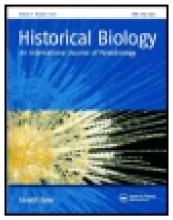
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First Tupinambinae teiid (Squamata, Teiidae) from the Palaeogene of South America

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The Lumbrera Formation in north-western Argentina is notable for preserving lizard remains from the Palaeogene of South America. In this paper, we offer a re-description and re-evaluation of the material belonging to the holotype of the teiid lizard *Lumbrerasaurus scagliai* (Donadío, 1985) recovered from sediments of the Lumbrera Formation near Pampa Grande, Salta province, Argentina. *Lumbrerasaurus scagliai* is considered here as an extinct genus of Tupinambinae. These remains are the earliest record of teiid lizards and confirm the presence of tupinambine teiids in the South American Palaeogene.

Keywords: Squamata; Teiidae; Tupinambinae; Palaeogene; South America

Introduction

Palaeogene squamates of Argentina are represented by a number of snake vertebrae primarily from Patagonian localities (Simpson 1933, 1935; Hoffstetter 1959; Albino 1987, 1993, 1996, 2011, 2012); however, the Lumbrera Formation has provided lizard remains from the northwest of the country (Figure 1) (Donadío 1985; Albino 1993). The lizard material from this formation was first described by Donadío (1985), who named a new genus and species of Teiidae, Lumbrerasaurus scagliai. Later, Sullivan and Estes (1997) suggested that Lumbrerasaurus is inadequately diagnosed and considered it as a nomen dubium. Nevertheless, nobody has studied the fossil remains since Donadío (1985). A re-description and reevaluation of the remains belonging to Lumbrerasaurus scagliai are presented in this paper. They represent the only documented record of squamate reptiles from the Palaeogene of a non-patagonian site of Argentina.

Institutional abbreviation

AC, Cátedra de Anatomía Comparada, Departamento de Zoología, Universidad Nacional de Córdoba, Córdoba, Argentina; FML, Fundación Miguel Lillo, San Miguel de Tucumán, Argentina; MLP, Museo de La Plata, La Plata, Argentina; MMP, Museo Municipal de Ciencias Naturales 'Lorenzo Scaglia', Mar del Plata, Argentina; MNHNC, Museo Nacional de Historia Natural, Santiago, Chile; MZUSP, Museu de Zoologia, Universidade de São Paulo, São Paulo, Brazil; QCAZR, Museo de Zoología, Pontífica Universidad Católica del Ecuador, Pichincha, Ecuador; USNM, National Museum of Natural History, Smithsonian Institution, Washington DC, USA; UMMZ, University of Michigan Museum of Zoology, Ann Arbor, USA; UNMdP,

Colección Herpetológica Laboratorio de Vertebrados, Facultad de Ciencias Exactas y Naturales, Universidad Nacional de Mar del Plata, Mar del Plata, Argentina; UNMdP-O, Colección Herpetológica – Sección Osteología, Facultad de Ciencias Exactas y Naturales, Universidad Nacional de Mar del Plata, Mar del Plata, Argentina; UNNE, Universidad Nacional del Nordeste, Resistencia y Corrientes, Argentina.

Materials and methods

The biochronological sequence of South American Land Mammal Ages (SALMAs of Pascual et al. 1996) for the early Cenozoic of South America was recently revised (Gelfo et al. 2009; Woodburne et al. 2013) and provides an up-to-date scheme for the record of early Cenozoic squamates. Squamate systematics follow Gauthier et al. (2012). Fossils were confronted with osteological specimens of different institutional collections (Appendix 1) and bibliography. Fossils and specimens were analysed using a stereoscope microscope and photographed with a digital camera; scanning electron microscopy was also used.

Vertebral terminology follows Hoffstetter and Gasc (1969), whereas cranial terminology is based on Oelrich (1956) with additions/modifications of Montero et al. (2004), Rage and Augé (2010), Smith (2006) and Smith and Gauthier (2013). Denton and O'Neill (1995, p. 240–241) describe a dual septum in the maxilla of *Prototeius stageri* and mentioned the distribution of this character among Teiidae and Polyglyphanodontia (= Borioteiioidea Nydam et al., 2007). Although *Prototeius stageri* was originally considered a Polyglyphanodontia, its relationship with these lizards is uncertain (Nydam et al. 2010).

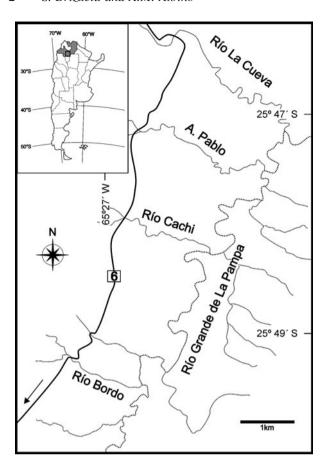


Figure 1. Location map of area from Lumbrera Formation from where fossils were recovered (modified from Carvajal et al. 1977). Salta Province in grey, square is the magnified area and arrow indicates the direction to Pampa Grande.

