Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

Contrasting effects of acidification and warming on dimethylsulfide concentrations during a temperate estuarine fall bloom mesocosm experiment

Robin Bénard[1], Maurice Levasseur[1], Michael Scarratt[2], Sonia Michaud[2], Michel Starr[2], Alfonso Mucci[3], Gustavo Ferreyra[4,5], Michel Gosselin[4], Jean-Éric Tremblay[1], Martine Lizotte[1], Gui-Peng Yang[6]

[1]Département de biologie, Université Laval, 1045 avenue de la Médecine, Québec, Québec G1V 0A6, Canada
[2]Fisheries and Oceans Canada, Maurice Lamontagne Institute, P.O. Box 1000, Mont-Joli, Québec G5H 3Z4, Canada
[3]Department of Earth and Planetary Sciences, McGill University, 3450 University Street, Montréal, Québec H3A 2A7, Canada
[4]Institut des sciences de la mer de Rimouski (ISMER), Université du Québec à Rimouski, 310 allée des Ursulines, Rimouski, Québec G5L 3A1, Canada
[5]Centro Austral de Investigaciones Científicas (CADIC), Consejo Nacional de Investigaciones Científicas y Técnicas, Bernardo Houssay 200, 9410 Ushuaia, Tierra del Fuego, Argentina
[6]Institute of Marine Chemistry, Ocean University of China, 238 Songling Road, Qingdao 266100, Shandong, China

Correspondence: Robin Bénard (robin.benard.1@ulaval.ca)

Abstract. The effects of ocean acidification and warming on the concentrations of dimethylsulfoniopropionate (DMSP) and dimethylsulfide (DMS) were investigated during a mesocosm experiment in the Lower St. Lawrence Estuary (LSLE) in the fall of 2014. Twelve mesocosms covering a range of pH_T (pH on the total hydrogen ion concentration scale) from 8.0 to 7.2, corresponding to a range of CO_2 partial pressures (pCO_2) from 440 to 2900 µatm, at two temperatures (in situ and +5 °C; 10 °C and 15 °C) was monitored during 13 days. All mesocosms were characterized by the rapid development of a diatom bloom dominated by *Skeletonema costatum*, followed by its decline upon the exhaustion of nitrate and silicic acid. Neither the acidification nor the warming resulted in a significant impact on the abundance of bacteria over the experiment. However, warming the water by 5 °C resulted in a significant increase of the average bacterial production (BP) in all 15 °C mesocosms as compared to 10 °C, with no detectable effect of pCO_2 on BP. Variations in total DMSP ($DMSP_t$ = particulate + dissolved DMSP) concentrations tracked the development of the bloom although the rise in $DMSP_t$ persisted for a few days after the peaks in chlorophyll *a*. Average concentrations of $DMSP_t$ were not affected by acidification or warming. Initially low concentrations of DMS (< 1 nmol L^{-1}) increased to reach peak values ranging from 30 to 130 nmol L^{-1} towards the end of the experiment. Increasing the pCO_2 reduced the averaged DMS concentrations by 66 % and 69 % at 10 °C and 15 °C, respectively, over the duration of the experiment. On the other hand, a 5 °C warming increased DMS concentrations by an average of 240 % as compared to in situ temperature, resulting in a positive offset of the adverse pCO_2 impact. Significant positive correlations found between bacterial production rates and concentrations of DMS throughout our experiment point towards temperature-associated enhancement of bacterial DMSP metabolism as a likely driver for the mitigating effect of warming on the negative impact of acidification on the net production of DMS in the LSLE and potentially the global ocean.

Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

1. Introduction

Dimethylsulfide (DMS) is ubiquitous in productive estuarine, coastal and marine surface waters (Barnard et al., 1982; Iverson et al., 1989; Kiene and Service, 1991; Cantin et al., 1996; Kettle et al., 1999). With an estimated average 28.1 Tg of sulfur (S) being transferred to the atmosphere annually (Lana et al., 2011), DMS emissions constitute the largest natural source of tropospheric S (Lovelock et al., 1972; Andreae 1990; Bates et al., 1992). The oxidation of atmospheric DMS yields hygroscopic sulfate (SO_4^{2-}) aerosols that directly scatter incoming solar radiation and act as nuclei upon which cloud droplets can condense and grow, thereby potentially impacting cloud albedo and the radiative properties of the atmosphere (Charlson et al., 1987; Andreae and Crutzen 1997; Liss and Lovelock, 2007; Woodhouse et al., 2013). The scale of the impact of biogenic SO_4^{2-} particles on global climate, however, remains uncertain (Carslaw et al., 2010; Quinn and Bates, 2011, Quinn et al., 2017). The strength of DMS emissions depends on wind- and temperature-driven transfer processes (Nightingale et al., 2000) but mostly on its net production in the surface mixed layer of the ocean (Malin and Kirst, 1997). Net changes in the aqueous DMS inventory are largely governed by microbial food webs (see reviews by Simó, 2001; Stefels et al., 2007) whose productivity is potentially sensitive to modifications in the habitats that sustain them. Given the complexity of the biological cycling of DMS, understanding how climate change related stressors could impact the production of this climate-active gas is a worthy but formidable challenge.

DMS stems, for the most part, from the enzymatic breakdown of dimethylsulfoniopropionate (DMSP) (Cantoni and Anderson, 1956), a metabolite produced by several groups of phytoplankton, with an extensive range in intracellular quotas between taxa (Keller et al., 1989; Stefels et al., 2007). Several species of the classes Haptophyceae and Dinophyceae are amongst the most prolific DMSP producers, but certain members of Bacillariophyceae (diatoms) and Chrysophyceae can also produce significant amounts of DMSP (Stefels et al., 2007). The biosynthesis of DMSP is highly constrained by abiotic factors and its up- or down-regulation may allow cells to cope with environmental shifts in temperature, salinity, nutrients and light intensity (Kirst et al., 1991; Karsten et al., 1996; Sunda et al., 2002), while its de novo synthesis and exudation may also serve as a sink for excess carbon (C) and sulfur (S) under unfavourable growth conditions (Stefels, 2000). Beyond active exudation in healthy cells (Laroche et al., 1999), cellular or particulate DMSP ($DMSP_p$) can be transferred to the water column as dissolved DMSP ($DMSP_d$) through viral lysis (Hill et al., 1998; Malin et al., 1998), autolysis (Nguyen et al., 1988; Stefels and Van Boeckel, 1993), and grazing by micro-, meso- and macrozooplankton (Dacey and Wakeham, 1986; Wolfe and Steinke, 1996). The turnover rate of $DMSP_d$ in the water column is generally very rapid (a few hours to days) as this compound represents sources of C and reduced S for the growth of microbial organisms (Kiene and Linn, 2000). Heterotrophic bacteria mediate most of the turnover of S-$DMSP_d$ through pathways that constrain the overall production of DMS: (1) enzymatic cleavage of $DMSP_d$ that yields DMS; (2) demethylation/ demethiolation of $DMSP_d$ that yields methanethiol (MeSH); (3) production of dissolved non-volatile S compounds, including SO_4^{2-}, following oxidation of $DMSP_d$; (4) intracellular accumulation of $DMSP_d$ with no further metabolization (Kiene et al., 1999, 2000; Kiene and Linn, 2000; Yoch, 2002). A compilation of ^{35}S-$DMSP_d$ tracer studies conducted with natural microbial populations shows that microbial DMS yields rarely exceed 40% of consumed

$DMSP_d$ in surface coastal and oceanic waters (see review table in Lizotte et al., 2017). Another potential fate of $DMSP_d$ is its uptake by non-DMSP producing eukaryotic phytoplankton such as certain diatoms (Vila-Costa et al., 2006b; Ruiz-González et al., 2012) and cyanobacteria such as *Synechococcus* and *Prochloroccocus* (Malmstrom et al., 2005; Vila-Costa et al., 2006b), but the overall turnover of $DMSP_d$ seems to be dominated by heterotrophic organisms.

