EOCENE CONCHOSTRACANS FROM THE LANEY MEMBER OF THE GREEN RIVER FORMATION, WYOMING, USA

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ABSTRACT—Two new species of conchostracans, *Cyclestherioides wyomingensis* and *Prolynceus laneyensis*, belonging to the Cyclestheriidae and Lynceidae, respectively, are described from the Eocene Laney Member of the Green River Formation of Wyoming, USA. These are the first Cenozoic fossil conchostracans to be formally reported from North America. *Cyclestherioides wyomingensis* has a close affinity to the extant *Cyclestheria hislopi*. The latter is a pan-tropical species, found between approximately 30°N and 35°S. Thus, *Cyclestherioides* may indicate that the Laney Member was deposited in a subtropical setting. The conchostracans and associated fauna of the Laney Member suggest that this member was deposited nearshore in a shallow lake. The concept of *Cyclestherioides* is emended and some fossil species referred to this genus are revised.

INTRODUCTION

THE WORLDWIDE record of Cenozoic fossil conchostracans is limited to only a few regions: China, with many localities ranging from Paleocene to Eocene (Hong et al., 1974; Chen, 1975; Chen and Shen, 1979, 1980, 1981; Shen and Chen, 1979; Shen and Zhang, 1979), Mongolia (late Paleocene–early Eocene; Trusova and Badamgarav, 1976), Brazil (Oligocene; Gallego and Mesquita, 2000), and Argentina (late Paleocene and Miocene– Pliocene; Gallego and Mesquita, 2000; Petrulevicius, 2001).

In the present paper, we describe two new species, *Cyclestherioides wyomingensis* and *Prolynceus laneyensis*, assigned to the Cyclestheriidae and Lynceidae, respectively, from the Anvil Wash locality, the Laney Member of the Green River Formation of Wyoming, USA. These are the first Cenozoic fossil conchostracans to be formally described from North America. These fossils were previously assigned with doubt to *Cyzicus* sp. (Buchheim, 1978) and briefly described by Grande (1984). At that time they were the only Paleogene conchostracans reported outside of China, Mongolia, and Brazil. Abundant conchostracan specimens provide a basis for this study of their taxonomy and the relationships between fossil and modern forms.

LOCALITY AND HORIZON

The conchostracan specimens were found in the Laney Member of the Green River Formation at the Anvil Wash study locality, in the Green River Basin of Wyoming (Fig. 1). The site consists of exposures along Anvil Wash banks at 41°05.743N 109°36.169W in Sweetwater County, southwest of Green River, Wyoming.

The Green River Formation accumulated in the lake basins of Utah, Colorado, and Wyoming (Fig. 1). In the Green River Basin, the formation was divided by Roehler (1992) into the following units, in ascending order (Fig. 2): the Luman Tongue, Tipton Shale Member, that interfingers with the Farson Sandstone Member; the Wilkins Peak Member; and the uppermost member, the Laney Member, unique among other things because of the occurrence of conchostracans (Fig. 2). The Laney Member is well known for its exquisitely preserved fossil fish, among other fossils, as well as for its oil shales and clay shales, but the member also contains other siliciclastic and carbonate rocks.

A middle Eocene age has been assigned to the Laney Member based on several fossil evidence. A fossil mammal from the correlative Bridge Formation indicates a Bridgerian age (McGrew, 1959; Roehler, 1973). In addition, a palynomorph assemblage found in equivalent rocks in the northern part of the basin supports an early middle Eocene age for the Laney Member (Biaggi, 2001). More recently, radiometric dating of tuffs of the Green River Formation corroborates a middle Eocene age for this member (Machlus et al., 2002).