Independently, the dual septum has been recognised (i.e. coded in data matrix) in other Polyglyphanodontia (Bicuspidon, Cherminsaurus, Gilmoreteius, Polyglyphanodon) and some Teiidae (Callopistes, Dracaena) (Denton and O'Neill 1995; Nydam et al. 2007). Denton and O'Neill (1995) do not name the anterior and posterior buttresses that constitute their dual septum. For teiids, our interpretation is that the first of these two structures would correspond (but much modified) to the crista transversalis, and the second to a previously unnamed vertical structure developed medially on the nasal process whose dorsal end continues posteriorly as the nasolacrimal ridge (lacrimal crest of Oelrich 1956 = cresta semilunar of Montero et al. 2004) (Figure 2). The septomaxilla largely articulates on the crista transversalis and extends posteriorly to the posterior buttress. The question on what is the anterior buttress of the Polyglyphanodontia (very developed crista transversalis, fused septomaxilla or something else) is out of the scope of this contribution.

Geological setting

The Lumbrera Formation corresponds to the upper unit of the Santa Barbara subgroup (Salta Group). del Papa (2006) recognised, based on facies associations, two sections in this formation: inferior and superior. Contact between the Inferior Lumbrera Formation and the Superior Lumbrera Formation is represented by an unconformity that separates two sedimentary cycles (del Papa 2006). Within the Inferior Lumbrera Formation, there are two green levels: the inferior Faja Verde I and the superior Faja Verde II. Only the former is continuous throughout the Inferior Lumbrera Formation (del Papa 2006).

The mammalian fauna of the Lumbrera Formation was initially referred to the Casamayoran SALMA (Carvajal et al. 1977; Pascual et al. 1978, 1981), which Kay et al. (1999), based on isotopic and magnetic polarity data from Patagonia, correlate with the late Eocene. More recently, Deraco et al. (2008), based on the presence of a leontiniid (Mammalia, Notungulata), consider the red levels above Faja Verde II, as 'Barrancan'-Mustersan (mid late Eocene). These authors consider the levels below the Fajas Verdes I and II as chronologically equivalent to the Patagonian Vaquian (early Casamayoran). However, the faunas from north-west of Argentina show strong endemism at the generic and specific level, complicating the correlation with other faunas, especially with those of Patagonia, on which the Casamayoran is sustained (Vucetich et al. 2007). Marshall et al. (1997) consider that the lower part of the Lumbrera Formation correlates with chrons 25n and 26n, and the Fajas Verdes (I and II) correlate with chron 24r. Therefore, an age of 58-55.5 Ma for the lower part of the Lumbrera Formation is inferred. This would put the Lumbrera Formation in the late Palaeocene. Some consideration must be taken into account regarding this age because the palaeomagnetic data from the Lumbrera Formation were analysed when laboratory techniques were less rigorous than those applied later (see Sempere et al. 1997, p. 725). In this paper, the more accepted criteria of Vucetich et al. (2007) are followed, according to which the Lumbrera Formation is early Eocene in age.

The material described here was collected in a joint expedition of MLP and MMP in 1979. It was retrieved 5 m below Faja Verde I in a canyon between rivers Bordo and Toro near Pampa Grande, Salta province (Donadío 1985). All elements were closely associated. Reference to Ruta Nacional 9 in Donadío (1985) and Carvajal et al. (1977) would correspond to Ruta Provincial 6 that connects the localities of Pampa Grande and La Viña, as inferred from course of the road and geographical coordinates provided in Carvajal et al. (1977) (Figure 1).

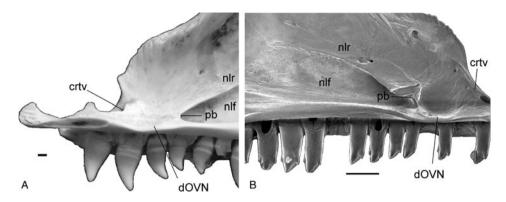


Figure 2. Antero medial anatomical features of Teiidae maxillae. (A) tupinambine *Tupinambis merianae* (UNMdP-O 75); (B) teiine *Ameiva ameiva* (UNMdP-O 27) crtv, *crista transversalis*; pb, posterior buttres; nlf, nasolacrimal fossa; nlr, nasolacrimal ridge; dOVN, depression for vomeronasal organ. Scale: 1 mm.