Whereas the role of bacteria in the production of DMS via $DMSP_d$ is well recognized, an increasing number of studies have shown that the phytoplankton-mediated enzymatic conversion of total DMSP ($DMSP_t$) into DMS can also be significant when communities are dominated by DMSP-lyase producing phytoplankton groups such as Dinophyceae and Haptophyceae (Niki et al., 2000; Steinke et al., 2002; Stefels et al., 2007; Lizotte et al., 2012), particularly under high doses of solar radiation (Toole and Siegel, 2004; Toole et al., 2006, 2008; Vallina et al., 2008). Removal processes of DMS from surface waters include photo-oxidation, bacterial degradation and efflux across the air-sea interface, the individual intensity of which depends on several factors such as light intensity, wind velocity, the depth of the surface mixed layer and the gross production of DMS (Brimblecombe and Shooter, 1986; Simó and Pedros-Alió, 1999; Nightingale et al., 2000; Hatton et al., 2004; Simó, 2004). Overall, production and turnover of DMS and its precursor DMSP are unequivocally linked with microbial activity, both autotrophic and heterotrophic. The associated biological processes and interactions amongst these microorganisms have been shown to be sensitive to fluctuations in abiotic factors and may thus be further modulated by multiple drivers of climate change.

Since the pre-industrial era, atmospheric CO_2 concentrations have risen from 280 ppm, and, according to the business-as-usual scenario RCP 8.5 and global ocean circulation models, are expected to reach 850–1370 ppm by 2100 (IPCC, 2013). The oceans have already absorbed about 28 % of the anthropogenic CO_2 emitted to the atmosphere (Le Quéré et al., 2015), leading to a pH decrease of 0.11 units in surface waters (Gattuso et al., 2015), a phenomenon called ocean acidification (OA). An additional decrease of pH by 0.3–0.4 units is expected by the end of this century, and could reach 0.8 units by 2300 (Caldeira and Wickett, 2005; Doney et al., 2009; Feely et al., 2009). In addition to the oceanic sink, a similar fraction of anthropogenic CO_2 emissions has been captured by terrestrial vegetation. while the anthropogenic CO_2 remaining (45% of total emissions) in the atmosphere (Le Quéré et al., 2013) has led to an estimated increased greenhouse effect of 0.3–0.6 W m^{-2} globally over the past 135 years (Roemmich et al., 2015). Ninety percent of this excess heat has been absorbed by the ocean, increasing sea surface temperatures (SST) ~0.1 °C per decade since 1951 and could increase SST by 3–5 °C before 2100 (IPCC, 2013). Leading experts in the field of global change have called upon the scientific community to address critical knowledge gaps, among which, a top priority remains the assessment of the impact of multiple environmental stressors on marine microorganisms (Riebesell and Gattuso, 2015).

The sensitivity of natural planktonic assemblages to OA, along with their production of DMSP and DMS, has been investigated in several experimental studies (see review table in Hussherr et al., 2017). The majority of these experiments have shown a decrease in both DMSP and DMS concentrations with increasing pCO_2 (Hopkins et al., 2010; Avgoustidi et al., 2012; Park et al., 2014; Webb et al., 2015). The decrease in DMSP production has largely been attributed to the deleterious impact of decreasing pH on the coccolithophore *Emiliania huxleyi*, the dominant DMSP producer in several of these studies. Nevertheless, OA does not always result in a concomitant decrease in DMSP and DMS production. For example, the pCO_2-