The stratigraphic section measured at the Anvil Wash study locality, which is approximately 23 m thick, consists mainly of kerogen-poor, laminated micrites (KPLM) (laminated carbonate rocks of clay-size grains and less than 2% total organic carbon) (Fig. 3). These laminated rocks contain abundant conchostracans, as well as gastropods and ostracodes. Other fossils include the herring *Knightia* (Jordan) (fish) (Buchheim and Surdam, 1981) and plant remains. Alternating with these laminated calcimicrites are thin (usually less than 10 cm thick) layers of dark-colored organic-rich mudstones that contain abundant plant fragments.

In the upper part of the section the conchostracans occur in a cyclic fashion between pulses of quartz-rich sediment. Near the top of the section a unit of papery, laminated calcimicrites contains a fish mass mortality bed and fish coprolites. The uppermost few meters consist of fine-grained siliciclastic rocks capped by Pleistocene coarse-grained deposits.

GLOBAL DISTRIBUTION OF CENOZOIC CONCHOSTRACANS

Cenozoic fossil conchostracans were first recorded in China by Wang *in* Hong et al., 1974. Subsequently, several localities were reported in the world (Chen, 1975; Zhang et al., 1976; Trussova and Badamgarav, 1976; Shen and Chen, 1979; Shen and Zhang, 1979; Chen and Shen, 1980, 1981; Gallego and Mesquita, 2000) (Table 1).

In Asia they were found from the upper Paleocene–lower Eocene and three faunas can be recognized (Zhang et al., 1976; Chen and Shen, 1981). The *Altaestheria* Trussova *in* Trussova and Badamgarav, 1976 from the Nalanburic Formation in the Nemegetin basin of Mongolia are considered to be synonymous with *Fushunograpta* Wang Si'en *in* Hong et al., 1974, based on their small shell configuration and slender and curvate radial lirae sculpture. The Nalanburic Formation yields fossil vertebrates that are regarded as equivalent to the Gashato fauna and thought to be of Late Paleocene age (Trussova and Badamgarav, 1976).

In South America there are three conchostracan horizons. *Tremembeglypta* (Gallego *in* Gallego and Mesquita, 2000) was reported from the Oligocene Tremembé Formation of the Taubaté Group, São Paulo State, Brazil. Petrulevicius (2001) mentioned conchostracans (*Euestheria*? Depéret and Mazeran, 1912) in the

TABLE 1-Stratigraphical distribution of the Cenozoic conchostracans.

Pliocene				
rnocene			-	Fushunograptidae?
Messinian				Loxomegaglytidae?
Tortonian				
Serravallian				
Langhian				
Burdigalian				
Aquitanian				
Chattian				Tremembeglypta
Rupelian				
Priabonian				
Bartonian			Cyclestherioides Prolynceus	
Lutetian				
Ypresian	Paraleptestheria menglaensis f.			
Thanetian	Fushunograpta changzhouensis f.	Fushunograpta (=Altaestheria)		Euestheriidae?
Selandian	Perilimnadia f.			
Danian				
	Messinian Tortonian Serravallian Langhian Burdigalian Aquitanian Chattian Rupelian Priabonian Bartonian Lutetian Ypresian Thanetian Selandian Danian	Messinian Tortonian Serravallian Langhian Burdigalian Aquitanian Chattian Rupelian Priabonian Bartonian Lutetian Ypresian Fushunograpta changzhouensis f. Selandian Danian	Messinian Tortonian Serravallian Langhian Burdigalian AquitanianHerein AquitanianChattian Rupelian Priabonian Bartonian Lutetian YpresianParaleptestheria menglaensis f.Paraleptestheria menglaensis f.Fushunograpta (=Altaestheria) Perilimnadia f.DanianPerilimnadia f.	Messinian Tortonian Serravallian Langhian Burdigalian AquitanianAutom Autom AquitanianChattian

Maíz Gordo Formation (late Paleocene) in Salta Province, Argentina. Neogene conchostracans in the Upper Member of the San Roque Formation in San Luis, Argentina, can be identified as two taxa, probably belonging to the Fushunograptidae and Loxomegaglyptidae (Gallego and Mesquita, 2000).

