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Spatial patterns in abundance, taxonomic composition and carbon biomass of nano- and microphytoplankton in Subarctic and Arctic Seas

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ABSTRACT

In the summers of 2007 and 2008, we studied assemblages of nano- and microphytoplankton from the subsurface chlorophyll maximum (SCM) across five broad oceanographic domains in the seas surrounding northern North America. These domains are the eastern Subarctic North Pacific (ESNP), Bering and Chukchi Seas (BE-CH), Beaufort Sea and Canada Basin (BS-CB), Canadian Arctic Archipelago (CAA), and Baffin Bay and Labrador Sea (BB-LS). Average abundance and total carbon biomass (C) of phytoplankton (> 2 µm) varied ∼10-fold and ∼20-fold, respectively, across the five domains. In the BE-CH, CAA and BB-LS, diatoms averaged 35–70% and dinoflagellates 11-45% of total phytoplankton C (> 2 µm), whereas in the ESNP and BS-CB, unidentified flagellates/coccoids (2–8 µm) represented a greater proportion of total C (27% and 39% respectively) than in the other domains.

In the BE-CH and BB-LS, phytoplankton C ($> 2 \mu m$) was dominated by dinoflagellates of the genus Gymnodinium, centric diatoms including Thalassiosira spp. and Chaetoceros spp., unidentified flagellates/coccoids (2–8 µm), and cryptomonads. In contrast, diatoms such as Thalassiosira spp. and its resting spores dominated C in the CAA, with dinoflagellates being less significant than in the BE-CH and BB-LS. Unidentified flagellates/coccoids (2–8 µm), Gymnodinium spp., and cryptomonads dominated in the ESNP, and particularly in the BS-CB, where diatoms contributed only 18% of the very low levels of total phytoplankton C ($> 2 \mu m$).

Phytoplankton C (> 2 µm) to chlorophyll a ratios (phyto C:chl a) averaged only 31 g C g chl a^{-1} in the oligotrophic BS-CB domain, and 51–150 g C g chl a^{-1} in the other domains, whereas ratios of biogenic silica to phytoplankton C (> 2 µm) (bSi:phyto C) were lowest in the eastern domains. Estimates of phytoplankton C were highly sensitive to the choice of C to cell volume equations (C:vol) adopted in the calculations, particularly in diatom-rich areas.

This study highlights how diatoms and dinoflagellates are the main drivers of large-scale variations in C biomass for phytoplankton ($> 2 \mu m$), whereas unidentified flagellates/coccoids (2–8 μ m) make a significant contribution to C biomass in oligotrophic domains, such as BS-CB, where diatoms and dinoflagellates are less abundant. Reduced surface water density (σ_T) was associated with deeper SCM layers, and with decreased C biomass of unidentified flagellates/coccoids (2–8 µm). These observations confirm recent studies highlighting the role of surface water stratification caused by melting sea ice in shaping nano- and microphytoplankton assemblages.

1. Introduction

Recent enhanced melting of sea ice has exposed greater areas of the Arctic Ocean to wind stress, irradiance and freshening of surface waters, resulting in increased coastal productivity in some areas [\(Tremblay](#page-27-0) [et al., 2011\)](#page-27-0), and latitudinal shifts of biological boundaries in others ([Grebmeier et al., 2006\)](#page-26-0). In the stratified southeast Beaufort Sea, seasonal nitrate $(NO₃⁻)$ consumption and net community production (NCP)

increased between 2003 and 2011, but decreased in the North Water Polynya in northern Baffin Bay between 1997 and 2011 as a result of increasing surface freshening and stratification ([Bergeron and](#page-25-0) [Tremblay, 2014\)](#page-25-0). In the Canada Basin, a gradual deepening of the subsurface chlorophyll maximum (SCM) has led to a shift in the size composition of the phytoplankton assemblages ([Li et al., 2009](#page-26-1)). With these rapidly changing oceanographic conditions in the Arctic and Subarctic, there is a pressing need to understand how large-scale

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Fig. 1. Study areas in Arctic and Subarctic Seas around northern North America showing the sampling stations during summers of 2007 (eastern sector and Canada Basin; C3O/JOIS cruises) and 2008 (western sector; C3O cruise). See Section [2.1](#page-1-1) and [Tables 1 and 2](#page-2-0) for details of regions sampled during each cruise. White circles represent full-profile sampling stations where phytoplankton samples were taken from the depth of the subsurface chlorophyll maximum (SCM). Grey circles represent surface stations where phytoplankton samples were taken from depths of 2–10 m. For further cruise details see [Carmack and McLaughlin \(2011\), Varela et al.](#page-26-2) [\(2013\) and Wyatt et al. \(2013\)](#page-26-2). The oceanographic domains consist of (1) the Eastern Subarctic North Pacific Ocean (ESNP), (2) the Bering and Chukchi Seas (BE-CH), (3) the Beaufort Sea and Canada Basin (BS-CB), (4) the Canadian Arctic Archipelago (CAA), and (5) Baffin Bay and Labrador Sea (BB-LS). Station markers shown outside of these domains represent excluded stations (see [Table 1\)](#page-2-0).

physical and chemical changes will shift biological communities in the future. The Arctic has traditionally been a challenging environment to study but improvements in technology, ship access and availability of historical datasets are promoting more synoptic oceanographic studies on physical processes (e.g. [Carmack and McLaughlin, 2011\)](#page-26-2), nutrient cycling (e.g. [Codispoti et al., 2013](#page-26-3)), primary productivity (e.g. [Ardyna](#page-25-1) [et al., 2011; Hill et al., 2013; Matrai et al., 2013; Varela et al., 2013; Hill](#page-25-1) [et al., 2017](#page-25-1)) and the composition of suspended biogenic particles (e.g. [Wyatt et al., 2013; Crawford et al., 2015](#page-27-1)).

Limited light penetration and lack of wind action due to sea-ice cover in the deep Arctic basins result in an upper water column characterized by low irradiance and low nutrient concentrations, with phytoplankton assemblages consequently limited in biomass and productivity (e.g. [Varela et al., 2013](#page-27-2)). Under these oligotrophic conditions, phytoplankton often aggregate within a distinct SCM where nutrient availability is optimized at light levels still adequate for photosynthesis ([Martin et al., 2012\)](#page-26-4). Seasonal thinning and melting of sea ice can promote phytoplankton productivity by exposing a greater sea surface area to irradiance and wind forcing over shallow shelves ([Tremblay](#page-27-0) [et al., 2011\)](#page-27-0), whereas in deeper basins melting can reduce productivity by enhancing the strong density stratification that inhibits vertical supply of nutrients ([McLaughlin and Carmack, 2010\)](#page-26-5). The Arctic Ocean is also an advective estuarine environment [\(Carmack and Wassmann,](#page-26-6) [2006; Yamamoto-Kawai et al., 2006](#page-26-6)) where chemical and biological changes in one area can transform the lateral supply of nutrients to other areas (e.g. [Tremblay et al., 2002](#page-26-7)).

Research on pelagic and sea-ice unicellular autotrophs in Arctic marine waters extends back to the late 1800s (e.g. see review by [Poulin](#page-26-8) [et al., 2011](#page-26-8)). For much of the 20th century descriptive studies focused primarily on taxonomic examinations of larger cells that could be readily collected and preserved ([Poulin et al., 2011\)](#page-26-8), but later studies began to address how phytoplankton assemblages are shaped by physical and chemical processes in Arctic and Subarctic waters (e.g. [Booth,](#page-25-2) [1988; Booth and Horner, 1997; Booth and Smith, 1997\)](#page-25-2). Analyses of the composition of phytoplankton assemblages are subject to particular drawbacks that hamper large-scale studies. Painstaking microscopic examination is highly skilled and labour intensive, limiting the sampling resolution achievable in oceanographic surveys. Changing taxonomic nomenclature can also potentially complicate comparisons with historical studies. The methodology employed for sampling and preserving phytoplankton has resulted in a historical bias towards microphytoplankton (20–200 µm), although there is now an increasing interest in the role of nanophytoplankton $(2-20 \mu m)$ and picophytoplankton (0.2–2 µm) in the Arctic Ocean (e.g. [Li et al., 2009;](#page-26-1) [Tremblay et al., 2009; Ardyna et al., 2011\)](#page-26-1), particularly in open

oligotrophic regions where heterotrophic protists are also clearly important (e.g. [Sherr et al., 1997; Kirchman et al., 2009\)](#page-26-9). Separation of autotrophic carbon (C) from heterotrophic and detrital C is challenging using light microscopy ([Booth, 1987](#page-25-3)) and consequently, other methods have been increasingly used to evaluate the importance of smaller autotrophs, such as fluorescence microscopy [\(Booth, 1987](#page-25-3)), chemotaxonomic analyses of pigments ([Vidussi et al., 2004; Hill et al., 2005;](#page-27-3) [Tremblay et al., 2009; Coupel et al., 2012; Coupel et al., 2015](#page-27-3)) and flow cytometry [\(Li et al., 2009; Tremblay et al., 2009; Ardyna et al., 2011;](#page-26-1) [Balzano et al., 2012\)](#page-26-1). Molecular techniques such as 18S rRNA gene are also being used on both autotrophic and heterotrophic protists [\(Lovejoy](#page-26-10) [et al., 2011; Lovejoy and Potvin, 2011](#page-26-10)). The presence of sea-ice unicellular autotrophs in close proximity to phytoplankton assemblages (e.g. [Gosselin et al., 1997](#page-26-11)), and the aggregation of phytoplankton within discrete SCM layers are also potential complications in the accuracy of sampling and enumeration of phytoplankton assemblages in the vertical dimension. Despite these drawbacks, some large-scale surveys of Arctic phytoplankton assemblages are now emerging, both in the form of extensive literature reviews (e.g. [Poulin et al., 2011\)](#page-26-8) and long oceanographic transects [\(Gosselin et al., 1997; Tremblay et al.,](#page-26-11) [2009; Ardyna et al., 2011\)](#page-26-11).

The International Polar Year (IPY) Canada Three Oceans (C3O) project undertaken during the summers of 2007 and 2008, combined with the ongoing Joint Ocean Ice Study (JOIS) project, provided an opportunity to compare phytoplankton assemblages across the Arctic and Subarctic Seas surrounding northern North America. In the present study, we used the domain classification of [Carmack and McLaughlin](#page-26-2) [\(2011\),](#page-26-2) which was already applied to contrast large-scale variations in physical oceanography [\(Carmack and McLaughlin, 2011](#page-26-2)), primary productivity and nutrient uptake [\(Varela et al., 2013\)](#page-27-2), and composition of biogenic particles ([Wyatt et al., 2013; Crawford et al., 2015](#page-27-1)). This investigation complements these previous studies by comparing the abundance, C biomass and taxonomy of nano- and microphytoplankton within the SCM among domains during the summer period.