induced decrease in DMS reported by Archer et al. (2013) in Arctic waters was accompanied by an increase in DMSP concentrations, indicating that DMS production is at least partly dependent on the turnover of DMSP, rather than on the DMSP pool. A modeling study showed that the specific implementation of the negative effect of OA on DMS net production in a coupled ocean-atmosphere model reduces global DMS production by 18 ± 3 %, resulting in an additional warming of 0.23–0.48 K by 2100 under the A1B scenario (Six et al., 2013). Schwinger et al. (2017) further showed that the OA-induced decreases in oceanic DMS emissions could result in a transient global warming of 0.30 K, mostly resulting from a reduction of cloud albedo. These first attempts to model the potential effect of OA on climate through its impact on DMS oceanic production show that OA may significantly affect climate by reducing marine emissions of DMS but also highlight the importance of carefully assessing the robustness of the DMS-OA negative relationship. This is particularly relevant considering that some experiments reveal a neutral or positive effect of increasing pCO_2 on DMS net production (Vogt et al., 2008; Kim et al., 2010; Hopkins and Archer, 2014). Regional or seasonal differences in phytoplankton taxonomy, microzooplankton grazing, and bacterial activity have been proposed as key drivers of the discrepancies between these experimental results.

Whereas studies of the impact of OA on DMS cycling have gained momentum, the importance of assessing how combined drivers of change may impact the structure and the functioning of ocean ecosystems, using multifactorial approaches, is now increasingly recognized (Boyd et al., 2015; 2018; Riebesell and Gattuso, 2015; Gunderson et al., 2016). Thus far, only two mesocosm studies assessed the combined effect of OA and warming on DMS dynamics by natural plankton assemblages. The two studies, both conducted with coastal waters, led to contrasting results. The first study showed an 80 % increase in DMS concentrations under high pCO_2 conditions (900 ppm vs. 400 ppm), and a reduction by 20 % of this stimulating effect when the increase in pCO_2 was accompanied by a 3 °C warming (Kim et al., 2010). However, the absence of a specific stand-alone warming treatment did not allow the authors to assess the sole impact of temperature on DMS net production. The second study showed decreasing DMS concentrations under both acidification and greenhouse conditions, with the lowest DMS concentrations measured under combined acidification and warming treatments (Park et al., 2014). The authors attributed these contrasting responses to differences in the phytoplankton assemblages, DMSP-related algal physiological characteristics, and microzooplankton grazing. Nevertheless, questions remain as to the combined effect of pCO_2 and warming on DMS net production since the temperature treatments were not conducted over the full range of pCO_2 tested (Kim et al., 2010; Park et al., 2014).

The combined influence of acidification and warming on the dynamics of the St. Lawrence Estuary phytoplankton fall bloom was investigated during a full factorial mesocosm experiment (Bénard et al., 2018). During this experiment, a bloom of *Skeletonema costatum* developed in all mesocosms, independently of the pCO_2 gradient (from 440 to 2900 µatm) and temperatures tested (10 and 15 °C). The increase in pCO_2 had no influence on the bloom but warming accelerated the growth rate of the diatoms and hastened the decline of the bloom (Bénard et al., 2018). Here, we report on the impacts of acidification and warming on DMSP and DMS concentrations with a focus on the dynamics of heterotrophic bacteria, a component of the marine food web known to affect the turnover of DMSP and DMS.

2. Material and methods

2.1 Mesocosm setup

The mesocosm experimental setup is described in detail in Bénard et al. (2018). Briefly, mesocosm experiments were conducted at the ISMER marine research station of Rimouski (Québec, Canada) in the fall of 2014. The twelve 2600 L cylindrical (2.67 m × 1.4 m), conical bottom, mesocosms were housed in two temperature-controlled, full-size shipping containers each containing six mesocosms (Aquabiotech Inc., Québec, Canada). Each mesocosm are mixed by a propeller secured near the top of each enclosure to ensure homogeneous vertical mixing of the water column. The mesocosms are sealed by a Plexiglas cover transmitting 50–85 % of solar UVB (280–315 nm), 85–90 % of UVA (315–400 nm), and 90 % of photosynthetically active radiation (PAR; 400–700 nm) of the natural incident light. Independent temperature probes (AQBT-Temperature sensor, accuracy ± 0.2 °C) were installed in each mesocosm, recording temperature every 15 minutes and either triggering a resistance heater (Process Technology TTA1.8215) or a glycol refrigeration system activated by an automated pump. The pH of the mesocosms was measured every 15 minutes by Hach® PD1P1 probes (± 0.02 pH units) linked to Hach® SC200 controllers. To maintain pH, two reservoirs of artificial seawater were equilibrated with pure CO_2 before the start of the experiment and positive deviations from the target pH values in each mesocosm activated peristaltic pumps that injected the CO_2 supersaturated seawater into the mesocosm water. This control system was able to maintain the pH in the mesocosms within ± 0.02 pH units of the targeted values during the initial bloom development by lowering the pH, but it could not increase the pH during the declining phase of the bloom.

