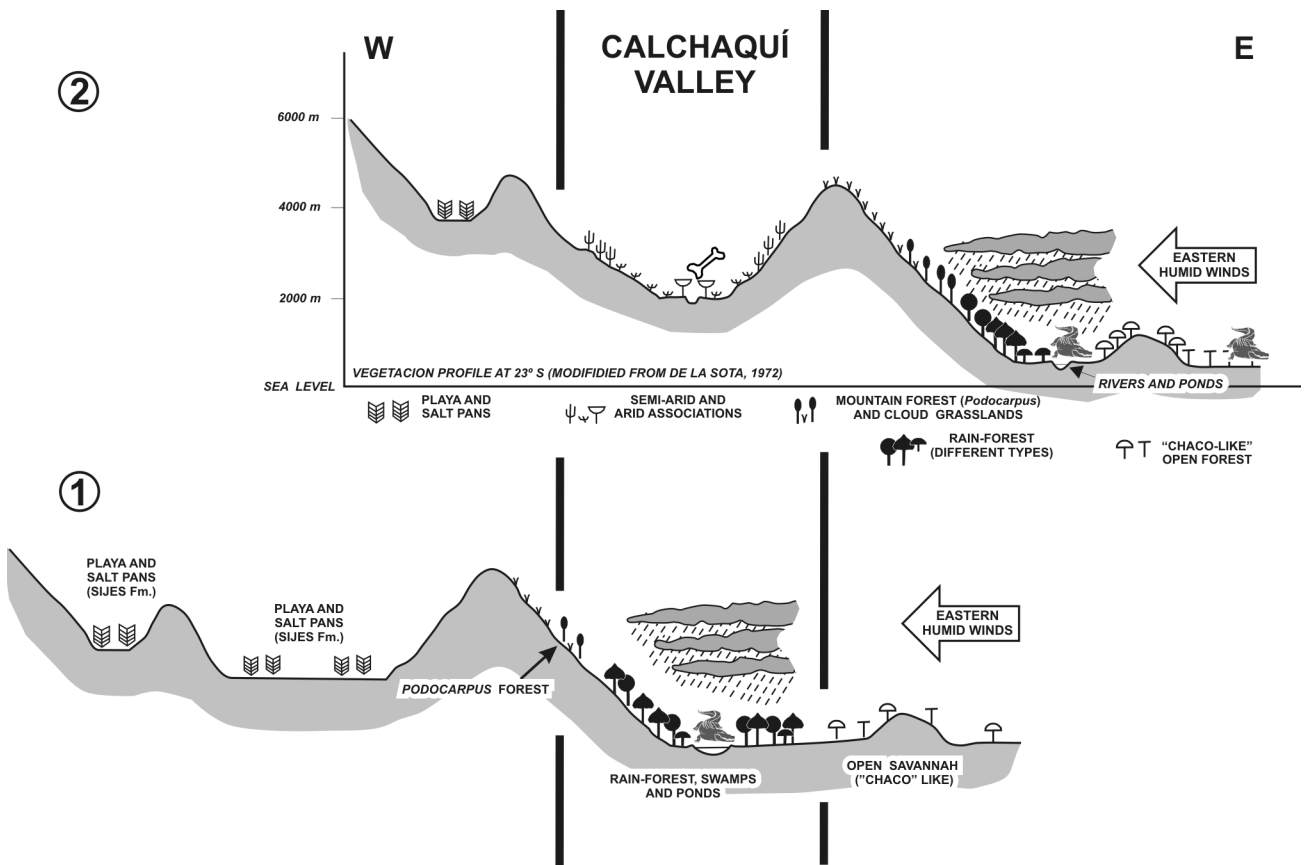


Figure 4. Schematic topographic profiles showing the paleoenvironment distribution (modified from Starck and Anzótegui, 2001). 1. Situation at 6 Ma (deposition time for the upper section of the Palo Pintado Fm). 2. Present day situation.

sis under a warm and humid climate (Galli *et al.*, 2011a,b). The formation of these clay minerals during the deposition of the Palo Pintado Formation is another evidence of benign climatic conditions, compared with the mostly arid ones previous to 10 Ma (during the deposition of the Angastaco Formation, Starck and Vergani, 1996; Starck and Anzótegui, 2001).