2. Methods

2.1. Study area and domain classification

Seawater samples were collected across the northwest Atlantic, Arctic, and northeast Pacific Oceans [\(Fig. 1\)](#page-1-0). The sampling stations were arranged into five broad-scale regional domains that coincide with distinct water masses revealed by temperature/salinity (T/S) correlations ([Carmack and McLaughlin, 2011\)](#page-26-2). These domains ([Fig. 1](#page-1-0), [Table 1\)](#page-2-0) consist of the Eastern Subarctic North Pacific Ocean (ESNP), the Bering

Table 1

Oceanographic domains and stations for cruises 2007–19, 2007–20, and 2008–02 in the summers of 2007 and 2008, operated by the Institute of Ocean Sciences, Fisheries and Oceans Canada (see [Varela et al., 2013](#page-27-2) for more details). Phytoplankton samples were taken from the subsurface chlorophyll maximum (SCM), where present, or from the depth indicated where no SCM was present. At surface only stations (*), the depth of the SCM was unknown, and samples were taken from a depth of between 2 and 10 m (∼50% surface irradiance). Excluded stations represent those that could not be included in any of the five broad regional domains defined according to oceanographic properties by [Carmack and McLaughlin \(2011\)](#page-26-2).

* Denotes surface station only, where samples were taken from 2 to 10 m depth. No associated vertical profiles or ice cover estimates were taken at these surface stations by [Varela et al.](#page-27-2) [\(2013\)](#page-27-2) or [Wyatt et al. \(2013\)](#page-27-1) and the depth of the SCM was unknown.

and Chukchi Seas (BE-CH), the South Beaufort Sea (off shelf) and Canada Basin (BS-CB), the Canadian Arctic Archipelago (CAA), and Baffin Bay and Labrador Sea (BB-LS).

We present phytoplankton data from the same stations as those described in [Varela et al. \(2013\), Wyatt et al. \(2013\) and Crawford et al.](#page-27-2) [\(2015\),](#page-27-2) with the exceptions of 5 stations that were not sampled for phytoplankton taxonomy (full-profile station RS-1 in the BS-CB domain,

and surface stations BC-4, BCL-6, BS-3, BS-1 in the BE-CH domain). In addition, several stations were excluded from the domain classification ([Table 1](#page-2-0), see also [Varela et al., 2013](#page-27-2)) because they were located in transitional zones where physical conditions were not consistent with neighbouring domains [\(Carmack and McLaughlin, 2011; Varela et al.,](#page-26-2) [2013\)](#page-26-2).

2.2. Sampling program

Sampling started in the Labrador Sea in early July 2007 aboard the CCGS Louis S. St-Laurent (C3O cruise 2007-19) and continued through Baffin Bay and the Canadian Arctic Archipelago [\(Fig. 1](#page-1-0) and [Table 1](#page-2-0); see also [Varela et al., 2013\)](#page-27-2). During the subsequent JOIS cruise (2007–20) in July–August 2007, still aboard the CCGS Louis S. St-Laurent, sampling was carried out in the southern Beaufort Sea (off shelf) and Canada Basin. The July 2008 C3O cruise (2008-02), aboard the CCGS Sir Wilfrid Laurier, sampled from the west coast of Vancouver Island (Canada), through the northeast Pacific, the Bering Sea, Chukchi Sea, and southern Beaufort Sea (off shelf).

Throughout all cruises, consistent methodologies were adopted for sampling, enumeration of phytoplankton taxa, and all calculations. Station identification and positions are presented in [Table 1,](#page-2-0) and a fuller account of background data for all stations is given in [Table 1](#page-2-0) of [Varela et al. \(2013\).](#page-27-2) Complementary physical information, and chemical and biological samples were collected at the same stations, consisting of full-profile stations (sampled from the surface to the bottom of the euphotic zone) and surface stations (only one surface depth sampled); data and methodologies are presented elsewhere ([Carmack and](#page-26-2) [McLaughlin, 2011; Varela et al., 2013; Wyatt et al., 2013; Crawford](#page-26-2) [et al., 2015](#page-26-2)). Mean properties for each domain are summarized from these studies in [Table 2](#page-3-0). The chlorophyll data used in the present study refer to total chl a $(> 0.7 \,\text{\ensuremath{\mu}m})$ as given in [Varela et al. \(2013\)](#page-27-2). Density

of seawater is presented as σ_T (kg m⁻³ – 1000). Stratification index was calculated as the difference in σ_T between the surface and the depth of 1% surface irradiance. The use of a fixed lower depth of e.g. 80 m (e.g. [Ardyna et al., 2011\)](#page-25-1) for this calculation was precluded by the shallow water encountered in some areas, such as the BE-CH. However, the index is relatively insensitive to the choice of lower depth because most of the stratification of σ_T occurs within the upper ∼30 m (see [Fig. 2](#page-4-0)).

Samples for nano- and microphytoplankton analyses from full-profile stations were collected from the depth of the SCM where present ([Table 1](#page-2-0)). When the SCM was absent, samples were taken at depths shown in [Table 1.](#page-2-0) At an additional 10 stations, full vertical profiles were not taken and "surface" only samples were collected at 2–10 m depth (within the surface mixed layer where present), which typically corresponded to the depth of 50% of incident surface irradiance. The depth of the SCM was therefore unknown at these "surface" stations; phytoplankton data derived from these stations are included in the raw counts shown in the [Appendix Tables 1](#page-20-0)–6, but are not included in station distribution data and domain averages [\(Figs. 3](#page-5-0)–7) which refer only to samples from the SCM.

2.3. Identification and enumeration of phytoplankton taxa

On the ship, water samples were taken using Niskin bottles, and subsamples were preserved with Lugol's iodine solution ([Parsons et al.,](#page-26-12) [1984\)](#page-26-12) for identification and enumeration of phytoplankton taxa. In the

Table 2

Physical, chemical and biological properties of the water column for the five oceanographic domains, summarized from data presented by [Varela et al. \(2013\)](#page-27-2). Measurements for each variable are domain averages for the "surface" (2–10 m) and for the depth of the subsurface chlorophyll maximum (SCM). Data from full-profile stations only. Standard error in parentheses.

		ESNP	BE-CH	BS-CB	CAA	BB-LS
Number of stations		9	5	10	6	8
Ice cover (%)		0	26.0(16.0)	83.0 (9.7)	50.0(21.0)	47.5 (15.8)
Depth of SCM (m)		22.11 (4.20)	17.40 (4.66)	58.10 (7.73)	37.67 (6.82)	11.50 (4.42)
Depth of 1% surface Irradiance (m)		46.50 (7.15)	26.40 (3.08)	60.90 (7.48)	35.83 (2.61)	35.71 (2.77)
Temperature (°C)	Surface	9.13(0.39)	3.76 (0.70)	1.21(1.04)	2.11(1.00)	1.76(1.11)
	SCM	7.57(0.49)	0.79(0.97)	$-0.22(0.72)$	$-1.16(0.20)$	0.87(1.10)
Salinity	Surface	32.47 (0.09)	31.26 (0.64)	24.07 (1.00)	28.79 (0.72)	31.83 (0.51)
	SCM	32.51 (0.09)	31.83 (0.52)	30.30 (1.25)	30.61 (0.76)	32.24 (0.45)
Density σ_T (kg m ⁻³)	Surface	25.11 (0.08)	24.82 (0.49)	19.20 (0.81)	22.95 (0.53)	25.39 (0.34)
	SCM	25.37 (0.11)	25.48 (0.42)	24.31 (1.04)	24.60 (0.62)	25.77 (0.28)
Stratification index (kg m ^{-3})		1.01(0.11)	1.49(0.59)	6.83(0.75)	2.08(0.26)	1.42(0.22)
chl a (μ g L ⁻¹)	Surface	0.41(0.10)	0.43(0.22)	0.05(0.01)	0.42(0.23)	0.73(0.15)
	SCM	0.48(0.11)	1.44(0.48)	0.18(0.02)	0.69(0.14)	1.03(0.10)
NO_3^- (umol L^{-1})	Surface	8.74 (2.17)	4.07(2.28)	0.27(0.09)	0.93(0.40)	1.25(0.41)
	SCM	9.76 (2.27)	6.83(2.55)	3.85(1.35)	7.19(1.52)	1.44(0.38)
$Si(OH)4$ (umol L ⁻¹)	Surface	15.61 (2.49)	16.27 (3.26)	4.05(0.89)	4.44(0.95)	3.00(0.86)
	SCM	19.59 (2.95)	25.61 (5.79)	12.44 (2.75)	14.10 (2.24)	3.65(0.92)
$PO43$ (umol L ⁻¹)	Surface	1.07(0.12)	0.93(0.26)	0.57(0.06)	0.72(0.13)	0.60(0.18)
	SCM	1.27(0.16)	1.61(0.30)	1.14(0.11)	0.77(0.09)	0.35(0.12)
NH_4 ⁺ (umol L ⁻¹)	Surface	0.52(0.22)	0.50(0.45)	0.05(0.00)	0.08(0.02)	0.14(0.05)
	SCM	0.84(0.26)	1.13(0.47)	0.10(0.05)	0.12(0.03)	0.16(0.07)
Urea (µmol N L^{-1})	Surface	0.46(0.12)	0.88(0.26)	0.65(0.12)	0.47(0.09)	0.56(0.06)
	SCM	0.58(0.19)	0.99(0.26)	0.87(0.16)	0.77(0.21)	0.60(0.07)
Dissolved inorganic carbon DIC (μ mol L ⁻¹)	Surface	2046.1 (19.4)	1983.8 (43.1)	1748.6 (29.6)	1915.6 (13.6)	2015.2 (32.3)
	SCM	2057.6 (10.8)	2073.5 (46.0)	2045.0 (44.3)	2054.3 (38.7)	2042.8 (30.7)

DIC samples were analyzed and provided by the Institute of Ocean Sciences, Fisheries and Oceans Canada (see [Varela et al., 2013](#page-27-2)).

Fig. 2. Oceanographic profiles of salinity, temperature, density and chlorophyll a (chl a). Data source is [Carmack and McLaughlin \(2011\) and Varela et al. \(2013\).](#page-26-2) Each panel represents a scatterplot of all sampling depths for all stations within each domain. Vertical dashed red lines represent 0°C. Horizontal dashed green lines represent all the depths of the subsurface chlorophyll maximum (SCM) in each domain where samples were taken for phytoplankton taxonomical enumeration. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

laboratory, 100 ml subsamples were settled for 48 h in sedimentation chambers [\(Utermöhl, 1958](#page-27-4)) and phytoplankton cells enumerated at $600 \times$ magnification with light microscopy (LM) using an Olympus IX-70 inverted microscope. Differentiation between autotrophic and heterotrophic cells was based upon the presence/absence of chloroplasts. We endeavored to count only autotrophic nano- and microphytoplankton, but given the limitations of LM we cannot guarantee that this was achieved in all cases. At least 300 individuals, all of them with full protoplasm, were counted in random fields, and cell abundances for each taxon were expressed as numbers of cells per litre of seawater. We classified phytoplankton groups based on the taxonomic Class level (-phyceae) to which they belong (as in [Kubiszyn et al.,](#page-26-13) [2017\)](#page-26-13). Cells were identified to genus/species level, where possible, generally according to [Tomas \(1997\) and Horner \(2002\).](#page-26-14) Since small naked mono- or biflagellated cells can be difficult to identify and many can lose their flagella after addition of fixatives, they were grouped together into two unidentified flagellate/coccoids categories (2–5 µm and 5–8 µm), which could therefore include a number of taxonomic groups. Total abundance of diatoms (Bacillariophyceae), dinoflagellates (Dinophyceae), cryptomonads (Cryptophyceae) and unidentified fla g ellates/coccoids (2–8 μ m) were then calculated. When identifiable, abundances of species/genera for other taxonomic groups were also noted, such as prasinophyceans (Prasinophyceae), chlorophyceans (Chlorophyceae), euglenophyceans (Euglenophyceae), chrysophyceans

(Chrysophyceae) and silicoflagellates (Dictyochophyceae). Picophytoplankton $(< 2 \mu m)$ such as *Synechococcus* were not generally counted although some picophytoplankton of around 2 μm may have been included in our counts when staining of the cells was irregular, denoting the presence of plastids. The diatom Chaetoceros gelidus can be easily misidentified as Chaetoceros socialis. Our identification was based upon a review of the recent literature concerning the morphological variation within the C. socialis complex ([Degerlund et al., 2012; Gaonkar et al.,](#page-26-15) [2017\)](#page-26-15) and the distinguishing features of C. gelidus ([Chamnansinp et al.,](#page-26-16) [2013\)](#page-26-16).