2.2 Experimental approach

Prior to the onset of the experiment, all the mesocosms were meticulously washed with diluted Virkon™, an anti-viral and anti-bacterial solution, according to the manufacturer's instructions (Antec International Limited), and thoroughly rinsed. The experimental approach is also detailed in Bénard et al. (2018). To fill the mesocosms, water from ~5 m depth was collected near the Rimouski harbour (48° 28' 39.9" N, 68° 31' 03.0" W) on the 27th of September 2014 (day -5). Initial conditions were: practical salinity (S_P) = 26.52, temperature = 10 °C, nitrate (NO_3^-) = 12.8 ± 0.6 µmol L^{-1}, silicic acid ($Si(OH)_4$) = 16 ± 2 µmol L^{-1}, and soluble reactive phosphate (SRP) = 1.4 ± 0.3 µmol L^{-1}. Following its collection, the water was screened through a 250 µm mesh while the mesocosms were simultaneously gravity-filled by a custom made "octopus" tubing system. The initial in situ temperature of 10 °C was maintained in all mesocosms for the first 24 h (day -4). On day -3, the six mesocosms in one of the containers were gradually heated to 15 °C while the mesocosms in the other container were maintained at 10 °C. No manipulations were carried on day -2 to avoid excessive stress, and acidification was carried out on day -1. The mesocosms were initially set to cover a gradient of pH_T (total proton concentration scale) of ~8.0 to 7.2 corresponding to a range of pCO_2 from 440 to 2900 µatm. Two mesocosms, one in each container (at each temperature), were not pH-controlled to assess the effect of freely fluctuating pH condition. These two mesocosms were called drifters since the in-situ pH was allowed to drift over time throughout the bloom development. To achieve the initially targeted pH_T, CO_2-

saturated artificial seawater was added to mesocosms M1, M3, M5, M7, M8, M10 (pH_T 7.2–7.6) while mesocosms M2, M4, M6, M9, M11, M12 (pH_T 7.8–8.0 and the drifters) were openly mixed to allow CO_2 degassing. Then, the automatic system controlling the occasional addition of CO_2-satured artificial seawater maintained the pH equal or below the targeted pH, except for the drifters.

2.3 Seawater analysis

Daily sampling of the mesocosms was carried out between 05:00 and 08:00 every day (EDT) as described in Bénard et al. (2018). Samples for carbonate chemistry, nutrients, DMSP and DMS were collected directly from the mesocosms prior to filling of 20 L carboys from which seawater for the determination of chlorophyll *a* (Chl *a*), bacterial abundance, and bacterial production (BP) was subsampled. Samples were collected directly from the mesocosms and the artificial seawater tank on days -3, 3 and 13 for practical salinity determinations. The samples were collected in 250 mL plastic bottles and stored in the dark until analysis was carried out on a Guildline Autosal 8400B salinometer in the months following the experiment.