Systematic palaeontology

SQUAMATA Oppel, 1811
SCLEROGLOSSA Estes, de Queiroz and Gauthier, 1988
AUTARCOGLOSSA Wagler, 1830
SCINCOMORPHA Camp, 1923
TEIOIDEA Estes, de Queiroz and Gauthier, 1988
TEIIDAE Gray, 1827
TUPINAMBINAE Presch, 1974
Lumbrerasaurus scagliai Donadío, 1985

Holotype: MMP 1418, partial lizard skeleton, that includes fragments of an anterior (MMP 1418-1) and posterior (MMP 1418-2) left maxilla, the proximal end of a left epipterygoid (MMP 1418-3), three dorsal vertebrae (MMP 1418-4 a, 1418-6), one sacral vertebra (MMP 1418-7) and two caudal vertebrae (MMP 1418-8, 1418-9). All remains belong to a same specimen and were previously described by Donadío (1985).

(Figure 3)

Revised diagnosis: Lumbrerasaurus scagliai is a teiid lizard bearing a unique combination of characters but not those mentioned by Donadío (1985) as diagnostic. Lumbrerasarus scagliai is distinguishable from Teiinae teiids (Ameiva, Aspidoscelis, Cnemidophorus, Dicrodon and Kentropyx) in presenting a low, vague posterior buttress, and vertebrae with narrow, thick zygosphene, triangular outline of the neural canal and weak sagittal crest. The amount of cementum at the base of the teeth in Lumbrerasaurus differs from the more abundant deposition in the tupinambines Dracaena and Paradracaena. Lumbrerasaurus further differs from Dracaena, Paradracaena and Tupinambis (except Tupinambis quadrilineatus) in presenting labiolingually compressed posterior teeth. Unlike other tupinambinae bearing labiolingually compressed teeth (Crocodilurus and Tupinambis quadrilineatus), those of Lumbrerasaurus are more compressed distally.

Re-description: The anterior and posterior maxillary fragments (MMP 1418-1, -2) show subpleurodont tooth implantation. Tooth positions are delimited by interdental septa (transverse walls from the parapet to the supradental shelf) of alveolar bone. Attachment of the teeth to the bone is mainly through their bases, which are, in turn, covered by abundant cementum. The latter only partially conceals both the interdental septa and *sulcus dentalis*.

Specimen MMP 1418-1 (Figure 3(A),(B)) is the anterior part of a left maxilla, where the anterior end of the premaxillary process is not preserved. The lateral surface is smooth, without ornamentation, and bears three labial foramina at tooth positions (b), (c) and one more posterior to this last. The posteroventral part of the fenestra exonarina is preserved. The facial process rises gradually in lateral view. Its dorsal surface is slightly convex and extends from the anterior superior alveolar foramen in an anterolabial direction as a weak crista transveralis. Medially and anteriorly on the palatal shelf, the articulation surface for the vomer is preserved. Posterior to this articulation surface, but dorsal on the palatal shelf, is the depression that dorsally housed the vomeronasal organ (VNO). This depression is limited laterally by a subtle step that as consequence of the distortion of the facial process forms a narrow and deep groove. Separating the depression for the VNO from the nasolacrimal fossa is possible to distinguish a faint, low and very short convexity, the posterior buttress. Only the anterior-most part of the nasolacrimal crest is preserved, which forms a relatively acute angle with the palatal shelf, and ends anteriorly at the posterior buttress. Only a small portion of the nasolacrimal fossa is preserved below the crest. The posteromedial border of the palatal shelf is not preserved. In ventral view, the maxilla is wide, with the parapet taller than the crest of the supradental shelf. Only three tooth positions are preserved, in which two teeth are present with their apical ends missing. The anterior-most tooth, (a), is tall and labiolingually compressed (i.e. suboval in

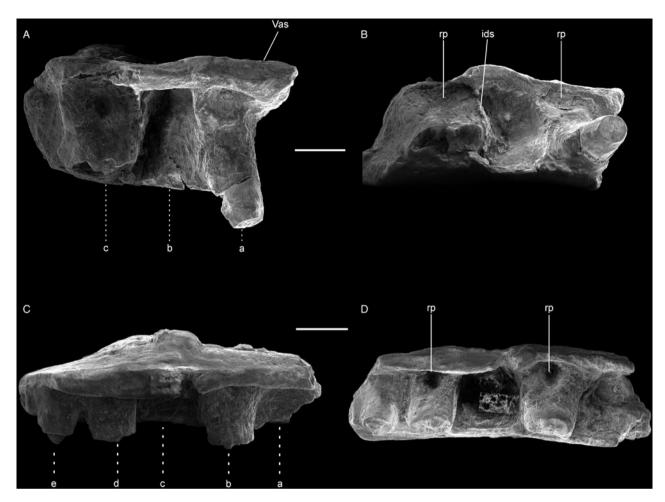


Figure 3. Cranial elements of the Tupinambinae *Lumbrerasaurus scagliai*, from the early Eocene (Lumbrera Formation) of north-west Argentina. (A, B) Anterior fragment of a maxilla (MMP 1418-1) in medial (A) and ventral (B) views; (C, D) posterior fragment of a maxilla (MMP 1418-2) in medial (C) and ventral (D) views. ids, interdental septum; rp, resorption pit; Vas, vomer articulation surface. Scale: 2 mm.