A scanning electron microscope (SEM) was also used for a more detailed examination and identification of certain taxa, principally diatoms. Samples were cleaned of organic matter according to the method of [Hasle and Fryxell \(1970\),](#page-26-17) prepared on stubs and coated with gold-palladium according to [Ferrario et al. \(1995\)](#page-26-18). SEM examination was conducted with a Hitachi S-3500N SEM in the Department of Biology at the University of Victoria, Canada.

2.4. Calculation of cellular biovolume and carbon biomass

Cell abundance and cell size were used to calculate total biovolume for all autotrophic cells $> 2 \mu m$ by adopting the geometric shapes suggested by [Hillebrand et al. \(1999\)](#page-26-19). Biovolume was then converted to total carbon biomass (C) for each taxon according to the C:volume

Fig. 3. Cell abundance of major groups of phytoplankton (> 2 µm) within the SCM across all full-profile stations in Subarctic and Arctic Seas, (a) total nano- and microphytoplankton, (b) diatoms (Bacillariophyceae) c) dinoflagellates (Dinophyceae), (d) unidentified flagellates/coccoids (2–8 µm), and (e) cryptomonads (Cryptophyceae). Abundance of each group is denoted at each station by the size of the coloured circles, with a reference circle shown to the left of each map. Note the variation in scale between panels.

(C:vol) equations of [Menden-Deuer and Lessard \(2000\)](#page-26-20). The data presented throughout the paper is based upon separate equations for the C density (in pg μ m⁻³) of diatoms and of other non-diatom protists ([Menden-Deuer and Lessard, 2000](#page-26-20)). However, in order to examine uncertainty of C estimates for selected representative taxa, we also estimated C biomass using the single C:vol equation for all protists proposed by [Montagnes et al. \(1994\)](#page-26-21).

2.5. Ratios of phytoplankton carbon to chlorophyll a, and biogenic silica to phytoplankton carbon

Ratios of total phytoplankton C ($> 2 \mu m$) to chlorophyll a (phyto C:chl *a*), and biogenic silica (SiO₂ or bSi) to phytoplankton C ($> 2 \mu m$) (bSi:phyto C) were calculated at each station using the chl a and bSi data concurrently collected from the depth of the SCM [\(Varela et al.,](#page-27-2) [2013; Wyatt et al., 2013\)](#page-27-2).

3. Results

3.1. Oceanographic properties

The physical and chemical properties of the five domains in the C3O/JOIS programs, already described by [Carmack and McLaughlin](#page-26-2) [\(2011\) and Varela et al. \(2013\),](#page-26-2) are summarized here in [Tables 1 and 2](#page-2-0), and [Fig. 2,](#page-4-0) together with the SCM depths where phytoplankton samples were taken for the present study.

The subarctic domains ESNP and BB-LS were characterized by weak stratification of the water column [\(Table 2](#page-3-0), [Fig. 2\)](#page-4-0). Based on the difference between the surface and the SCM depth, the upper water column was also relatively homogenous in concentrations of nitrate (NO₃⁻), silicic acid (Si(OH)₄), orthophosphate (PO₄³⁻), ammonium $(NH₄⁺)$ and urea-N [\(Table 2,](#page-3-0) see also [Varela et al., 2013\)](#page-27-2). Clearly, the BB-LS domain differed from the ESNP in having lower temperatures ([Table 2](#page-3-0), [Fig. 2\)](#page-4-0), and the presence of extensive ice cover at the Baffin Bay stations [\(Tables 1 and 2](#page-2-0)). Depths of the SCM were within the upper 40 m ([Fig. 2](#page-4-0), [Tables 1 and 2](#page-2-0)), and always shallower than the average depth of 1% surface irradiance in both domains. The BE-CH domain was characterized by moderate ice cover at the northern stations ([Table 1](#page-2-0)), and increasing density stratification as a result of lower surface temperature and salinity [\(Table 2](#page-3-0), [Fig. 2\)](#page-4-0). Average depth of the SCM was < 20 m and at most stations shallower than the 1% surface irradiance depth ([Table 2](#page-3-0)). Extensive ice cover, with low surface temperature and salinity, characterized the BS-CB and CAA domains ([Tables 1 and 2](#page-2-0), [Fig. 2](#page-4-0)), leading to pronounced density stratification of the upper water column. However, the CAA differed from the BS-CB in having weaker stratification, and shallower average SCM depth and 1% surface irradiance depth.

Other than the stratification resulting from low salinity surface waters, particularly in the BS-CB and CAA, the most striking difference between domains was the gradient in nutrient concentrations between

Fig. 4. Carbon biomass (C) of major groups of phytoplankton (> 2 µm) within the SCM across all full-profile stations in Subarctic and Arctic Seas, (a) total nano- and microphytoplankton, (b) diatoms (Bacillariophyceae), (c) dinoflagellates (Dinophyceae), (d) unidentified flagellates/coccoids (2–8 µm), and (e) cryptomonads (Cryptophyceae). Total C biomass for each group is denoted at each station by the size of the coloured circles, with a reference circle shown to the left of each map. Note the variation in scale between panels.

the ESNP in the west and the BB-LS in the east, particularly in terms of dissolved NO_3^- and Si(OH)₄ ([Table 2](#page-3-0)). These nutrient and chl α concentrations and distributions are described in more detail by [Varela](#page-27-2) [et al. \(2013\)](#page-27-2).

3.2. Spatial variation in abundance and carbon biomass of nano- and microphytoplankton

[Figs. 3 and 4](#page-5-0) show abundance and carbon biomass (C) data for all stations where samples were taken from the SCM, including those omitted from domains (see [Table 1\)](#page-2-0). Total phytoplankton (> 2 µm) abundances reached 7.13 \times 10⁶ cells L⁻¹ off the west coast of Vancouver Island at station NP-1, which is a coastal station omitted from the domains (see [Table 1](#page-2-0)), and 11.2×10^6 cells L^{−1} in the Davis Strait, northern Labrador Sea, at station LS-7 ([Fig. 3a](#page-5-0)). In contrast, the lowest total cell abundances of \sim 0.19 × 10⁶ cells L⁻¹ occurred at stations CB-2a in the southwestern Canada Basin and CB-15 in the north-central Canada Basin [\(Fig. 3a](#page-5-0)). Large-scale variations in total cell abundance ([Fig. 3a](#page-5-0)) were largely driven by variations in the abundance of unidentified flagellates/coccoids (2–8 µm) that reached a maximum of 2.94×10^6 cells L⁻¹ at station BB-1 in Baffin Bay [\(Fig. 3d](#page-5-0) and [Appendix](#page-19-0)). The highest abundance of autotrophic dinoflagellates of 0.26 × 10⁶ cells L⁻¹ occurred at station LS-4 in the Labrador Sea with minimum numbers of < 0.001 \times 10⁶ cells L⁻¹ at station CB-15 in the north-central Canada Basin [\(Fig. 3](#page-5-0)c and [Appendix\)](#page-19-0). Maximum diatom

numbers reached 1.60×10^6 cells L⁻¹ at BL-2 [\(Fig. 3](#page-5-0)b and [Appendix](#page-19-0)), a station located on the continental shelf east of the Chukchi Sea and excluded from both the BE-CH and BS-CB domains ([Table 1](#page-2-0)). Higher diatom numbers (2.70 \times 10⁶ cells L⁻¹) were observed at surface station CB-28aa (data in [Appendix](#page-19-0) only) located in the south Beaufort Sea. Cryptomonads exceeded ~0.2 × 10⁶ cells L⁻¹ at station BS-5 in the southern Bering Sea and at several stations in the BB-LS domain, peaking at 0.63×10^6 cells L⁻¹ at station BB-1 in eastern Baffin Bay off the coast of Greenland ([Fig. 3e](#page-5-0) and [Appendix\)](#page-19-0).

Total phytoplankton cell C ($> 2 \mu m$) showed a similar distribution [\(Fig. 4](#page-6-0)a) to that of cell abundance [\(Fig. 3a](#page-5-0)), although unidentified flagellates/coccoids (2–8 µm) had little influence on the overall distribution pattern [\(Fig. 4a](#page-6-0) and d), which was largely driven by the C biomass of diatoms and autotrophic dinoflagellates [\(Fig. 4b](#page-6-0) and c). Dinoflagellate C reached 107.5 µg C L⁻¹ at station BC-2 in the eastern Chukchi Sea off the coast of Alaska, and was 100.8 µg C L⁻¹ at station NP-1 off Vancouver Island [\(Fig. 4](#page-6-0)c), where diatom C was only 0.41 µg C L⁻¹ (Fig. 4b). Diatom C varied greatly, with the highest level for the entire study area of 331.6 µg C L−¹ recorded at station CAA-10 in the Canadian Arctic Archipelago [\(Fig. 4](#page-6-0)b). Diatom C biomass levels of < 0.1 µg C L^{-1} were measured at half of the offshore stations in the BS-CB domain, with no diatom C measured at the depth of the SCM at stations BI-2 in the eastern Canada Basin and CB-4 in the western Canada Basin [\(Fig. 4b](#page-6-0)). Cryptomonad C reached 15.4 µg C L^{-1} at station BS-5 in the southern Bering Sea, and 18.7 µg C L⁻¹ at station BB-1 in eastern Baffin Bay [\(Fig. 4e](#page-6-0)).

Fig. 5. Average cell abundance and carbon biomass (C) for major phytoplankton groups ($>$ 2 µm) within the SCM in each oceanographic domain. Error bars denote \pm one standard error (SE). (a) Total nano- and microphytoplankton, (b) diatoms (Bacillariophyceae), (c) dinoflagellates (Dinophyceae), (d) unidentified flagellates/coccoids (2–5 µm), (e) unidentified flagellates/coccoids (5–8 µm), and (f) cryptomonads (Cryptophyceae).