2.3.1 Carbonate chemistry and nutrients

Analytical methods used to determine the carbonate parameters are described in detail in Bénard et al. (2018). Briefly, pH was determined every day by transferring samples from the mesocosms to 125 mL plastic bottles without headspace. The samples were analyzed within hours of collection on a Hewlett-Packard UV-Visible diode array spectrophotometer (HP-8453A) and a 5 cm quartz cell using phenol red (PR; Robert-Baldo et al., 1985) and *m*–cresol purple (mCP; Clayton and Byrne, 1993) as indicators after equilibration to 25.0 ± 0.1 °C in a thermostated bath. The pH on the total proton scale (pH_T) was calculated according to Byrne (1987), with the salinity of the sample and the HSO_4^- association constants given by Dickson (1990). The reproducibility of pH measurements, based on replicate measurements of the same samples and values derived from both indicators, was on the order of 0.003. Total alkalinity (TA) samples were collected every 3–4 days in 250 mL glass bottles to which a few crystals of $HgCl_2$ were added before sealing with ground glass stoppers and Apiezon® Type-M high-vacuum grease. The TA determinations were carried out within one day of sampling by open-cell automated potentiometric titration (Titrilab 865, Radiometer®) with a pH combination electrode (pHC2001, Red Rod®) and a dilute (0.025 M) HCl titrant solution calibrated against Certified Reference Materials (CRM Batch#94, provided by A. G. Dickson, Scripps Institute of Oceanography, La Jolla, USA). The average relative error, calculated from the average relative standard deviation on replicate standards and sample analyses, was < 0.15 %. The computed pH_T at 25 °C, measured TA, silicic acid and SRP concentrations were used to calculate the in situ pH_T, pCO_2 and saturation state of the water in each mesocosm using CO_2SYS (Pierrot et al., 2006) and the carbonic acid dissociation constants of Cai and Wang (1998).

The samples for the determination of NO_3^-, $Si(OH)_4$, and SRP were filtered through Whatman GF/F filters, collected in acid washed polyethylene tubes and stored at -20 °C. Analysis was carried out using a Bran and Luebbe Autoanalyzer III using the colorimetric methods of Hansen and Koroleff (2007). The analytical detection limit was 0.03 µmol L^{-1} for NO_3^- plus nitrite (NO_2^-), 0.02 µmol L^{-1} for NO_2^-, 0.1 µmol L^{-1} for $Si(OH)_4$, and 0.05 µmol L^{-1} for SRP.

2.3.2 Biological variables

Chl *a* determination methods are presented in Bénard et al. (2018). Succinctly, duplicate 100 mL samples were filtered onto Whatman GF/F filters. The filters were soaked in a 90 % acetone solution at 4 °C in the dark for 24 h, the solution was then analyzed by a 10-AU Turner Designs fluorometer (acidification method: Parsons et al., 1984). The analytical detection limit for Chl *a* was 0.05 µg L^{-1}.

Samples for the determination of free-living heterotrophic bacteria were kept in sterile cryogenic polypropylene vials and fixed with glutaraldehyde Grade I (final concentration = 0.5 %, Sigma Aldrich; Marie et al., 2005). Duplicate samples were placed at 4 °C in the dark for 30 min, then frozen at -80 °C until analysis by a FACS Calibur flow cytometer (Becton Dickinson) equipped with a 488 nm argon laser. Before enumeration, the samples were stained with SYBR Green I (0.1 % final concentration, Invitrogen Inc.) to which 600 µl of a Tris-EDTA 10 × buffer of pH 8 were added (Laboratoire MAT; Belzile et al., 2008). Fluoresbrite beads (diameter 1 µm, Polysciences) were also added to the sample as an internal standard. The green fluorescence of SYBR Green I was measured at 525 ± 5 nm. Bacterial abundance was determined as the sum of low and high nucleic (LNA and HNA) counts (Annane et al., 2015).

Bacterial production was estimated in each mesocosm on days 0, 2, 4, 6, 8, 10, 11 and 13 by measuring incorporation rates of tritiated thymidine (^{3}H-TdR), using an incubation and filtration protocol based on Fuhrman and Azam (1980, 1982). Twenty mL water subsamples were transferred from glass Erlenmeyers to five sterile glass vials; three as "measured" values and two as blanks. In all blank vials, 0.2 mL of formaldehyde 37 % were added, immediately after the sampling to stop all biological activities. Then, 1 mL of ^{3}H-TdR solution (4 µmol L^{-1}), prepared from commercial solution (63 Curie $mmol^{-1}$; 1 mCurie mL^{-1}, 10 µmol L^{-1} ^{3}H-TdR, MP Biomedicals), was added in all vials. Samples were incubated 2.5 h at experimental temperatures (10 or 15 °C), and then 0.2 mL of formaldehyde 37 % were immediately added in the three "measure" vials. Bacteria were then collected by filtration (diameter 25 mm; 0.2 µm porosity) and filters were treated according to Fuhrman and Azam (1980, 1982). ^{3}H-TdR incorporation was measured using a scintillation counter (Beckman LS5801) and results were expressed in dpm. Blank values were subtracted to "measured" values to remove background radioactivity. ^{3}H-TdR incorporation rates were converted in mole of ^{3}H-TdR incorporated per unit of volume and time, before converting to rate of carbon production using the carbon conversion factor of Bell (1993).