In Gemany middle Eocene conchostracans were reported from the Lake Messel (Richter and Baszio, 2001), but have not been formally described. However, as far as is known no credible Quaternary fossil conchostracans have yet been reported anywhere in the world. Thus it can be seen that the conchostracans of the Laney Member are a new Eocene assemblage (*Cyclestherioides– Prolynceus*) belonging to Cyclestheriida and Laevicaudata, respectively. They have a close affinity to the extant *Cyclestheria* Sars, 1887 and *Lynceus* Müller, 1776 (via Barnard, 1929).

SYSTEMATIC PALEONTOLOGY

All illustrated specimens came from the Laney Member of the Green River Formation at Anvil Wash section, Sweetwater County, Wyoming, and are housed in Nanjing Institute of Geology and Palaeontology, Chinese Academy of Sciences (NIGPAS 134166–134186) and in Palaeozoological Collection of the Facultad de Ciencias Exactas y Naturales y Agrimensura, UNNE, Corrientes, Argentina (PZ-CTES N° 7340–7350).

Order CONCHOSTRACA Sars, 1867 (via Tasch, 1969) Suborder CYCLESTHERIDA Sars, 1899 (via Martin and Davis, 2001)

Family CYCLESTHERIIDAE Sars, 1899 (via Tasch, 1969) Genus CYCLESTHERIOIDES Raymond, 1946, emended, Shen and Gallego

Emended diagnosis.—Carapace thin, subcircular through subovate, umbo subterminal or subcentral, a few fine or weak

growth lines, more or less equal width growth bands, small reticulate sculpture.

Discussion .--- Fossil Cyclestherioides was first recognized in the Upper Permian of Australia (Raymond, 1946). Subsequently, many authors ascribed species to this genus. However, most of the species possess thick valves, obvious growth lines, and numerous, even, narrow, growth bands, although they do share a subcircular outline and subcentral umbo, which are characteristics of Cyclestherioides. Novojilov (1958, 1960) placed Cyclotunguzites Novojilov, 1958, Wetlugites Novojilov, 1958, and Spaerestheria Novojilov, 1954 in the Cyclestheriidae based on their subcircular configuration, but the first two genera have a large umbo, which is a key character of the Palaeolimnadiidae, and the last genus has numerous well-developed growth lines and narrow growth bands. All three genera possess thick valves. Cyclestheria krivickii (Novojilov, 1958, p. 35, fig. 32, pl. 3, fig. 35) from Tatarian, Guff Khatanga, Russia, does not fit the type of valve of extant Cyclestheria in terms of carapace configuration and linear sculpture. Cyclestherioides franconica (Reible, 1962, p. 217, fig. 18, pl. 9, fig. 6) from the Upper Triassic of Germany has approximately 30 distinct growth lines and densely spaced growth bands in large adults and should not be assigned to this genus. Tasch (1987) erected Cyclestherioides (C.) alexandriae from the Lower Jurassic of the Transantarctic Mountains (p. 69, pl. 12, figs. 1, 2; pl. 14, figs. 7, 8; pl. 18, fig. 6). It probably belongs to Carapacestheria Shen, 1994 in terms of shell configuration and thick valve. The placement of Cyclestherioides (C.) pintoi (Tasch, 1987, p. 114, pl. 34, fig. 3a, Middle Permian Estrada Nova Formation, Brazil) in this genus is also questionable, because it bears more than 30 dense growth bands. Assignment of Cyclestherioides dalongkouensis Liu, 1989 and C. triangularis Liu, 1989 from the



FIGURE 1—Location map, showing study location and geology. The study site falls within the Laney Member (Tgl) of the Green River Formation, in the southern part of the Green River Basin, in the Eocene of the Rocky Mountain region (southwestern Wyoming, Utah, and Colorado).