section), compression that is more accentuated on the distal end of the tooth (Figure 3(A),(B)). A very thin and weak crest crosses the distal face of the tooth, from the base to the apex where it is more evident (Figure 3(B)). The base of tooth (c) is larger than that of tooth (a), and it is labiolingually compressed as in the last one. Resorption zones are located lingually at the base of the teeth, and they develop into deep, suboval pits excavated in the cementum.

Fragment MMP 1418-2 (Figure 3(C),(D)) represents the posterior end of the same left maxilla. Its lateral surface is straight and preserves the last labial foramen at tooth position (a). Dorsally, the preservation is poor, but the facial process border is distinguishable. Medially, the superior alveolar canal is exposed and filled with sediment. Posteriorly, and poorly defined, are the superior alveolar foramen medial to the groove that housed the maxillary process of the jugal; this groove extends anteriorly clearly reaching tooth (d). In medial view, a poorly preserved palatal shelf is observed, upon which the

articulation surfaces for the palatine and ectopterygoid are preserved posteriorly. Ventrally is the dental series, medial to which develops the *sulcus dentalis*. As in MMP 1418-1, the parapet is taller than the crest of the supradental shelf. There are five tooth positions in which only three bases are preserved (positions b, d and e). Tooth bases are labiolingually compressed, more so distally. There is less cementum at the bases in this fragment than in MMP 1418-1. Two resorption pits are preserved lingual to teeth (b) and (d). The pits are more circular, somewhat smaller and deeper than those in MMP 1418-1.

MMP 1418-3 is a fragment of a long flat bone that probably corresponds to the ventral end of a left epipterygoid (Figure 4). The fragment is medially flat and laterally slightly convex. In lateral view, its ventral end is expanded, more pronounced anteriorly. The anterior lateral outline of this fragment is slightly concave. The posterior part of the lateral surface is higher, and we interpret it as the crest that crosses the epipterygoid dorsoventrally.

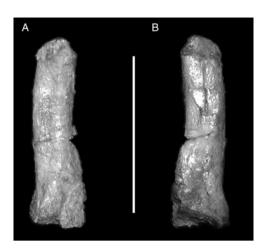


Figure 4. Epipterygoid fragment (MMP 1418-3) of the Tupinambinae *Lumbrerasaurus scagliai*. from the early Eocene (Lumbrera Formation) of north-west Argentina. (A) lateral view; (B) medial view. Scale: 10 mm.

The dorsal vertebra MMP 1418-4 (Figure 5(A),(B)) is large and distorted. The latter resulted in the obliteration of the neural canal due to the collapse of the neural arch onto the vertebral centrum. The condyle is missing, while the left transverse process is only partially preserved. The prezygapophyses are anterolaterally oriented, and their articulation surfaces are oval and steeply inclined. The anteromedial end of the neural arch is not preserved. Nonetheless, from the thick proximal (right) part, an important anterior extension is inferred. Although distorted, the transition between the prezygapophyseal and zygosphenal articulation surfaces is continuous. The neural spine is broken anteriorly, though its mark on the neural arch is visible on the anterior part. Though well-developed, the neural spine was probably low. Laterally, posterior to the prezygapophyses, are the transverse processes that distally bear oval, vertically orientated synapophyses. The vertebral centrum has a triangular outline, with a poorly developed sagittal crest along its entire ventral length. On either side of the sagittal crest, there are small foramina, which are interpreted as subcentral foramina. The cotyle is oval, wider than high and oriented anteroventrally.

The dorsal vertebra MMP 1418-5 (Figure 5(C),(D)) is also dorsoventrally distorted and somewhat smaller than the previous one. In this vertebra, the neural canal is displaced to the left of the centrum. The right prezygapophysis and postzygapophysis and the transverse processes are missing. The left prezygapophysis is large and anterolaterally oriented. Its articulation surface is poorly preserved. Medial to the prezygapophyses is the proximal part of the zygosphene, which is poorly preserved. The broken base of the neural spine extends along the entire dorsal surface of the neural arch. The

articulation surface of the left postzygapophysis is oval and posterolaterally oriented. The ventral outline of the vertebral centrum is triangular, ending in a marked precondylar constriction. The ventral surface of the centrum is similar to that described for vertebra MMP 1418-5 with a small subcentral foramen on the right. The cotyle is deep, transversally oval and ventrally inclined. The condyle is similarly oval and inclined.