In this paleoenvironmental scenario, the studied fossil was found in cross-laminated sandstone beds (At) representing the transverse bedforms setting (SB). The occurrence of *Caiman cf. latirostris* also supports the hypothesis that the climatic conditions in Valle Calchaquí during the late Miocene were comparatively more benign than those inferred for contemporaneous units deposited to the east (Guanaco Formation, Orán Group).

From a more regional perspective, the Palo Pintado Formation was deposited at the front of a contemporaneous orogenic belt in an overall setting similar to the present-day Chaco Plains, specially their western sectors close to the Andean foothills (Starck and Vergani, 1996; Starck and Anzótegui,

2001). Paleoenvironment dynamics in northwestern Argentina during the late Neogene was strongly controlled by the advance of the Andean Orogen front. The eastward thrust belt propagation pushed in the same direction the humid environments developed at the bottom of the foothills (Fig. 4). To the west, behind the orogenic front, intermontane valleys were coevally originated. These were higher and under dryer climatic conditions (*e.g.*, the present-day Valle Calchaquí). Thus, during the last 6 Ma (since the late Miocene), the orogenic front propagated approximately 150 km to the east (at a rate of about 2.5 cm/year) pushing the related climatic belts. The distribution of *Caiman cf. latirostris* seems to accompany that displacement; thus, the current geographic distribution of the species can be explained in terms of a retraction process related to climatic changes controlled by orogenic causes.

Caiman latirostris is currently found in Paraguay, southern and eastern Brazil, much of Bolivia (“Amazonas Boliviano”) and northern and northeastern Argentina. In Argentina, this species extends further south than other species of the genus (Larriera and Imhof, 2000) and has been recorded in the