2.3.3 DMSP and DMS concentrations

For the quantification of $DMSP_t$, duplicate 3.5 mL samples of seawater were collected into 5 mL polyethylene tubes. Samples were preserved by adding 50 µL of a 50 % sulfuric acid solution (H_2SO_4) to the tubes before storage at 4 °C in the dark until analysis in the following months. Samples for the quantification of $DMSP_d$ were taken daily, but a technical problem during storage and transport of the samples led to a loss of all samples. The DMSP samples were injected into a purge and trap (PnT) system before being completely flushed using 1–5 mL Milli-Q™ water into the helium purged chamber heated to 70 °C. DMSP concentrations were determined by a mole to mole conversion to DMS following hydrolysis with a 5 M NaOH solution injected

in the chamber prior to the sample, and trapping the gas sample in a loop immersed in liquid nitrogen. The loop was then heated in a water bath to release the trapped sample and analyzed using a Varian 3800 Gas Chromatograph equipped with a pulsed flame photometric detector (PFPD, Varian 3800) and a detection limit of 0.9 nmol L^{-1} (Scarratt et al., 2000; Lizotte et al., 2012).

Samples for the quantification of DMS were directly collected from the mesocosms into 20 mL glass vials with a butyl septa and aluminum crimp. The samples were kept in the dark at 4 °C until analysis was carried out within hours of collection using the PnT system described above.

2.4 Statistical analyses

The statistical analyses were performed using the nlme package in R (R Core Team, 2016). The data were analyzed using a general least squares (gls) approach to test the linear effects of the two treatments (temperature, pCO_2), and their interaction on the variables (Paul et al., 2016; Hussherr et al., 2017; Bénard et al., 2018). The analyses were conducted on the averages of the measured parameters over the whole duration of the experiment, and separate regressions for pCO_2 were performed for each temperature when the latter had a significant effect. The residuals were checked for normality using a Shapiro-Wilk test ($p > 0.05$) and data were transformed (square root or natural logarithm) if necessary. In addition, squared Pearson's correlation coefficients (r^2) with a significance level of 0.05 were used to evaluate correlations between key variables.

3. Results

3.1 Physical and chemical conditions during the experiments

The S_p was 26.52 ± 0.03 on day -4 in all mesocosms and remained constant throughout the experiment, averaging 26.54 ± 0.02 on day 13 (Bénard et al., 2018). The temperature of the mesocosms in each container remained within ± 0.1 °C of the target temperature throughout the experiment and averaged 10.04 ± 0.02 °C for mesocosms M1 through M6, and 15.0 ± 0.1 °C for mesocosms M7 through M12 (Fig. 1a). The pH_T remained relatively stable throughout the experiment in the pH-controlled treatments, but decreased slightly as the experiment progressed, deviating by an average of -0.14 ± 0.07 units relative to the target pH_T on the last day (Fig. 1b). The pH variations corresponded to changes in pCO_2 from an average of 1340 ± 150 µatm on day -3, and ranged from 564 to 2902 µatm at 10 °C and from 363 to 2884 µatm at 15 °C on day 0 following the acidification (Fig. 1c). The in situ pH_T in the drifters (M6 and M11) increased from 7.896 and 7.862 on day 0, at 10 °C and 15 °C respectively, to 8.307 and 8.554 on day 13, reflecting the balance between CO_2 uptake and metabolic CO_2 production over the duration of the experiment. On the last day, pCO_2 in all mesocosms ranged from 186 to 3695 µatm at 10 °C and from 90 to 3480 µatm at 15 °C.