3.3. Domain averages of abundance and carbon biomass of nano- and microphytoplankton

[Fig. 5](#page-7-0) summarizes the average cell abundance and C biomass for the major phytoplankton groups. Average total cell abundance [\(Fig. 5](#page-7-0)a, [Table 3\)](#page-9-0) was highest (3.61 \pm 1.15 \times 10⁶ cells L⁻¹) in BB-LS, lowest (0.34 \pm 0.04 × 10⁶ cells L⁻¹) in the BS-CB, and varied between 0.85 and 1.50×10^6 cells L⁻¹ in the other three domains (ESNP, BE-CH and CAA). Average abundances of flagellates/coccoids (2–5 µm and 5–8 µm) ([Fig. 5d](#page-7-0) and e) and dinoflagellates [\(Fig. 5c](#page-7-0)) were highest in the BB-LS. The high abundance of dinoflagellates observed at excluded station NP-1 ([Fig. 3](#page-5-0)c and [Appendix](#page-19-0)) did not contribute to the ESNP domain average ([Fig. 5c](#page-7-0)). Average abundance of diatoms varied between 0.31×10^6 and 0.56 x10⁶ cells L⁻¹ in the BE-CH, CAA and BB-LS domains [\(Fig. 5b](#page-7-0)), with much lower numbers in the ESNP and BS-CB. Average abundance of cryptomonads was highest in the BB-LS domain ([Fig. 5f](#page-7-0)), and was very low in the BS-CB and CAA.

Whereas variations in total phytoplankton ($> 2 \mu m$) abundance across domains ([Fig. 5a](#page-7-0), [Table 3\)](#page-9-0) were largely driven by the contribution of flagellates/coccoids (2–8 µm) ([Fig. 5d](#page-7-0), e, [Table 3](#page-9-0)), variations in total phytoplankton C were dominated by the contribution of dino-flagellates and diatoms [\(Fig.5a](#page-7-0), b and c, [Table 3\)](#page-9-0). Unidentified flagellates/coccoids (2–8 μm) averaged < 10 μg C L⁻¹ in all domains, and < 5 µg C L⁻¹ in ESNP, BS-CB and CAA [\(Fig.5](#page-7-0)d, e, [Table 3\)](#page-9-0). Total C biomass was highly variable between stations ([Fig. 4\)](#page-6-0), with the average ([Fig. 5a](#page-7-0)) highest in the CAA (97.9 µg C L^{-1}) and slightly lower in the BE-CH (94.2 µg C L⁻¹) and BB-LS (64.9 µg C L⁻¹) domains ([Table 3](#page-9-0)). Average total C biomass was 19.1 μg C L⁻¹ in the ESNP and only 4.7 μg CL^{-1} in the BS-CB. Cryptomonads made a relatively small contribution

to overall biomass with average C of < 6 µg L⁻¹ for all domains ([Fig. 5f](#page-7-0), [Table 3](#page-9-0)).

[Fig. 6](#page-8-0) shows several 'minor' taxa that were identified in the samples but not included in [Figs. 3](#page-5-0)–5. These taxa occurred in relatively low abundance with contributions to total C biomass that were largely insignificant. Except for Phaeocystis spp. ([Fig. 6](#page-8-0)a) with a very high abundance at station LS-7 in the BB-LS, average C biomass for these groups did not exceed 1 µg L⁻¹ for any other domain [\(Fig. 6](#page-8-0) – note the variation in scales between panels). However, distributions of these taxa showed interesting trends. For example, the prasinophyceans Pyramimonas spp. occurred in much higher numbers in the BB-LS domain than in other domains ([Fig. 6](#page-8-0)b), whereas the chrysophycean Dinobryon balticum [\(Fig. 6](#page-8-0)c) showed elevated abundance at some CAA stations. Abundances of silicoflagellates Dictyocha speculum were generally low, with domain averages < 1000 cells L^{-1} , and with a maximum average C biomass in the BS-CB domain [\(Fig. 6](#page-8-0)d). The abundance of euglenids, mainly Eutreptiella sp., showed no clear pattern [\(Fig. 6e](#page-8-0)), whereas the chlorophycean Chlamydomonas sp. was more abundant in the western domains, and less so in BS-CB [\(Fig. 6](#page-8-0)f).

3.4. Relative contribution of major taxonomic groups

The relative contribution of each taxonomic group to total cell abundance and C content was calculated for each station and then averaged by domain ([Fig. 7](#page-9-1) and [Table 4\)](#page-10-0). The 'others' group shown in [Fig. 7](#page-9-1) represented the sum of the minor taxa shown in [Fig. 6.](#page-8-0) Unidentified flagellates/coccoids (2–8 µm) contributed on average 61–90% of total cell abundance over all domains [\(Fig. 7a](#page-9-1), [Table 4\)](#page-10-0), but represented only 9–39% of total phytoplankton C ([Fig. 7b](#page-9-1), [Table 4](#page-10-0)).

Fig. 6. Average cell abundance and carbon biomass (C) for several less common phytoplankton taxa within the SCM in each oceanographic domain. Error bars denote ± 1 SE. Note the variability in scales, since the abundance of these taxa span several orders of magnitude. These taxa were not included in the major groups shown in [Fig. 5.](#page-7-0) (a) Phaeocystis spp. (Prymnesiophyceae), (b) Pyramimonas spp. (Prasinophyceae), (c) Dinobryon spp., mainly Dinobryon balticum (Chrysophyceae), (d) silicoflagellate Dictyocha speculum (Dictyochophyceae), (e) euglenids (Euglenophyceae, mainly Eutreptiella sp.) and (f) Chlamydomonas sp. (Chlorophyceae).

Despite relatively low abundances, dinoflagellates contributed significantly to total C biomass in the western domains (ESNP and BE-CH) and in the BS-CB. Diatoms averaged only 4–30% of total cell abundance ([Fig. 7a](#page-9-1), [Table 4\)](#page-10-0), but represented 18–70% of total phytoplankton C ($> 2 \mu m$) over the five domains ([Fig. 7b](#page-9-1), [Table 4](#page-10-0)), with the greatest contribution to total C in the BE-CH, CAA and BB-LS domains. On average, the overall combined contribution of diatoms and dinoflagellates to total C biomass was ∼55% in BB-LS and ESNP, > 80% in BE-CH and CAA, and ~49% in BS-CB. Flagellates/coccoids (2–8 μm) contributed 39% of total C biomass in the BS-CB, which was the highest relative contribution of the group for any domain.

Over the whole study area, there was a strong relationship, within the SCM, between total particulate organic carbon (POC) [\(Wyatt et al.,](#page-27-1) [2013\)](#page-27-1) and total phytoplankton C ($> 2 \mu m$) estimated from cell counts in the present study [\(Fig. 8](#page-11-0)a). Relationships were also evident between total POC in the SCM and the C biomass of diatoms ([Fig. 8](#page-11-0)b), dino-flagellates [\(Fig. 8](#page-11-0)c) and unidentified flagellates/coccoids $(2-8 \mu m)$ ([Fig. 8d](#page-11-0)). In stations where POC in the SCM exceeded \sim 10 μmol C L⁻¹, diatoms and dinoflagellates made the dominant contribution to total C ([Fig. 8b](#page-11-0) and c). However, where POC was ~1–10 µmol C L⁻¹, unidentified flagellates/coccoids (2-8 µm) made a greater relative contribution to total C ([Fig. 8d](#page-11-0)).

3.5. Dominant taxa

Individual taxa can have significance either through consistent presence at many stations within a domain, or else, through a more sporadic occurrence, with high biomass at only a few stations within a

domain. The upper panel in [Table 5](#page-12-0) shows the average C biomass for the dominant phytoplankton taxa present at ≥50% of stations within each domain (50% is simply a nominal value chosen to highlight 'more common' taxa). The percentage of total stations in the domain at which each the taxa was present is also shown. The lower panel in [Table 5](#page-13-0) shows taxa present at < 50% of stations, but at biomass of > 1 µg C L⁻¹ when averaged for the whole domain (here also, 1 µg C L⁻¹ is simply a nominal average biomass chosen to highlight a 'significant' contribution). In other words, the lower panel of [Table 5](#page-13-0) highlights taxa typically present at elevated biomass but only at one or two stations within a domain. Some other taxa (e.g. those in [Fig. 6\)](#page-8-0) do not appear in [Table 5](#page-12-0) because of infrequent presence and/or average C biomass of $\lt 1$ µg L^{-1} .

Unidentified flagellates/coccoids (2–8 µm) and cryptomonads were almost ubiquitous through all domains, and because they could not be examined any further taxonomically, their total C contributions reported in [Table 5](#page-12-0) (upper panel) are the same as the domain averages shown in [Fig. 5](#page-7-0). Cryptomonads contributed more C in the BB-LS and ESNP domains [\(Fig. 5f](#page-7-0) and [Table 5](#page-12-0) upper panel), and showed a very low C biomass of only 0.09 µg C L⁻¹ in the BS-CB.

Among the diatoms in the ESNP domain, only Pseudo-nitzschia spp. had a consistent presence (67% of stations), although at low biomass ([Table 5](#page-12-0) upper panel). The diatom Neodenticula seminae [\(Fig. 9](#page-14-0)a and b) was present in the ESNP but did not contribute significantly to C biomass (therefore does not feature in [Table 5\)](#page-12-0), and was not observed in other domains. Corethron hystrix was also noted sporadically in ESNP but not in other domains [\(Table 5](#page-13-0) lower panel). Gymnodinium spp. ([Table 5](#page-12-0) upper panel) and Tripos muelleri [\(Fig. 9](#page-14-0)c, [Table 5](#page-13-0) lower panel)

Fig. 7. Average percentage contribution of the major taxonomic groups to total abundance and carbon biomass (C) of phytoplankton $(> 2 \mu m)$ within each oceanographic domain. Taxa not belonging to the four major groups are included in the 'others' category detailed in [Fig. 6](#page-8-0). Relative contributions were calculated for each station and then averaged for each domain. Data with standard errors are shown in [Table 4.](#page-10-0)

were the only dinoflagellates that made significant contributions to C biomass in the ESNP.

The BE-CH domain was characterized by a rich diversity of diatoms. Some, such as Chaetoceros gelidus and Cylindrotheca closterium, were present at most stations [\(Table 5](#page-12-0) upper panel). Other diatoms were

present at a single station [\(Table 5](#page-13-0) lower panel), such as Coscinodiscus wailesii [\(Fig. 9d](#page-14-0)), and Pauliella taeniata ([Fig. 9](#page-14-0)e and f). Several unidentified centric species were also notable [\(Table 5](#page-12-0) upper panel). Of the dinoflagellates, Gymnodinium spp. and Gyrodinium spp. occurred at most stations ([Table 5](#page-12-0) upper panel), with Alexandrium sp. and Peridiniella catenata ([Fig. 9](#page-14-0) g) significant at individual stations [\(Table 5](#page-13-0) lower panel).

Total phytoplankton C ($> 2 \mu m$) was very low in the BS-CB domain, with unidentified flagellates/coccoids (2–8 μ m) and *Gymnodinium* spp. providing the most significant contributions [\(Table 5](#page-12-0) upper panel). Although diatoms were present in low abundance ([Figs. 3 and 5](#page-5-0)b) in BS-CB, no individual diatom species or genus made a noticeable contribution to total C biomass other than Chaetoceros gelidus ([Table 5](#page-12-0) upper panel, [Fig. 9](#page-14-0) h). However, diatoms contributed on average ∼18% of total C in the BS-CB ([Fig. 7,](#page-9-1) [Table 4\)](#page-10-0) suggesting that the bulk of this must have been contributed by species present at very low abundance (see [Appendix\)](#page-19-0). Prasinophyceans were present in the BS-CB domain, but did not contribute significantly to total phytoplankton C ([Fig. 6](#page-8-0)). The chrysophycean Dinobryon balticum [\(Fig. 9i](#page-14-0)) was abundant at surface stations MK-3 and BL-4 in the BS-CB domain (see [Appendix](#page-19-0)), but did not contribute significantly to total C biomass within the SCM [\(Fig. 6](#page-8-0)). The only phytoplankton taxon with average C higher in BS-CB than elsewhere was the silicoflagellate Dictyocha speculum [\(Fig. 9j](#page-14-0)), albeit with low C biomass (average < $0.5 \mu g \text{ C L}^{-1}$).