The dorsal vertebra MMP 1418-6 is only partially preserved but includes the left side of the zygosphene (Figure 5(E),(F)). It extends anteriorly covering the cotyle in dorsal view, with the anterior margin notched. Due to distortion, the zygosphenal articulation surface is horizontal and parallel to that of the prezygapophysis. Transition between these surfaces is continuous as in MMP 1418-4. Only the anterior part of the neural spine is preserved, which is bent laterally because of distortion. The left prezygapophysis is large, orientated anterolaterally and bears a large, subcircular articulation surface. Posterior to the prezygapophysis is a well-defined interzygapophyseal crest. The transverse process is short and bears a vertically oriented, rectangular synapophysis. The cotyle is much deformed, but clearly deep.

Sacral vertebra MMP 1418-7 (Figure 5(G),(H)) does not preserve its dorsal part, except for the left prezygapophysis and the anterodorsal surface of the left pleurapophysis. The prezygapophysis is small and anteriorly orientated. Its articulation surface is subtriangular and almost vertically orientated because of distortion. In ventral view, the vertebral centrum is anteroposteriorly short and rectangular. A marked precondylar constriction precedes the condyle. At both sides of the centrum, there are the proximal parts of the pleurapophyses. These are strong and depressed. Ventrally they show shallow longitudinal grooves, but no ventral foramina are observed.

Caudal vertebra MMP 1418-8 preserves part of the left side and the vertebral centrum. The latter is long, cylindrical and constricted at the middle, with a well-developed precondylar constriction. Posterior to the cotyle, there are two subcentral foramina, of which the right one is larger. In lateral view, upon the neural arch, is a well-defined interzypapophyseal ridge. Between the neural arch and vertebral centrum is the mark of the transverse process/pleurapophysis, which was anteroposteriorly long and strongly depressed.

The caudal vertebra MMP 1418-9 preserves the right side and the vertebral centrum but not the condyle. The centrum is short but more robust than the one previously described. All other aspects are as in MMP 1418-8.

Fragment MMP 1418-10 could correspond to the posterior part of a dorsal vertebra. Part of the condyle is inferred as well as the right posteroventral border of the neural arch.

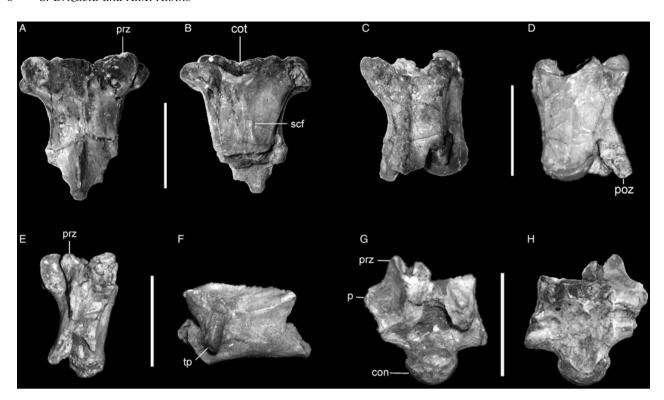


Figure 5. Vertebrae of the Tupinambinae *Lumbrerasaurus scagliai*. from the early Eocene (Lumbrera Formation) of north-west Argentina. (A, B) Dorsal vertebra (MMP 1418-4) in dorsal (A) and ventral (B) views; (C, D) dorsal vertebra (MMP 1418-5) in dorsal (C) and ventral (D) views; (E, F) dorsal vertebra (MMP 1418-6) in dorsal (E) and lateral (F) views; (G, H) sacral vertebra (MMP 1418-7) in dorsal (G) and ventral (H) views. con, condyle; cot, cotyle; p, pleurapophysis; poz, postzygapophysis; prz, prezygapophysis; scf, subcentral foramen; tp, transverse process. Scale: 10 mm.

Comparisons

From the observations of the articulation surface for the vomer on the maxilla MMP 1418-1, it is possible to affirm that maxilla and vomer meet at the anterior margin of the fenestra exochoanalis, which is a squamate synapomorphy (Rieppel et al. 2008; Gauthier et al. 2012). The nasolacrimal fossa, ventral to the nasolacrimal ridge, as observed in the fossil maxilla MMP 1418-1 is a synapomorphy of Autarchoglossa (Gauthier et al. 2012). The anteriorly extended maxillary process of the jugal (relative to the tooth row), as inferred from MMP 1418-2, allows the exclusion of the fossils from most Xantusiidae (except Lepidophyma flavimaculatum), Rhineuridae and the clade Aiolosaurus oriens + Varanoidea (Gauthier et al. 2012). The presence of a jugal can further exclude the fossils from the Krypteia because they lack this bone (Gauthier et al. 2012).