Lower Triassic Dalongkou Formation of Xinjiang, China (p. 134, pl. 35, figs. 3, 4; p. 134, pl. 35, fig. 5), and *Cyclestheria rossica* Novojilov, 1959 from the Lower Triassic of Russia (*via* Molin and Novojilov, 1965, p. 61, pl. 8, fig. 2) to this genus are also questionable.

However, the following species may belong to the cyclestheriids: Cyclestherioides detykteica Novojilov in Molin and Novojilov, 1965 (p. 51, pl. 5, figs. 5, 6) from the Upper Permian Bu-Formation, Tungus, Russia; galiktin *Cyclestherioides* (?Cyclestherioides) sp. from the ?Upper Permian Middle Member of the Peine Formation, northern Chile (Gallego and Breitkreuz, 1994, pl. 2, fig. 6); Euestheria sparsa Shen in Zhang et al., 1976 (pl. 19, figs. 2, 3) from the Middle Triassic (Anisian) Badong Formation, Hubei, China; and Euestheria? cf. E. atsuensis (Kobayashi, 1952 in Zhang et al., 1976, pl. 18, fig. 16) from the Upper Triassic (Norian) Xujiahe Formation, Sichuan, China. Cyclestheriids seem to range from Late Permian to Holocene, although they are absent from the fossil record in the middle and late Mesozoic.

Based on morphological and molecular data, cyclestherids differ significantly from other spinicaudate conchostracans and should not be placed among spinicaudates. It is transitional between spinicaudatans and cladocerans (Spears and Abele, 2000; Martin and Davis, 2001; Olesen, 2002). Ax (1999) first suggested the term "Cladoceromopha" for the clade containing *Cyclestheria* plus Cladocera. Martin and Davis (2001) recommended removal



FIGURE 2—Regional stratigraphic relationships. Rocks and fossils studied occur in the Eocene Laney Member of the Green River Formation (GRF), southwestern Wyoming (after Roehler, 1992).

of *Cyclestheria* from the Spinicaudata. They placed it in a new suborder, the Cyclestherida. We follow this classification, but retain use of the order Conchostraca, which includes five suborders (Shen, 2003): Laevicaudata, Spinicaudata, Cyclestherida, Leaiida (Devonian to Permian), and Estheriellina (Triassic to Early Cretaceous).

CYCLESTHERIOIDES WYOMINGENSIS Shen and Gallego, new species

Figure 4.1–4.9; Appendix 1

Diagnosis.—Carapace thin, subcircular to ovate, small to moderate, 7–10 weak growth lines, the second one more obvious than others, irregularly small reticulations on growth bands.

Description.—Carapace thin, subcircular to ovate, small to moderate, 2.9–5.6 mm long, 2.4–4.5 mm high, height to length ratio of 0.7:0.9; dorsal margin long, gently arched upwards, umbo slightly rising above dorsal margin and situated at subterminal to subcentral of dorsal margin; anterior and posterior margins rounded, ventral margin slightly curved downward; growth lines 7–10, weak, fine, the second one more obvious than others, those of the middle part of valve weakest; more or less equal width growth bands; ornamented with small irregular transverse reticulations.

Etymology.—From Wyoming.

Types.—Holotype, NIGPAS 134177 (Fig. 4.6). Paratypes NIG-PAS 134166, 134168 (Fig. 4.1, 4.8); PZ-CTES N° 7340–7347. Over 183 individuals, among which 78 came from GR2 bed and 105 from GR3 bed. Most of the valves are complete or slightly displaced.

Occurrence.—The Eocene Laney Member of the Green River Formation, Anvil Wash section, Wyoming, USA.