The CAA was characterized by high biomass of relatively few diatom taxa, principally Thalassiosira spp. [\(Fig. 9k](#page-14-0), [Table 5](#page-12-0) upper panel) and particularly resting spores of T. antarctica var. borealis ([Fig. 9](#page-14-0)l, m, [Fig. 5](#page-7-0) lower panel), together with other unidentified centric diatoms ([Table 5](#page-12-0) upper panel). Dinoflagellates were less significant in the CAA with only Alexandrium sp. [\(Table 5](#page-13-0) lower panel) and Gymnodinium spp. ([Table 5](#page-12-0) upper panel) contributing an average of $> 1 \mu$ g C L⁻¹. Average biomass of flagellates/coccoids (2–8 µm) in the CAA was lower than in the BE-CH domain [\(Table 5](#page-12-0) upper panel).

Phytoplankton assemblages in the BB-LS domain [\(Table 5](#page-12-0) upper panel) were dominated by flagellates/coccoids (2–8 µm), cryptomonads ([Fig. 9](#page-14-0)n), Gymnodinium spp. ([Fig. 9o](#page-14-0), p), and various diatoms such as Thalassiosira spp., including, among others, T. nordenskioeldii ([Fig. 9](#page-14-0)q, [Table 5](#page-13-0) lower panel) and T. antarctica var. borealis [\(Fig. 9](#page-14-0)r, [Table 5](#page-13-0) lower panel), Chaetoceros gelidus [\(Table 5](#page-13-0) lower panel) and Bacterosira bathyomphala ([Table 5](#page-13-0) lower panel). Also present were pennate diatoms of the genus Fragilariopsis [\(Fig. 9s](#page-14-0), [Table 5](#page-12-0) upper panel), the dinoflagellate Pyrophacus cf. horologium [\(Table 5](#page-13-0) lower panel), and single cells of Phaeocystis spp. (prymnesiophycean, [Table 5](#page-13-0) lower panel). Pyramimonas spp. were also present at most stations, but with relatively low C biomass ([Table 5](#page-12-0) upper panel).

Table 3

Average abundance and carbon biomass (C) of total phytoplankton (> 2 µm), and each major taxonomic group in each oceanographic domain. Only data from the SCM for all full-profile stations were used. Standard error appears in parentheses. The 'others' group refer to the sum of the minor taxa shown in [Fig. 6.](#page-8-0)

Table 4

Average percentage contribution of the major taxonomic groups to total abundance and carbon biomass (C) of phytoplankton (> 2 µm) in each oceanographic domain. Only data from the SCM for all full-profile stations were used. The % abundance or % C biomass for each group was calculated for each station within each domain, and then averaged for the domain in question; this calculation allows presentation of the standard error for each group (in parentheses). Data also appears graphically in [Fig. 7.](#page-9-1) The 'others' group refer to the sum of the minor taxa shown in [Fig. 6.](#page-8-0)

3.6. Ratios of phytoplankton carbon to chlorophyll a, and biogenic silica to phytoplankton carbon

Using the taxonomic analysis of the phytoplankton assemblage presented here, and the concentrations of biogenic particles for the same stations (from [Varela et al., 2013; Wyatt et al., 2013\)](#page-27-2), we calculated ratios of phytoplankton C to chl a (phyto C:chl a), and biogenic silica to phytoplankton C (bSi:phyto C). Domain averages for phyto C:chl a were in the range 51 to 87 g C g chl a^{-1} in the ESNP, BB-LS and BE-CH domains, but only 31 g C g chl a^{-1} in the BS-CB [\(Fig. 10](#page-15-0)a). In the CAA, the high average phyto C:chl a ratio (150 \pm 76 g C g chl a⁻¹) was substantially skewed by stations CAA-2 and CAA-10, where very high levels of phytoplankton C were not accompanied by significant increases in chl a. These two stations coincided with elevated abundance of diatom resting spores, conceivably with higher C:chl a ratios than in vegetative cells. Ratios of bSi:phyto C decreased from an average of 0.49 (\pm 0.19) mol Si mol C⁻¹ in the BE-CH domain, to 0.22 (\pm 0.09) mol Si mol C−¹ in the BB-LS domain ([Fig. 10b](#page-15-0)).

4. Discussion

4.1. Approach and strengths of this study

The present study maximized regional coverage, minimized the role of seasonal variability by sampling only in summer (July–August), and maintained consistency in methodology and taxonomic definitions throughout all regions. However, some limitations to this approach remain. Firstly, in order to maximize regional coverage whilst at the same time maintaining the logistics of sample processing to a realistic level, only samples from the depth of the SCM were processed, as in several other studies (e.g. [Ardyna et al., 2011](#page-25-1)). Although phytoplankton within the SCM make a large contribution to total water column production in Arctic waters [\(Martin et al., 2012](#page-26-4)), it is unclear to what extent these assemblages within the SCM are representative of the water column. Reasonable agreement between SCM and surface samples has been documented, at least for some taxa [\(Booth, 1988; Joo](#page-25-2) [et al., 2012](#page-25-2)), although in the Beaufort Sea in summer, [Balzano et al.](#page-25-4) [\(2012\)](#page-25-4) have reported increased numbers of nanophytoplankton in the SCM compared to the surface. The SCM can also at times include sea-ice unicellular autotrophs, as well as phytoplankton, and species such as Nitzschia frigida were sporadically observed in the present study. Secondly, with such high variability of physical, chemical and biological properties over such a wide geographical area, potential temporal effects on our data cannot be fully separated from spatial variability, as noted by [Varela et al. \(2013\)](#page-27-2). Finally, light microscopy provides limited taxonomic information on smaller phytoplankton [\(Booth, 1987\)](#page-25-3) and preservatives may have differential effects on taxonomic groups ([Booth,](#page-25-2) [1988\)](#page-25-2). Therefore, some heterotrophic protists may have been included in our "unidentified flagellate/coccoid (2–8 µm)" fraction.

Robust comparisons between studies, both at a pan-Arctic scale and at the smaller scale, have been hampered by methodological inconsistencies. Definitions and categories for cell size, trophic status and taxonomic classification often differ between studies, and data for individual taxa have been reported in terms of cell abundance ([Booth,](#page-25-2) [1988; Booth et al., 1993; Booth and Horner, 1997; Schloss et al., 2008;](#page-25-2) [Sukhanova et al., 2009; Ardyna et al., 2011; Joo et al., 2012\)](#page-25-2), total biovolume (e.g. [Joo et al., 2012](#page-26-22)), or total C biomass ([Booth, 1988;](#page-25-2) [Booth et al., 1993; Booth and Horner, 1997; Gosselin et al., 1997;](#page-25-2) [Sukhanova et al., 2009](#page-25-2)). This variety of approaches has further complicated between-study comparisons of C contributions by various taxa. The present study has attempted to minimize these inconsistencies and provides comparative estimates of abundance and biomass across a wide geographical area.

4.2. Oceanographic features of the domains

To facilitate interpretation of taxonomic data, we provide here a summary of the physical, chemical and biological features of the five domains, originally more fully described by [Carmack and McLaughlin](#page-26-2) [\(2011\)](#page-26-2) and later by [Varela et al. \(2013\)](#page-27-2).

In the ESNP domain, temperature, salinity and density at the surface and within the SCM indicated relatively low stratification within the euphotic zone, with low concentrations of chl a and high concentrations of $NO₃$ consistent with phytoplankton constrained by the availability of iron over much of this domain (e.g. [Boyd et al., 1996; Boyd et al., 2004](#page-26-23)). Water column stratification was also relatively weak in the BE-CH, although stronger than in the ESNP. A pronounced SCM and notable drawdown in surface $NO₃⁻$ concentrations were evident in the BE-CH during the July study period, consistent with a domain strongly influenced by seasonal cycles, early sea-ice melt, and enriched by nutrient-laden Pacific water earlier in the season (e.g. [Walsh et al., 1989; Carmack and McLaughlin,](#page-27-5) [2011\)](#page-27-5). The BS-CB domain was characterized by extensive ice-cover, low salinity surface waters, and a deep euphotic zone. Average summer surface concentrations of NO₃⁻ of 0.27 μmol L⁻¹ and chl *a* of 0.05 μg L⁻¹ were consistent with previous observations of a highly stratified oligotrophic system with a deep SCM and phytoplankton limited principally by irradiance and the vertical supply of $NO₃⁻$ [\(Carmack and McLaughlin, 2011\)](#page-26-2). Although weaker than in the BS-CB, stratification was also evident in the vertical distributions of density and $\rm NO_3^-$ in the CAA. Surface $\rm NO_3^-$ and chl a were higher in the CAA than in the BS-CB, and SCM layers were shallower, possibly as a result of mixing processes that enrich surface waters and favor phytoplankton growth in this topographically complex domain [\(Carmack](#page-26-2) [and McLaughlin, 2011](#page-26-2)). The upper water column of the BB-LS was characterized by weaker density stratification than in the CAA, with relatively high chl a concentrations in both surface waters and the SCM. Compared to the western domains, summer concentrations of $NO₃⁻$ and Si(OH)₄ in the BB-LS were low in both surface waters and within the SCM [\(Varela et al.,](#page-27-2) [2013\)](#page-27-2).

Fig. 8. Relationship between total particulate organic carbon (POC) at the depth of the SCM (from [Wyatt et al., 2013](#page-27-1)) and carbon biomass (C) for the major phytoplankton taxa (> 2 µm) sampled from within the SCM during the present study. (a) Total nano- and microphytoplankton C, (b) diatom C, (c) dinoflagellate C and (d) unidentified flagellates/ coccoids C (2–8 µm; sum of 2–5 µm and 5–8 µm fractions). Solid red lines represent 1:1 relationship between POC and total phytoplankton C ($> 2 \mu$ m). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

4.3. Comparison of phytoplankton assemblages with previous studies

The most extensive study of phytoplankton assemblages in the open ocean ESNP to date ([Booth, 1988\)](#page-25-2) noted vertical variability in total phytoplankton C and taxonomic composition throughout the euphotic zone in August 1984 at Ocean Station Papa (OSP, 50°N 145°W). The numerical dominance by unidentified 2–8 µm flagellates/coccoids (78% of total abundance, on average) in our study was in good agreement with that of [Booth \(1988\)](#page-25-2) where small cells (< 5 µm) were shown to represent 67% of total abundance. These unidentified flagellates/coccoids averaged 4.8 µg C L⁻¹ of the 19.1 \pm 3.5 µg C L⁻¹ for total nanoand microphytoplankton C in our study, compared to \sim 4 µg C L⁻¹ in the < 5 μm fraction of the total C of $~\sim$ 8–12 μg C L^{−1} (estimated from the graphs of [Booth, 1988\)](#page-25-2). Diatoms represented on average 4.1 ± 1.8 ug C L⁻¹ in the present study, compared to < 1 to 9 ug C L⁻¹ in [Booth \(1988\)](#page-25-2). [Booth et al. \(1993\)](#page-25-5) extended the 1984 study of [Booth](#page-25-2) [\(1988\)](#page-25-2) to include data from the summer of 1988 and reported an average of 20 µg C L⁻¹ for total phytoplankton C in the upper 60 m at OSP, with large diatoms (> 24 µm) contributing an average of ∼7 to 8 µg C L−¹ [\(Booth et al., 1993\)](#page-25-5). The dominance of total C biomass by unidentified flagellates/coccoids (2–8 µm) in the ESNP is consistent with studies showing that the bulk of new production and nitrate uptake occurs in the $<$ 5 μ m fraction [\(Varela et al., 2013](#page-27-2)). With relatively low concentrations of bSi, and with ∼80% total chl a in the < 5 µm fraction [\(Wyatt et al., 2013](#page-27-1)), our observations are consistent with an open ocean system numerically dominated by small phytoplankton ([Varela and Harrison, 1999; Harrison, 2002; Peña and Varela, 2007](#page-27-6)), and with growth of large diatoms limited by iron availability ([Boyd](#page-26-23) [et al., 1996; Boyd et al., 2004\)](#page-26-23).