The subpleurodont tooth implantation is characteristic of the Teiidae, although it is known in members of other groups (e.g. the scincid *Tiliqua rugosa*, the amphisbaenian *Rhineura hatcherii*) (Gao and Fox 1996; Zaher and Rieppel 1999; Kearney et al. 2005; Nydam et al. 2007). The maxilla also shows lingual resorption zones, which corresponds to the iguanid method of tooth replacement

observed in most Scincomorpha, among others (see Edmund 1960, 1969). Resorption zones in Teiidae are not always lingual: they are sometimes found in a distolingual position (Edmund 1960, 1969). This variation is found at different levels, from generic, species, individual and even within the same specimen (Edmund 1960). Resorption pits of the fossils do not extend apically beyond the cementum. This type of 'low' resorption pit is observed among Teiidae. This contrast with the more apical developed pits observed in some pleurodont iguanids (e.g. *Tropidurus*) and seems to be related to implantation types.

The articulation surface for the palatine and ectopterygoid is wide, undifferentiated and embracing all the posterior end of the fossil maxilla MMP 1418-2. This allows the inference that the maxilla was excluded from the suborbital fenestra (Donadío 1985). This condition has long been recognised among Teiidae (although not exclusively) by various authors (e.g. Estes et al. 1988), further supporting affinity with them.

The Polyglyphanodontia, a Cretaceous Laurasian radiation of lizards, has for a long time been considered related to Teiidae (e.g. Estes 1983). This relationship between extinct Polyglyphanodontia and extant Teioidae (Teiidae and Gymnophthalmidae) [note the difference in

spelling, from the original of Estes et al. (1988) and emended in Gauthier et al. (2012)] is now considered to be more distant (e.g. Nydam et al. 2007; Nydam and Caldwell 2008; Gauthier et al. 2012). As mentioned before, Denton and O'Neill (1995) described a dual septum in the Prototeius (and noted in some Polyglyphanodontia) that limits anteriorly and posteriorly the depression for the VNO (on the palatal shelf). The anterior buttress is evident and medially expanded in Polyglyphanodontia (e.g. Bicuspidon see figure 2(A),(B) in Nydam and Cifelli 2002), whereas in the Teioidea, it is much reduced to the crista transversalis. Furthermore, the posterior buttress is located in different position among Teiidae and Gymnophthalmidae. While in Teiidae it is located posterior to the fenestra exonarina (in lingual view) (Figure 2), in Gymnophthalmidae, it is found just below, or much closer, to this fenestra (Bell et al. 2003, figure 4; Brizuela 2010). The morphology of the posterior buttress varies among Teiidae. In Teiinae the posterior buttress is marked, generally more developed dorsally (where it receives the septomaxilla), whereas in Tupinambinae, it is a low and vague ridge (Figure 2) (Brizuela 2010, figure 89). The morphology of the posterior buttress of MMP 1418-1 more closely resembles that of Tupinambinae. The dorsal morphology between the crista transversalis and posterior buttress, with a medial depression and a more lateral step, all much thinner than the posterior part of the palatal shelf, is also observed in tupinambine teiids (Figure 2(A)).

Regarding the postcranial elements, the morphology observed among the fossil vertebrae is like in extant Teiidae, although not exclusively (Hoffstetter and Gasc 1969). The shared characters are: procoelia, large size, well-developed zygosphene-zygantrum system with a developed tenon and notched anterior zygosphenal border, no prezygapophysial process, triangular vertebral centrum (in outline), dorsoventrally compressed cotyle and condyle, and marked precondylar constriction. Among Teiidae, several differences have been observed at subfamiliar level (Donadío 1985; Brizuela 2010; Albino and Brizuela 2014). In Tupinambinae there is a low, wide and vague ventral sagittal crest unlike the well-defined sagittal crest observed in Teiinae. In Tupinambinae, the neural canal is small and triangular in cross section, but it is relatively larger an oval in Teiinae. Finally, the zygosphene of adult Tupinambinae is narrower (in dorsal view) and thicker (in anterior view) than in Teiinae. The fossil shares the same characters observed in Tupinambinae. The large size of fossil vertebrae also favours Tupinambinae relationships. Small, subcentral foramina are observed on the dorsal vertebra (MMP 1418-4, contra Donadío 1985). These are smaller than those of the caudal vertebra (MMP 1418-8). Small foramina are common among dorsal vertebrae of extant teiids, although irregularly present. Donadío (1985) considered that the presence/absence of these foramina (on dorsal and caudal vertebrae) and the shape of the articulation surfaces of the zygapophyses have systematic value. These characters were later dismissed as non-diagnostic by Sullivan and Estes (1997). The variability present in these vertebral structures observed in the referred specimens (see Appendix 1), and the distortion of the fossils, support Sullivan and Estes's (1997) claim.