Discussion.—The new species is distinguished from all other species assigned to this genus (Raymond, 1946; Molin and Novojilov, 1965; Zhang et al., 1976) by weak and fine growth lines in which the second one is more obvious than the others. It closely



FIGURE 3—The Anvil Wash stratigraphic section. *1*, Measured stratigraphic section (scale bar = 1 m); 2, expanded stratigraphy of units 10-16, kerogen-poor, laminated micrite (KPLM) alternates in cyclic fashion with claystone.

resembles extant *Cyclestheria hislopi* (Daday, 1926) in carapace configuration, the weak growth lines, and thin valve, but differs in that its second growth line is stronger than the others. *Paracyclestheria sinensis* (Shen and Dai, 1987) differs from the new species in its very weak growth lines.

Suborder LAEVICAUDATA Linder, 1945 (*via* Tasch, 1969) Family LYNCEIDAE Stebbing, 1902 Genus PROLYNCEUS Shen and Chen, 1984

PROLYNCEUS LANEYENSIS Shen and Gallego, new species Figure 4.10–4.16; Appendix 1

Diagnosis.—Subcircular to ovate, small, one peripheral growth line.

Description.—Carapace subcircular to ovate, small, 1.2–1.5 mm long, 1.05–1.1 mm high, height to length ratio of 0.72:0.88; dorsal margin long, slightly arched upward; anterior margin broadly curved, posterior margin rounded; ventral margin slightly curved downward; one peripheral growth line; sculpture unclear.

Etymology.—From the Laney Member of the Green River Formation.

Types.—Holotype NIGPAS 134181(Fig. 4.16); paratype NIG-PAS 134184 (Fig. 4.11, 4.12), PZ-CTES N° 7348–7350. Twelve specimens, among which two came from bed GR2 and 10 came from bed GR3. All specimens are complete valves, but are not well preserved.

Occurrence.—As for Cyclestherioides wyomingensis n. sp.

Discussion.—Fossil lynceids include two genera: *Paleolynceus* Tasch, 1956 and *Prolynceus*. The former occurs in the Lower Cretaceous Usqi-Karl Formation of Transbaikal, Siberia (Tasch, 1956; Krasinetz, 1964). The latter occurs in the upper Middle Jurassic Tuchengzi Formation of Beipiao, Liaoning Province, China (Shen and Chen, 1984). As for shell configuration they are very similar to those of extant forms, but *Paleolynceus* bears a larger valve (reaching 14.4 mm long) (Krasinetz, 1964) than other fossil lynceids (usually 1.3–5 mm long) (Royan and Alfred, 1971; Saunders and Wu, 1984). *Prolynceus* has a peripheral growth line, vertically prolonged shell gland situated at the anteromedial part of the valve, and a telson that is similar to the extant lynceids (Shen and Chen, 1984).

The new species is similar to *Prolynceus lineatus* (Shen and Chen, 1984, p. 315, pl. 1, figs. 6, 7) in having a peripheral growth line, although the appendages are not found from available collections. *P. laneyensis* differs from *P. lineatus* in its oval configuration and much smaller valve. *Prolynceus laneyensis* is distinguished from most modern species of *Lynceus* by the presence of a peripheral growth line.

PALEOENVIRONMENTAL IMPLICATIONS

Extant conchostracans can be found on all continents except Antarctica, and although some species inhabit permanent bodies of water, most can be found in ephemeral pools or temporary

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FIGURE 4-Magnification, 1, 4-6, 8-11, ×10; others as indicated. 1-9, Cyclestherioides wyomingensis Shen and Gallego, n. sp. 1, Paratype, right valve, scanning photo, GR3-2, NIGPAS 134166; 2, 3, irregular and small reticulations on the growth bands of Figure 1, $\times 40$, $\times 60$, scanning photo; 4, internal mold of the left valve, scanning photo, GR3-8, NIGPAS 134171; 5, external mold of the right valve, scanning photo, GR3-21, NIGPAS 134180; 6, holotype, external mold of the displaced right and left valve, GR3-15, NIGPAS 134177; 7, external mold of the right valve, ×15.3, scanning photo, GR3-1, NIGPAS 134165; 8, paratype, external mold of the left valve, scanning photo, GR3-4, NIGPAS 134168; 9, internal mold of the right valve, GR3-17, NIGPAS 134179. 10-16, Prolynceus laneyensis Shen and Gallego, n. sp. 10, Internal mold of the right valve, GR3-20, NIGPAS 134186; 11, 12, paratype, internal mold of the right valve, $\times 20$ (12), GR3–18, NIGPAS 134184; 13, external mold of the right valve, ×20, GR3-19, NIGPAS 134185; 14, internal mold of the right valve, ×20, GR3-4c, NIGPAS 134182; 15, internal mold of the right valve, ×20, GR3-7, NIGPAS 134183; 16, holotype, internal mold of the right valve, ×40, scanning photo, GR3-4b, NIGPAS 134181.