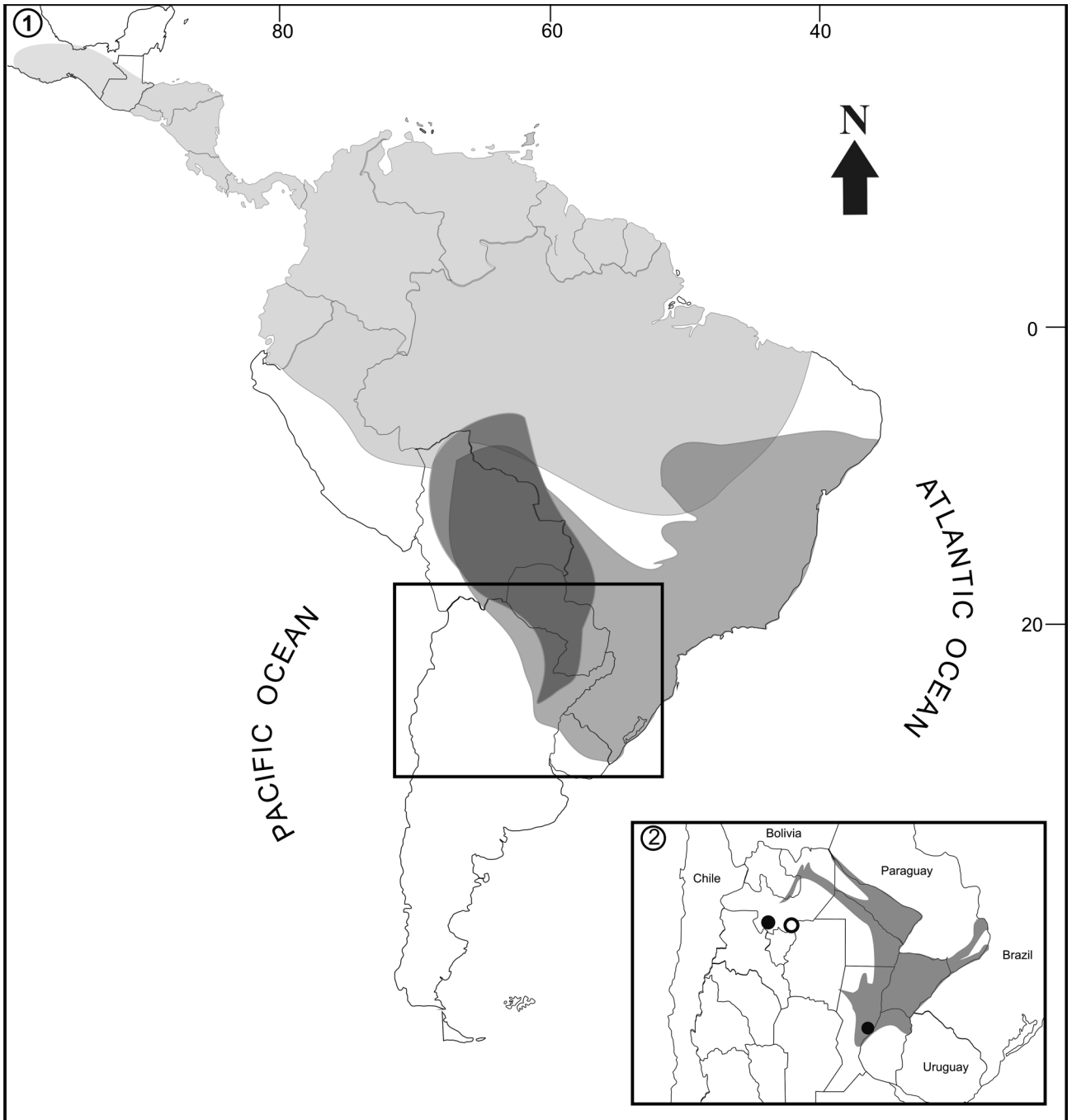


Figure 5. 1. Geographic distribution of *Caiman* genus in South America; the three areas in gray, from light gray to dark gray, indicate the present distribution of *C. crocodilus*, *C. latirostris* and *C. yacare*, (respectively) **2.** Detail of geographic distribution of *Caiman latirostris* in Argentina; gray area indicates the present distribution of the species; black circles indicate the distribution of *Caiman* cf. *latirostris* in the upper Miocene (left) and white circle indicates the distribution of *C. latirostris* in the Pliocene.

provinces of Chaco, Corrientes, Formosa, Santa Fe, Entre Ríos, Misiones, Salta and Jujuy (Fig. 5). In Salta Province, living populations of *Caiman latirostris* are restricted to the northeast (basins of the Bermejo and Pilcomayo rivers). It inhabits environments with abundant vegetation and lentic water bodies such as the pantanal and floodplains of rivers and

creeks developed under a subtropical to warm-temperate and humid climate (Fig. 3; Freiberg and Carvalho, 1965; Cei, 1993 Piña *et al.*, 2004; Rueda-Almonacid *et al.*, 2007; Barrios, 2013). These environmental and climatic conditions were those probably existing between 10 and 5 Ma in the Valle Calchaquí.

CONCLUSIONS

This study confirms the occurrence of *Caiman cf. latirostris* in the late Miocene in northwestern Argentina. It also extends the paleogeographic distribution of the genus and explains the presence of this taxon in younger rocks in eastern locations within Salta Province (e.g., Rosario de la Frontera, Pliocene–early Pleistocene and Aguas Blancas, Pleistocene; Paterson, 1936; Barrios, 2013). Thus, the idea that the biogeographic history of *Caiman* in the area was a story of expansion and (mainly) retraction clearly appears (in view of the environmental and climatic changes occurred during the late Neogene) linked to tectonic events (see Starck and Anzótegui, 2001).

Sedimentological and facial analyses carried out in the area (Quebrada Salta) suggest a preliminary hypothesis on the Palo Pintado Formation depositional subenvironments (Supplementary Online Information: Appendices 1 and 2). For the time-interval analyzed, the basin infill was deposited by an intermediate-sinuosity, sand-gravel wandering fluvial system. The sinuosity reflects low regional topographic gradients. Other outstanding characteristics of the studied sequences are the abundance of At lithofacies, several channel styles often associated to swamps and the development of grey paleosoils, all indicative of fluvial systems developed under tropical climates (Miall, 2006). These climatic conditions lasted in the Valle Calchaquí area until at least 5 Ma, allowing the development of a varied flora and fauna, including the studied *Caiman cf. latirostris*.

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