The maxilla fragments were originally recovered without the apical end of the teeth. Nonetheless, and based on the cross section of the preserved tooth bases, Donadío (1985) inferred a heterodont dentition in the fossil species. It could be argued that the subtle differences in the shape of the resorption pits could also correspond to a heterodont dentition. Unfortunately, the absence of the apical ends of functional or replacement teeth and the similarity of the tooth bases along the dental series make it impossible to verify Donadío's statement. Other aspects of the dentition are noteworthy. The suboval outline of the tooth bases, with a more marked distal constriction, is a character absent in other teiids, extant or extinct. This character was also dismissed by Sullivan and Estes (1997), but it seems unlikely to be an artefact because it is observed in all preserved tooth bases (Figure 3). The suboval outline of the fossil teeth differs from the more circular outline present in the tupinambines Dracaena and adult Tupinambis (except Tupinambis quadrilineatus). The subtle, but clear distal crest of tooth (a) of MMP 1418-1, seems to accompany this distal compression and would argue against artefactual compression. The dental characteristics of the fossils could be interpreted as similar to those observed in 'taxon 5' described by Carvalho (2001) for the Palaeogene of Itaboraí (Brazil) (Itaboraian SALMA early Eocene: Vucetich et al. 2007). 'Taxon 5' is referred by Carvalho (2001) to an unnamed new genus and species of Teiidae similar to the extant Crocodilurus. This fossil taxon is characterised by its heterodonty, which is composed of uni-, bi- and tricuspid teeth with cusps mesiodistally arranged. The particularity of this dentition is that the first accessory cusp that develops is the distal one, not the mesial one. Although Carvalho (2001) does not make reference to the cross section of the teeth, they appear to be oval based on his figures (Carvalho 2001, figures 24 and 25). The first appearance of the distal cusp could imply a more marked labiolingual constriction of the distal end of the tooth section, resulting in an outline similar to that of fossils studied here. However, in view of the insufficiency of the material from the Lumbrera Formation and the lack of a formal description of 'taxon 5', study of their possible relationships would be speculative.

Unlike Sullivan and Estes (1997), we feel that the Tupinambinae assignation of the studied specimen is supported by both cranial (subpleurodont implantation, lingual replacement zones with 'low' resorption pits and maxilla excluded from the suborbital fenestra) and

vertebral characters (developed zygosphene-zygantrum, precondylar constriction, among others), and that there is no justification to considered *Lumbrerasaurus scagliai* as a *nomen dubium* because it is characterised by an unique combination of characters.

Discussion

Previous reports concerning South American Palaeogene squamates include a great diversity of lizards and snakes from Brazil (Albino 1990; Rage 1998, 2001, 2008; Carvalho 2001), few representatives from Bolivia and Peru (Rage 1981, 1991), a marine palaeopheid snake from Ecuador (Hoffstetter 1958), a giant boid from Venezuela (Head et al. 2009) and a number of snakes from Argentina mainly distributed in Patagonian localities (Simpson 1933, 1935; Hoffstetter 1959; Albino 1987, 1993, 1996, 2011, 2012; Albino and Carlini 2008). The notable diversity of Palaeogene snakes from Patagonia, with at least one genus of madtsoiid (*Madtsoia*) and three genera of boids (*Boa*, Chubutophis and Waincophis), contrasts with the absence of lizards during the same period (Albino 2011). A similar pattern is observed in Patagonian Mesozoic deposits, where snakes are represented by more than six genera and continental lizards only by two taxa (Albino 2007, 2011). However, Cretaceous lizards from non-Patagonian localities are represented by interesting and diverse remains from various localities of Brazil (Estes and Price 1973; Evans and Yabumoto 1998; Candeiro et al. 2009; Nava and Martinelli 2011; Simões 2012; Simões et al. 2014). Also, the diverse but undescribed Itaboraian lizard fauna of Brazil (Carvalho 2001) demonstrates that the principal clades of lizards probably were diversified in South America at least by the early Eocene (Albino 2011). These deposits are also rich in snakes (Albino 1990; Rage 1998, 2001, 2008). In southern mid-latitudes, the fauna from Tiupampa in Bolivia (early Palaeocene, Gelfo et al. 2009) has provided a possible iguanian lizard and diverse snake taxa (de Muizon et al. 1983; Rage 1991). Early Eocene deposits of the Lumbrera Formation in north-western Argentina provide remains of a teiid lizard, corroborating that lizards and snakes were present in the Palaeogene of South America outside of Patagonia. It is not possible to know at present the reason for the disparity between Patagonian and non-Patagonian Palaeogene lizard occurrence. It could reflect the lack of lizards in southernmost Palaeogene environments, poorer samplings of small remains in Patagonian Palaeogene deposits, or insufficient conservation of fossil specimens for recognition as lizards.