waters (Frank, 1988). They generally live in small, temporary or permanent inland ponds, floodplain pools, roadside ditches, rice paddy fields, margins of big lakes, estuaries, lagoons, and even inundated road ruts (Tasch, 1969; Chen and Shen, 1985). Most fossil conchostracans are believed to have lived in freshwater pools. Although some species can tolerate either high salinity or brackish water, they have never been found in full marine environments.

The extant *Lynceus* is basically benthic, mostly resting on the bottom but often swimming freely and slowly for short periods. It is probably omnivorous (Fryer, 1987) and is widely distributed in the world. In Europe, findings reach northwards to Scandinavia (~70°N) (Straškraba, 1965). Lynceus brachvurus Müller, 1776 in Straškraba, 1965 (p. 208) is present from about 9°E throughout Europe and Siberia to East Asia, and northwards as far as the arctic zone (Brtek and Thiéry, 1995). This species is eurytopic, living in ditches, temporary pools, and inundated fields. Lynceus manchuricus Daday, 1913 was found in a temporary pond in Inner Mongolia, China, associated with the notostracans (Uéno, 1940). Lynceus taianensis (Han and Shu, 1995) was collected from a stew in Taian, Shandong Province. In northern America, lynceids were found in a shallow (<1.0 m) ephemeral pond in Leon County, Florida, where the water temperature was 24°C and air temperature 25°C on 9 April 1984. Water in the pond also contained aquatic insects, frog tadpoles, and large anostracans (Martin et al., 1986). In South America, lynceids extend across from Venezuela, through Brazil, and then to southern Argentina (Crespo, 1993).

Cyclestheriidae includes two modern genera, *Cyclestheria* and *Paracyclestheria* (Shen and Dai, 1987), and the fossil genus *Cyclestherioides*. *Cyclestheria*, consisting of only one species (*C. hislopi*), occurs in India, Southeast Asia, southern China, Africa, South America, Central America, and Australia (Daday, 1926; Tasch, 1969; Halloy, 1981; Poi de Neiff and Bruquetas, 1983; Shen and Dai, 1987; Olesen et al., 1997). *Paracyclestheria* was found in Yunnan, China.

Cyclestheria hislopi is a pantropical species found between approximately 30°N and 35°S (Olesen et al., 1997). In Argentina, this species was collected from a flood zone in Paraná River Basin, where a fluvial lowland with abundant marshy and aquatic vegetation is located. This species was found between the roots of the water hyacinth Eichhornia crassipes (Martius) Solms-Laubach, 1883. Environmental data are conductivity 56-58 uS/cm, transparency 28 cm, and surface temperature 27.9°C (Halloy, 1981), recorded from an oxbow lake on the northern reaches of the Paraná River Valley. The limnological characteristics were: depth 3.4 m, pH between 7.0 and 7.4, conductivity 90 and 170 uS/cm, with no substantial dissolved oxygen variations (0.75-1.0 mg/l) (Poi de Neiff and Bruquetas, 1983). Poi de Neiff (2003) found C. hislopi living on Eichhornia azurea (Swartz) Kunth, 1842 along the Paraguay River floodplain. In Pantanal of southern Brazil the species is extremely frequent during the dry season (low-water period) (Por and Rocha, 1998). In India, this species was found in paddy fields of the Punjab, and ponds and small ditches near Ludhiana (Battish, 1981). Paracyclestheria sinensis was found in a pond (22°N, 101°E) in Xishuangbanna, Yunnan Province, China, with the following environmental data: elevation crassipes 860 m a.s.l., depth 0.2-0.4 m, pH 6, and surface temperature 20.5°C on 17 March 1957 (Shen and Dai, 1987).