Total phytoplankton C of 264 and 424 µg C L⁻¹ was measured in the SCM at two stations in summer 1994 in the Chukchi Sea ([Booth and](#page-26-24) [Horner, 1997\)](#page-26-24), and 285 µg C L⁻¹ at 36 m depth on the Chukchi Shelf in summer 2002 [\(Sukhanova et al., 2009](#page-26-25)). The domain average in the present study of 94.2 \pm 26.6 µg C L⁻¹ was lower but consistent with that of ∼100 µg C L−¹ for late summer 2008 on the Chukchi Shelf ([Coupel et al., 2012](#page-26-26)). Our measurements of C biomass for unidentified flagellates/coccoids (2–8 µm) of 2.4–10.1 µg C L⁻¹ in the BE-CH were similar to those of 1.8 and 6.3 µg C L⁻¹ for the 2–10 µm 'flagellate' fraction for two stations in the Chukchi Sea [\(Booth and Horner, 1997](#page-26-24)). These ranges were much lower however than the highly variable nanophytoplankton fraction (34.3 \pm 34.2 μg C L⁻¹) for the Chukchi Shelf in late summer 2008 ([Coupel et al., 2012\)](#page-26-26), and the 4–10 µm flagellate fraction (0.6–159 µg C L⁻¹) reported by [Sukhanova et al. \(2009\)](#page-26-25) for similar depths in the Bering Strait and the Chukchi Shelf in summer 2002. Total nano- and microphytoplankton C in the BE-CH was ∼5-fold higher than in the ESNP in the present study, with the unidentified flagellate/coccoid fraction (2–8 µm) only ∼30% higher than in the ESNP. The elevated total C in the BE-CH was therefore principally driven by larger cells, with diatoms representing ∼38% and dinoflagellates ~45% of total C. Diatom C averaging 36.2 \pm 14.3 µg C L⁻¹ for the BE-CH in our study was in reasonable agreement with the 55.8 µg C L−¹ for diatoms on the Chukchi shelf during late summer 2008 ([Coupel et al., 2012](#page-26-26)). However, average diatom C biomass was lower than the maximum of 125 µg C L⁻¹ on the Chukchi Shelf in summer 2002 [\(Sukhanova et al., 2009](#page-26-25)), and two measurements of $~\sim$ 410 and $~\sim$ 260 μg C L⁻¹ on the central Chukchi Shelf in 1994 ([Booth](#page-26-24) [and Horner, 1997](#page-26-24)). Dinoflagellate C measurements on the Chukchi Shelf of 4.2 and 1.3 µg C L⁻¹ ([Booth and Horner, 1997](#page-26-24)), and < 1 µg C L^{-1} ([Sukhanova et al., 2009\)](#page-26-25) were much lower than our domain average of 47.5 (\pm 19.5 μg C L⁻¹) and that of 10.2 μg C L⁻¹ for the summer of 2008 [\(Coupel et al., 2012\)](#page-26-26). The abundance range of 0.007–0.037 × 10⁶ cells L⁻¹ for the dominant dinoflagellates *Gymno*dinium spp. was a little lower in our study than that of 0.012–0.100 \times 10⁶ for an unidentified *Gymnodinium* sp. in the Chukchi Sea in summer 2008 reported by [Joo et al. \(2012\),](#page-26-22) who also corroborated our observation of elevated abundance of cryptomonads in the south Bering Sea. Our abundance measurements for Chaetoceros spp., Thalassiosira spp. and Fragilariopsis spp. in the BE-CH were also of a similar order to those reported for the SCM by [Joo et al. \(2012\)](#page-26-22). Domination of total C by large cells (e.g. Gymnodinium, Thalassiosira, Coscinodiscus, and unidentified centric diatoms) in the BE-CH is

Table 5
Upper Panel: Average carbon biomass (C) for phytoplankton taxa present at ≥50% of stations in each of the Arctic and Subarctic domains. For each domain, the column to the right indicates the percentage of statio Upper Panel: Average carbon biomass (C) for phytoplankton taxa present at $\geq 50\%$ of stations in each of the Arctic and Subarctic domain. For each domain, the column to the right indicates the percentage of stations in taxa was present.

Lower Panel Average carbon biomass (C) for phytoplankton taxa present at $<$ 50% of stations in each of the Arctic and Subarctic domains, but with a domain average of > 1 µg C L⁻¹ (a nominal figure chosen to represent a Lower Panel: Average carbon biomass (C) for phytoplankton taxa present at < 50% of stations in each of the Arctic and Subarctic domains, but with a domain average of > 1 µg C1~1 (a nominal figure chosen to represent a 'si biomass). These species/genera occur more sporadically than the taxa shown in the upper panel, but can represent high levels of C at certain stations. Note that some taxa can appear in certain domains in the upper panel of

 8 Average total phytoplankton C for each domain is higher than the sum of the totals in upper and lowers panels, because of less abundant taxa not shown here. The dominant taxa shown in the two panels typically account average total phytoplankton C biomass for each domain. average total phytoplankton C biomass for each domain. 3e

Fig. 9. Light micrographs (LM) and scanning electron micrographs (SEM) of selected phytoplankton taxa observed across Subarctic and Arctic Seas around northern North America. Scale bars are given on each micrograph. (a) Diatom Neodenticula seminae (SEM). Frustule in girdle view. (b) Diatom Neodenticula seminae (SEM). Valve in internal view. (c) Dinoflagellate Tripos muelleri (LM). d) Diatom Coscinodiscus wailesii (LM). Cell in girdle view. (e) Diatom Pauliella taeniata (LM). Cells in ribbon-shaped colonies in apical view. (f) Diatom Pauliella taeniata (SEM). Valve with raphe in internal view. (g) Dinoflagellate Peridiniella catenata (LM). Cells in a colony. (h) Diatom Chaetoceros gelidus (LM). Cells forming spherical colonies. (i) Chrysophycean Dinobryon balticum (LM). Cells in loricae forming branched colonies. (j) Silicoflagellate Dictyocha speculum (LM). Silica skeleton. (k) Cells of diatom Thalassiosira spp. (LM). (l) Single cell of a centric diatom, and resting spore of Thalassiosira antarctica var. borealis indicated with arrow (LM). (m) Resting spores of the diatom T. antarctica var. borealis (SEM). (n) Cryptomonad cells (LM). (o) and (p) Dinoflagellate Gymnodinium spp. (LM). (q) Diatom Thalassiosira nordenskioeldii (SEM). (r) Diatom T. antarctica var. borealis (LM). Cells in long chain. (s) Diatom Fragilariopsis oceanica (SEM). Valve in internal view.

consistent with studies showing that the $>$ 5 μ m size fraction contributes $> 60\%$ of new production, nitrate uptake, and total chl a ([Varela et al., 2013, Wyatt et al., 2013](#page-27-2)), and with bSi > 4-fold higher than in the ESNP [\(Wyatt et al., 2013\)](#page-27-1).

In the BS-CB domain, the phytoplankton assemblage was numerically dominated (∼90%) by unidentified flagellates/coccoids (2–8 µm), with average abundance falling within the range observed for nanophytoplankton by [Balzano et al. \(2012](#page-25-4)) within the SCM of the deeper

Fig. 10. Ratios of total (> 2 µm) phytoplankton carbon (total phyto C), total chlorophyll a (chl a), and total particulate biogenic silica (bSi) within the SCM over the five oceanographic domains. (a) ratio of phyto C:chl a (b) ratio of bSi:phyto C. The average ratio $(± 1 SE)$ is based upon data from all full-profile stations for each domain. Data for chl a taken from [Varela et al. \(2013\)](#page-27-2) and data for bSi from [Wyatt et al. \(2013\)](#page-27-1).

offshore stations in the Beaufort Sea in summer 2009. Total nano- and microphytoplankton C averaged 4.70 \pm 0.83 µg C L⁻¹, a value within the summer ranges of 1.9 µg C L⁻¹ for the SCM in the Canada Basin in 1994 [\(Gosselin et al., 1997](#page-26-11)), and 9.2 \pm 7.6 µg C L⁻¹ for the Beaufort Sea SCM in 2009 [\(Coupel et al., 2015\)](#page-26-27). The unidentified flagellates/

coccoids fraction (2–8 µm) averaged 1.66 \pm 0.24 µg C L⁻¹ for the BS-CB domain, agreeing well with that of 1.1 \pm 0.2 µg C L⁻¹ for the < 10 µm fraction in the Canada Basin [\(Booth and Horner, 1997](#page-26-24)). Diatom C of 0.86 \pm 0.38 µg C L⁻¹ and dinoflagellate C of 1.59 \pm 0.42 µg C L^{-1} were of the same order as biomass of the unidentified flagellate/ coccoid (2–8 µm) fraction in this study, but somewhat lower than the measurements of 4.5 µg C L⁻¹ and 2.7 µg C L⁻¹ respectively observed by [Coupel et al. \(2012\).](#page-26-26) The C biomass of unidentified flagellate/coccoids (2–8 µm) in the BS-CB was ∼26% of that observed in the BE-CH, whereas combined C biomass of diatoms and dinoflagellates represented only ∼3% of that in the BE-CH domain. This very low contribution of diatoms and dinoflagellates to total C in the BS-CB is consistent with previous findings showing that the $>$ 5 μ m fraction contributed < 27% of new production, nitrate uptake, and total chl a ([Varela et al., 2013, Wyatt et al., 2013](#page-27-2)). Concentrations of bSi averaged only ∼4% of those observed in the BE-CH [\(Wyatt et al., 2013](#page-27-1)), consistent with our maximum abundances of Chaetoceros spp. of only ∼5% of those observed in the BE-CH. Few other diatom genera were observed in the BS-CB, with abundances for various species of Chaetoceros of a similar order of magnitude to those observed by [Joo et al. \(2012\)](#page-26-22). Our observations support previous studies identifying the BS-CB as an oligotrophic environment numerically dominated by nanophytoplankton and picophytoplankton ([Booth and Horner, 1997; Gosselin](#page-26-24) [et al., 1997](#page-26-24)).