As stated above, the South American Palaeogene includes a possible teiid lizard from the early Eocene of Brazil, although the material remains unpublished (Carvalho 2001). Teiid remains described in this paper suggest that this lizard group may have been broadly distributed from north-western Argentina to the Atlantic

coast of Brazil during the early Cenozoic. Also, the presence of tupinambine teiids in the Eocene indicates that the divergence of the two subfamilies of Teiidae (Teiinae and Tupinambinae) had already occurred at the beginning of the Cenozoic. This palaeontological conclusion is consistent with the molecular data indicating that the most basal tupinambine lineage (i.e. *Callopistes*) would have diverged during the Palaeocene (Giugliano et al. 2007).

Quattrocchio (2006) identified in the Lumbrera Formation a subtropical vegetation including elements of temperate regions. The teiid lizard studied in this paper, as well as the presence of sebecid crocodiles and pelomedusid turtles in the same sediments (Gasparini 1984; Gasparini et al. 1986), is consistent with these assumptions. Today, *Tupinambis rufescens* is the unique tupinambine teiid living in the area of Pampa Grande of the Salta province (Cei 1993).

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Note

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Appendix 1. Reference specimens consulted for this study.

ANGUIDAE: 'Ophiodes vertebralis', UNMdP 162. IGUA-NIDAE: 'Pristidactylus nigroiugulus', UNMdP-O 68; 'Tropidurus spinulosus', UNMdP-O 77; 'Tropidurus catalanensis', UNMdP-O 69. SCINCIDAE: 'Mabuya frenata', UNMdP 1599. GYMNOPHTHALMIDAE: 'Anadia bogotensis', UMMZ 192802; 'Bachia dorbignyi' UMMZ 172147; 'Cercosauria schreibersii', UNNE 8285, AC C119, C182, C253, C316, C318, C255, C081, C137, C155, C180, C181, C292; UNMdP-O 72, UNNE 8284, 'Vanzosaura rubricauda', UNMdP 649. TEIIDAE: 'Ameiva ameiva', UMMZ 149442, 176270, UNMdP 796, 1235, UNMdP-O 27, UNNE 1374, 1365, USNM 257529, 257530, 2414, 292419, 290898, 'Ameiva bifrontata', UMMZ 149443, 149444, 150782, 149449; 'Ameiva exsul', UMMZ 149448, 200865; 'Aspidoscelis cadilipes' UMMZ 149550, 149457, 149459, 149460; 'Aspidoscelis communis' UMMZ 149574; 'Aspidoscelis cozumela', UMMZ 149464; 'Aspidoscelis deppii', UMMZ 149551, 149552; 'Aspidoscelis exsanguis' UMMZ 138822, 149569; 'Aspidoscelis gularis', UMMZ 149553; 'Aspidoscelis sexlineata', UMMZ 149576, USNM 313460; 'Callopistes maculatus' USNM 4112, 4113; 'Cnemidophorus lacertoides', UNMdP-O 80; 'Cnemidophorus lemniscatus', UMMZ 189452, 'Cnemidophorus longicauda' UNMdP-O 51, 52, 81; 'Cnemidophorus ocelifer', UMMZ 149522; 'Cnemidophorus serranus', AC C371, UNMdP-O 82; 'Cnemidophorus tergolaevigatus', UNMdP-O 83; 'Crocodilurus amazonicus', MZUSP 8526, 12624, 92074; 'Dracaena guianensis', MZUSP 40082, USNM 220278; 'Dracaena paraguayensis', MZUSP 62849; 'Dicrodon guttulatum', QCAZR 2139; 'Kentropyx viridistriga', UNNE 4647, 1062; 'Kentropyx calcarata'. MZUSP 81580, 82978, USNM 292412; 'Teius teyou', UNMdP-O 15, 16, 22, 23, AC C365, UNNE 1398, 8715; 'Teius oculatus', AC C360, C368, C369, UNMdP-O 53, UNNE 4464, 4462; 'Teius suquiensis', AC C366, C367, UNMdP-O 84, 85; 'Tupinambis sp.', MZUSP 43042, 73309; 'Tupinambis duseni', MZUSP 92131; 'Tupinambis merianae', UMMZ 190109, UNMdP-O 1-14, 17-20, 24-26, 28, 32, 33, 35; 'Tupinambis quadrilineatus'. MZUSP 84921, 84928; 'Tupinambis rufescens', FML 7545, 7547, 1336, 2560, MZUSP 93082, 93083, 93084, UNMdP-O 36, USNM 342486; 'Tupinambis teguixin', MZUSP 92087, 92149, USNM 220279.