The occurrence of the fossil pulmonate gastropod *Biomphalaria aequalis* White, 1880, along with these conchostracans, is paleoecologically very revealing. *Biomphalaria* sp. is typical of the Eocene of the Rocky Mountains but disappears from the region after this time. The extant *B. glabrata* Say, 1818, a modern-day equivalent, living in quiet, shallow, mud-bottomed ponds, but can survive desiccation in seasonal water bodies. It is found in the American tropics, and from laboratory experiments it is known that its optimum temperature for oviposition is 25°C (Taylor, 1988).

Therefore, *Cyclestherioides*, with the pulmonate gastropod in the Laney Member, indicates a subtropical setting. This substantiates the findings derived from the occurrence of warm temperate and subtropical macro- and microfloras, suggesting that the region received ample rainfall, perhaps seasonally, and that temperatures were relatively mild.

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APPENDIX 1-Qualitative and Quantitative data.

Specimens	Configuration	Preservation status	L(mm)	H(mm)	H/L	Gl	Notes
Cyclestherioides wyomi	ingensis n. sp.						
NIGPAS 134165	subcircular	ext.m.right valve	4.0	2.8	0.70	8	
NIGPAS 134166	circular	right valve	4.9	4.1	0.84	?	paratype
NIGPAS 134167	circular	int.m.left valve	4.7	4.2	0.89	?	1 71
NIGPAS 134168	circular	displaced valve	4.8	4.1	0.84	c.7	paratype
NIGPAS 134169	circular	displaced valve	3.7	3.2	0.86	?	1 71
NIGPAS 134170	subcircular	right valve	5.6	4.5	0.8	c.10	
NIGPAS 134171	circular	int.m.left valve	4.4	3.9	0.89	c.9	
NIGPAS 134172	circular	left valve	2.9	2.6	0.90	c.9	
NIGPAS 134173	circular	left valve	4.1	3.5	0.85	>6	
NIGPAS 134174	subcircular	right valve	3.3	2.6	0.78	?	
NIGPAS 134175	circular	left valve	3.2	2.7	0.84	?	
NIGPAS 134176	circular	right valve	2.9	2.4	0.83	7	
NIGPAS 134177	ovate	ext.m.displaced valve	4.1	3.6	0.88	c.7	holotype
NIGPAS 134178	circular	displaced valve	4.8	?	?	c.7	
NIGPAS 134179	circular	int.m.right valve	3.8	3.2	0.84	7	
NIGPAS 134180	circular	ext.m.right valve	4.1	3.5	0.85	c.8	
Prolynceus laneyensis	n. sp.						
NIGPAS 134181	subcircular	ext.m.left valve	1.2	1.1	0.88	1	holotype
NIGPAS 134182	subcircular	int.m.right valve	1.5	1.2	0.83	1	
NIGPAS 134183	subcircular	int.m.right valve	1.3	1.1	0.88	1	
NIGPAS 134184	ovate	int.m.right valve	1.5	1.1	0.73	1	paratype
NIGPAS 134185	ovate	ext.m.right valve	1.5	1.1	0.72	1	I
NIGPAS 134186	ovate	int.m.right valve	1.5	1.0	0.69	1	

L = length of the valve; H = height of the valve; Gl = number of growth lines; ext.m = external mold; int.m = internal mold; ? = not clear.