Fewer records of phytoplankton assemblages are available for the CAA. [Ardyna et al. \(2011\)](#page-25-1) reported relative abundance, as size-fractionated chl a, for several taxa during summer 2005 in extended Arctic transects that included the CAA. Flagellates contributed ∼50% on average to the abundance of total protist cells $> 2 \mu m$, whereas diatoms varied from $< 10\%$ at some stations to $> 60\%$ at others (average ∼30%) and dinoflagellates contributed only ∼2% ([Ardyna et al.,](#page-25-1)

> Fig. 11. Relationships between % ice cover and (a) % surface irradiance within the SCM, and (b) the depth of the SCM. Relationships between surface water $\sigma_{\rm F}$ and (c) % surface irradiance within the SCM, and (d) depth of the SCM. Dotted line represents non-significant linear regression $(p > .05)$ and solid line represents significant regression ($p < .05$). Data points for ice cover and σ_T of represent surface water properties for full-profile stations (see [Tables 1 and 2\)](#page-2-0).

Fig. 12. Relationships between carbon biomass (C) of major taxonomic groups of phytoplankton (> 2 µm) within the SCM, and surface temperature (a, b, c, d), surface NO₃⁻ (e, f, g, h), surface σ_T (i, j, k, l) and surface DIC (m, n, o, p) across the five oceanographic domains. Total = C in total nano- and microphytoplankton. Dotted line represents non-significant linear regression (p > .05) and solid line represents significant regression (p < .05). Data points represent surface water properties for full-profile stations (see [Table 2\)](#page-3-0).

[2011\)](#page-25-1). Our abundance data for the CAA domain were similar, with ∼64% unidentified flagellates/coccoids (2–8 µm), ∼28% diatoms and ∼3% dinoflagellates. Our measurement of total nano- and microphytoplankton C in the CAA was similar to but more variable (97.9 \pm 53.5 µg C L⁻¹) than that in the BE-CH, whereas C biomass levels of unidentified flagellate/coccoids (2–8 µm) and dinoflagellates were much lower than in the BE-CH. The CAA domain was characterized by the high spatial variability in diatom C biomass, with a

maximum estimate of \sim 332 µg C L⁻¹ at station CAA-10. These observations are consistent with high and variable bSi concentrations and with total chl a, nitrate uptake and new production dominated by the > 5 µm size fraction [\(Varela et al., 2013, Wyatt et al., 2013\)](#page-27-2). In late summer and autumn, large areas of the high Arctic including the CAA have been characterized either as oligotrophic, dominated by eukaryotic picophytoplankton $(< 2 \mu m)$ and unidentified nanoflagellates (2–20 μm), or as eutrophic, dominated by centric diatoms, such as

Table 6

Estimates of C quota per cell for a range of representative taxa observed during this study. Carbon estimates using the protist and diatom specific C:vol equations given by [Menden-Deuer](#page-26-20) [and Lessard \(2000\)](#page-26-20) were compared with those using the single equation adopted by [Montagnes et al. \(1994\)](#page-26-21). The percent difference in C estimates between the two approaches is given in the final column.

Chaetoceros spp. ([Ardyna et al., 2011](#page-25-1)), and our data support these observations.

Our study did not include stations in the North Water Polynya (NOW) of the BB-LS domain, a highly productive area in the north of Baffin Bay (e.g. [Mostajir et al., 2001; Vidussi et al., 2004; Garneau et al.,](#page-26-28) [2007\)](#page-26-28). [Mostajir et al. \(2001\)](#page-26-28) also included surface stations further south in Baffin Bay that were more comparable with stations in our BB-LS domain, although their study was carried out in fall 1999 with samples taken only from the ocean surface. Our total nano- and microphytoplankton abundance for the BB-LS varied between 0.56 and 11.2×10^6 cells L⁻¹, with abundances in Baffin Bay itself (i.e. not including the Labrador Sea) of between 2 and 4 \times 10^6 cells L⁻¹. This is in reasonable agreement with abundances of 0.81 to 4.35 \times 10⁶ L⁻¹ for cells < 20 µm in size reported by [Mostajir et al. \(2001\)](#page-26-28) for Baffin Bay (not including the NOW region). Total phytoplankton C ($> 2 \mu m$) averaged 64.9 \pm 14.9 μg C L⁻¹ in the BB-LS in our study, which was much lower than in the BE-CH and CAA domains. Unidentified flagellates/coccoids (2–8 μm) contributed 9.6 \pm 1.8 μg C L⁻¹ of this total, a greater C biomass than in other domains. Previous studies have shown that the $> 5 \mu m$ size fraction contributed $> 50\%$ of total chl a, new production and nitrate uptake [\(Varela et al., 2013, Wyatt et al.,](#page-27-2) 2013) in the BB-LS, with bSi averaging $< 50\%$ of concentrations observed in the CAA [\(Wyatt et al., 2013](#page-27-1)). Our observations for the BB-LS support the notion of a diverse nano- and microphytoplankton summer assemblage with a lower diatom C biomass than in the CAA. The greater relative influence of unidentified flagellates/coccoids (2–8 µm), dinoflagellates and cryptomonads in the BB-LS could also be linked to lower summer concentrations of $Si(OH)_4$ in the sector of this domain that we studied [\(Varela et al., 2013\)](#page-27-2). Blooms of Phaeocystis spp. have been shown to be a characteristic feature further south in the Labrador Sea ([Fragoso et al., 2016, 2017](#page-26-29)) and it has been speculated that variations in Si^{*} (Si(OH)₄ minus NO_3^- concentrations) may be involved in the competition between Phaeocystis and diatoms in this area [\(Fragoso](#page-26-29) [et al., 2016, 2017](#page-26-29)). Our observations of high abundance of Phaeocystis spp. at station LS-7 in the Labrador Sea, combined with relatively low concentrations of $Si(OH)_4$ in the BB-LS [\(Varela et al., 2013](#page-27-2)) are consistent with this notion.

4.4. Overview of large-scale variability in nano- and microphytoplankton

The C biomass of unidentified flagellates/coccoids (2–8 µm) was relatively low and varied ∼6-fold among domains, whereas average diatom and dinoflagellate C biomass varied ∼100 fold and ∼30-fold respectively. In the BE-CH and CAA, diatoms and dinoflagellates together contributed > 80% of total nano- and microphytoplankton C biomass. Thus, large-scale spatial variability of total C was driven by the abundance of these larger cells. Despite their relatively low C biomass, the unidentified flagellate/coccoid (2–8 µm) fraction provided a greater proportion of total within the less productive ESNP and BS-CB domains where they contributed ∼27 and 39% to total C, respectively. With potentially higher intrinsic growth rates under low nutrient conditions relative to larger cells, and efficient grazing by heterotrophic protists, these smaller cells may contribute significantly to fixation, cycling and trophic transfer of C in these oligotrophic areas.

The diatoms Thalassiosira spp. and Chaetoceros spp. were the dominant taxa driving major variations in total phytoplankton C (> 2 µm) among Arctic domains. This observation is consistent with the findings of [Poulin et al. \(2011\)](#page-26-8) who provided a pan-Arctic review of historical taxonomic records of phytoplankton presence. They indicated that the dominant taxa in the Canadian and Alaskan Arctic were centric diatoms of the genera Thalassiosira and Chaetoceros, and pennate diatoms of the genera Cylindrotheca and Fragilariopsis. The most significant dinoflagellate genus recorded by [Poulin et al. \(2011\)](#page-26-8) was Protoperidinium, present in the Scandinavian and Russian Arctic, but not in the Canadian Arctic. Protoperidinium spp. were observed sporadically in the present study, but were considered to be heterotrophic forms (see [Appendix](#page-19-0)). Our observations of the dominant athecate genera Gymno-dinium and Gyrodinium contrast with the review of [Poulin et al. \(2011\)](#page-26-8) that did not mention the Gymnodinium genus. It is not known whether this difference results from temporal shifts in dominant genera as a result of environmental change in Arctic waters, or whether fragile athecate dinoflagellates were under-represented by certain sampling and preservation methods in the historical record. It is notable however, that our observations of high abundance of Gymnodinium spp. in the BE-CH were corroborated by the study of [Joo et al. \(2012\)](#page-26-22) from the same period as our study. Due to limitations of light microscopy, we did

not provide detailed taxonomical information on nanophytoplankton, and molecular techniques provide tremendous promise for examining the diversity of the pico- and nanophytoplankton diversity (e.g. [Poulin](#page-26-8) [et al., 2011](#page-26-8)).

Several taxa were recorded in low abundance and did not contribute significantly to total nano- and microphytoplankton C, but their presence appeared to be associated with particular domains or water masses. Cryptomonads, for example, were present in lower abundance in Arctic than in Subarctic domains, whereas silicoflagellates were more abundant in Arctic regions such as BS-CB. Elevated abundances of Phaeocystis spp. were mainly associated with the Labrador Sea, whereas the chlorophyceaen Chlamydomonas sp. was associated more with western domains (ESNP and BE-CH), and the prasinophyceaens Pyramimonas spp. more prominent to the east. Other small chlorophyceaens and prasinophyceaens could have been included in our unidentified flagellate/coccoid (2–8 µm) category, and therefore we must be cautious in extrapolating spatial patterns of smaller cells (2–8 µm) enumerated using light microscopy. Based on molecular techniques, small prasinophyceaens such as Micromonas spp. appear to have a more pan-Arctic distribution ([Lovejoy et al., 2007](#page-26-30)) than implied by the Pyramimonas spp. observed in our study, and other small phytoplankton have been shown to be cold-adapted and endemic to Arctic waters (e.g. [Lovejoy et al., 2007\)](#page-26-30).

Phytoplankton growth has been shown to be limited by iron in central ESNP waters (e.g. [Boyd et al., 1996](#page-26-23)), by irradiance, nitrate and iron in late summer in the Beaufort Sea ([Taylor et al., 2013\)](#page-26-31), but enhanced by upwelling of nutrients into shallow water in the BE-CH ([Carmack and McLaughlin, 2011; Varela et al., 2013\)](#page-26-2). Given the large spatial scales involved and the variability in oceanographic features within and among domains, it is perhaps not surprising that there is little evidence for any single physical or chemical variable driving spatial changes in nano-and microphytoplankton C. The depth of the SCM varied considerably across the study area [\(Fig. 2,](#page-4-0) [Tables 1 and 2\)](#page-2-0) but was deeper in the BS-CB and CAA domains where ice cover was greatest. However, relationships were insignificant between % icecover and both average irradiance within the SCM ([Fig. 11a](#page-15-1)), and the depth of the SCM ([Fig. 11](#page-15-1)b). By contrast, a weak but significant positive relationship ($p < .05$) was evident between density of surface water and average irradiance in the SCM [\(Fig. 11c](#page-15-1)), and a stronger inverse relationship ($p < .001$) between surface water density and the depth of the SCM ([Fig. 11](#page-15-1)d). Our data are therefore consistent with a deeper SCM in areas of reduced surface σ_T and increased stratification, with consequentially reduced irradiance within the SCM, particularly in the BS-CB domain.

Stratification of Arctic surface waters has increased in recent years as a result of sea-ice melt and increased riverine input ([McLaughlin and](#page-26-5) [Carmack, 2010](#page-26-5)). In some areas, this increased stratification dampens vertical mixing of the upper water column with a consequential deepening of the nitracline and the SCM [\(McLaughlin and Carmack, 2010](#page-26-5)). This has been associated with a concomitant decrease in the abundance of nanophytoplankton and an increase in the relative contribution of picophytoplankton ([Li et al., 2009\)](#page-26-1). In the present study, both in surface water and within the SCM, temperature was relatively invariant across the BE-CH, BS-CB, CAA and BB-LS domains. Indeed, there was no evidence of a significant relationship between surface temperature and C biomass of the major taxonomic groups [\(Fig. 12a](#page-16-0)–d). Similarly, surface water $NO₃⁻$ showed no significant relationship ([Fig. 12](#page-16-0)e-h) with the C content of major taxonomic groups. This presumably results from low nitrate concentrations being a feature of both reduced vertical supply to stratified surface waters – with consequently low phytoplankton biomass (e.g. BS-CB) – and of high biological drawdown with consequently high biomass (e.g. BE-CH). The significant relationship between surface

water σ_T and depth of SCM ([Fig. 11](#page-15-1)d), did not extend to a relationship between surface water σ_T and total C biomass, or the C biomass of diatoms or dinoflagellates [\(Fig. 12](#page-16-0)i–k). However, a significant re-lationship (p < .001) was evident ([Fig. 12](#page-16-0)l) between surface water σ_T and the C biomass of unidentified flagellates/coccoids (2-8 µm). Similarly, relationships between surface water DIC and the C biomass of major taxonomic groups [\(Fig. 12](#page-16-0)m–p) showed a significant relationship (p < .001, [Fig. 12](#page-16-0)p) only with unidentified flagellates/coccoids (2–8 µm). It is not known whether this represents some form of direct effect of lower DIC, or simply results from the relatively conservative relationship between DIC and salinity/density. Our observations support the contentions of [Li et al. \(2009\)](#page-26-1), in suggesting that the deeper SCM (and decreased irradiance within the SCM) associated with surface water stratification, also appears to be associated with a significant decrease in at least the 2–8 µm size-fraction of the nanophytoplankton assemblage (2-20 µm). Density stratification clearly plays a key role in the vertical mixing that shapes phytoplankton communities in the Beaufort Sea, CAA and Baffin Bay as suggested by [Ardyna et al. \(2011\)](#page-25-1).

4.5. Spatial variability in C:chl a and bSi:C ratios in nano- and microphytoplankton

The wide range of phyto C:chl a ratios, from 31 in BS-CB to 51–150 in the other domains, agrees with the modelled range of 20 to $>$ 160 g C g chl a^{-1} between latitudes of 0° and 60°N ([Taylor et al., 1997\)](#page-26-32), the measured range of 19–150 g C g chl a^{-1} in a North Pacific coastal environment (Welschmeyer and Lorenzen, 1984), and the range of ∼10 and > 200 g C g chl a^{-1} measured in phytoplankton cultures [\(Geider,](#page-26-33) [1987; Taylor et al., 1997\)](#page-26-33). Ratios of C:chl a appear to be driven principally by complex interactions between irradiance, nutrients and temperature ([Geider, 1987; Taylor et al., 1997\)](#page-26-33), but the impact of these factors on ratios over the wide geographical area covered by the present study is unknown. Although some evidence of high C:chl a ratios have been reported at low temperatures in Arctic waters (review of [Geider,](#page-26-33) [1987\)](#page-26-33), our data show a wide range of ratios across a variety of Arctic domains where temperature is relatively invariant. Minimum phyto C:chl a ratios were observed in the deep SCM of the BS-CB domain, as has been noted previously ([Booth and Horner, 1997; Sherr et al., 2003](#page-26-24)), and may be indicative of phytoplankton acclimated to low irradiance ([Smith and Sakshaug, 1990; Sherr et al., 2003\)](#page-26-34).

Ratios of bSi:phyto C averaged ∼0.22 to 0.49 across domains, somewhat higher than the bSi:C range (average 0.13) given by [Brzezinski \(1985\)](#page-26-35) for cultured diatoms. However, our observed range may have been influenced by the low POC:PON ratios observed in certain regions [\(Crawford et al., 2015](#page-26-36)), and by variations in the presence of non-siliceous taxa among domains. [Baines et al. \(2010\)](#page-25-6) observed bSi per unit cell volume to be ∼6-fold higher for diatoms in the $Si(OH)₄$ rich waters of the Antarctic Zone of the Southern Ocean, compared to $Si(OH)_4$ poor waters of the eastern equatorial Pacific. Similarly, observations of lower ratios of bSi:phyto C in our eastern domains could be consistent with lower summer concentrations of Si (OH)4 in these areas [\(Varela et al., 2013](#page-27-2)), potentially constraining the growth and/or silicification of some diatom species. Caution should be exercised in interpretation of these data because of the potential contribution of detrital Si to the bSi:C ratio ([Krause et al., 2010](#page-26-37)). Clearly, however, the plasticity in both bSi:phyto C and phyto C:chl a ratios across these broad oceanographic domains has implications for biogeochemical processes, modelling, and remote sensing studies.

4.6. Potential biases in estimation of C from cellular volume

The use of constant C:vol conversion factors over large ranges in

phytoplankton size can result in systematic errors in biomass estimates ([Menden-Deuer and Lessard, 2000\)](#page-26-20). In the present study, estimates of cellular C for each phytoplankton taxon were calculated using separate equations for diatoms and for all other protists ([Menden-Deuer and](#page-26-20) [Lessard, 2000](#page-26-20)). In order to examine potential biases when comparing C estimates with other studies, estimates of C calculated from a single equation for all protists [\(Montagnes et al., 1994](#page-26-21)) were compared with our estimates for several representative taxa in [Table 6](#page-17-0). The choice of equation clearly has a significant impact on the estimated C quota per cell, and this varies both with cell size and taxonomic group ([Table 6](#page-17-0)). Cellular C quotas estimated using the [Menden-Deuer and Lessard](#page-26-20) [\(2000\)](#page-26-20) equations were 43–72% higher than using the [Montagnes et al.](#page-26-21) [\(1994\)](#page-26-21) equation for typical flagellates and dinoflagellates, 60% higher for small diatoms, and between 24 and > 60% lower for medium and large diatoms ([Table 6\)](#page-17-0) (and even lower for larger diatoms, data not shown).

Because the relative proportion of taxonomic groups varied across domains, variations in C quota per cell also have important implications for estimates of total nano- and microphytoplankton C. In the ESNP and BS-CB domains, where the relative contribution of unidentified flagellates/coccoids (2–8 µm) was greater, average total phytoplankton C ($> 2 \mu m$) estimated using the [Montagnes et al. \(1994\)](#page-26-21) equation was ∼1–12% lower than using the equations of [Menden-Deuer and Lessard](#page-26-20) [\(2000\).](#page-26-20) In the BB-LS and BE-CH domains, a more balanced mix of diatoms, dinoflagellates and unidentified flagellates/coccoids (2-8 μ m) was present, and the average total C calculated using the [Montagnes](#page-26-21) [et al. \(1994\)](#page-26-21) equation was ∼16–33% higher than using [Menden-Deuer](#page-26-20) [and Lessard \(2000\)](#page-26-20) equations. In the CAA domain, where the relative contribution of diatoms was greatest, the average total C calculated using the [Montagnes et al. \(1994\)](#page-26-21) equation was almost a factor of 2 higher than using the [Menden-Deuer and Lessard \(2000\)](#page-26-20) equations. The present study has adopted the most appropriate equations that take into account variations in C density between diatoms and other autotrophic protists ([Menden-Deuer and Lessard, 2000](#page-26-20)). However, allometry ignores species-specific variability with the final aim of providing average estimates for comparative purposes [\(Menden-Deuer and Lessard, 2000](#page-26-20)). Our phytoplankton C estimates will therefore be influenced by the taxonomic similarity between assemblages in each domain and the taxa represented in the study of [Menden-Deuer and Lessard \(2000\).](#page-26-20) These considerations demonstrate that the C:vol equations adopted should be a major consideration when comparing C biomass estimates both historically and in future studies.

5. Conclusions

Temporal and spatial comparisons of phytoplankton assemblages among studies can be hampered by methodological, taxonomic and size category inconsistencies. Our study attempted to minimize some of

Appendix

these issues by adopting a consistent methodology and approach to provide a snapshot of nano- and microphytoplankton assemblages across five Subarctic and Arctic domains during the summers of 2007 and 2008.

Average abundance of nano- and microphytoplankon varied ∼10 fold, and total C biomass varied ∼20-fold among domains. Spatial variation in biomass was driven mainly by diatoms and dinoflagellates, together accounting for 55% to > 80% of total phytoplankton C (> 2 µm) in all domains except for the oligotrophic BS-CB.

Estimates of C quota for various taxa were highly sensitive to the C:vol equations adopted, with a consequential impact on estimates of total phytoplankton C (> 2 µm). Diatom-rich areas are particularly susceptible to uncertainties, because of specific issues with allometric scaling of C:vol equations.

Variations in taxonomic composition in each domain led to significant variations in phyto C:chl a and bSi:phyto C ratios, with important implications for modelling and biogeochemical studies in these areas.

The depth of the SCM, and to some extent the taxonomic composition of phytoplankton within it, were shown to be significantly influenced by the density of surface water. These observations support recent studies highlighting the role of thinning and melting sea ice in shaping phytoplankton assemblages, with potential consequences for primary productivity and transfer of C to higher trophic levels.

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Samples for identification of nano- and microphytoplankton were collected from the depth of the subsurface chlorophyll maximum (SCM) where present (see [Table 1](#page-2-0) for sampling depths). Data correspond to cruises 2007–19, 2007–20, and 2008–02 in the summers of 2007 and 2008 in Arctic and Subarctic Seas (see [Fig. 1](#page-1-0) and [Table 1](#page-2-0) for station and domain locations). Additional samples from surface stations were taken in the upper 2–10 m of the water column (see [Table 1](#page-2-0)); these surface data are presented here only (and nowhere else in this paper) to support other chemical and biological data for surface stations shown in [Varela et al. \(2013\).](#page-27-2) Note that in contrast to the text, tables and a few figures, abundance here is expressed in cells L^{−1} × 10³ because of the many uncommon taxa present at low abundances.

X denotes species recorded as present in low abundance but not enumerated.

∗ Denotes heterotrophic dinoflagellates not included in phytoplankton C biomass estimates.

Appendix Table 1 Abundance of phytoplankton taxa at stations in the Eastern Subarctic North Pacific (ESNP) domain.

Abundance of phytoplankton taxa at stations in the Bering and Chukchi Seas (BE-CH) domain. No surface stations were sampled within this domain.

Abundance of phytoplankton taxa at stations in the Beaufort Sea and Canada Basin (BS-CB) domain.

Abundance of phytoplankton taxa at stations in the Canadian Arctic Archipelago (CAA) domain.

Appendix Table 5 Abundance of phytoplankton taxa at stations in the Baffin Bay and Labrador Sea (BB-LS) domain. No surface stations were sampled within this domain.

Abundance of phytoplankton taxa at stations excluded from the domain classification.

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