Nitrate (NO_3^-) and silicic acid ($Si(OH)_4$) concentrations averaged 9.1 ± 0.5 µmol L^{-1} and 13.4 ± 0.3 µmol L^{-1} on day 0, respectively (Bénard et al., 2018). The two nutrients displayed a similar temporal depletion pattern following the development of the phytoplankton bloom. NO_3^- concentrations reached undetectable levels (< 0.03 µmol L^{-1}) in all mesocosms by day 5.

Likewise, $Si(OH)_4$ fell below the detection limit (< 0.1 µmol L^{-1}) between day 1 and 5 in all mesocosms except for those whose pH_T was set at 7.2 and 7.6 at 10 °C (M5 and M3) and in which $Si(OH)_4$ depletion occurred on day 9.

3.2 Phytoplankton, bacterial abundance and production

Chl *a* concentrations were below 1 µg L^{-1} following the filling of the mesocosms (day -4), and had already increased to an average of 5.9 ± 0.6 µg L^{-1} on day 0 (Fig. 2a). At 10 °C, Chl *a* quickly increased to reach maximum concentrations around 27 ± 2 µg L^{-1} on day 3 ± 2, and decreased progressively until the end of the experiment. Increasing the temperature by 5 °C resulted in a more rapid development of the bloom and a speedier decrease of Chl *a* concentrations during the declining phase of the bloom. The maximum Chl *a* concentration reached at the peak of the bloom was, however, not significantly affected by the difference in temperature. We found no significant effect of the pCO_2 gradient on the mean Chl *a* concentrations measured over the days 0–13, nor during the development phase and the declining phase of the bloom as described in Bénard et al. (2018) (Fig. 2a–b; Table 1).

The free-living bacterial abundance was ~1.2×10^9 cells L^{-1} on day -4, and increased rapidly to reach $3.1 \pm 0.6 \times 10^9$ cells L^{-1} on day 0 (Fig. 2c). This initial increase in abundance probably resulted from the release of dissolved organic matter (DOM) during pumping of the seawater and filling of the mesocosms. The subsequent decrease in bacterial abundance during the development phase of the bloom suggests that the initial pool of DOM was fully utilized and that freshly released DOM was scarce. As expected, bacterial abundance increased during the declining phase of the bloom at 10 °C. Under warmer conditions, bacterial abundance decreased earlier during the initial bloom development than what was observed at 10 °C, but was also marked by an earlier peak during the decline of the bloom, then was followed by a second, more variable peak in abundance. These daily variations in abundances probably reflect changes in the balance between bacterial growth and loss by grazing. When averaged over the experiment, we observed no effect of the treatments on the mean bacterial abundance (Fig. 2c–d; Table 1). At 10 °C, bacterial production was low at the beginning of the experiment and increased gradually during the development and declining phases of the bloom to reach peaks values of 9.3 ± 0.9 µg C L^{-1} d^{-1} (Fig. 2e). Bacterial production increased faster at 15 °C and reached maximal production rates of 19 ± 1 µg C L^{-1} d^{-1} on day 11. Results of the gls model show no effect of the pCO_2 gradient on bacterial production, but a positive effect of warming was observable throughout the experiment (Fig. 2f; Table 1).

3.3 $DMSP_t$ and DMS

At in situ temperature, $DMSP_t$ concentrations averaged 9 ± 2 nmol L^{-1} on day 0 and increased regularly in all mesocosms up to day 10 before they plateaued or slightly decreased over the last 2–3 days (Fig. 3a). These results reveal that DMSP accumulation persisted for several days after the bloom peaks, to reach a maximum value between days 8–13 of 366 ± 22 nmol L^{-1}. At 15 °C, $DMSP_t$ concentrations similarly increased after the maximum Chl *a* concentrations were reached, but increased faster than at in situ temperature. The maximum $DMSP_t$ concentrations were 396 ± 19 nmol L^{-1} at 15 °C, a value

that is not statistically different from the peak values measured at 10 °C (Fig 4a; Table 2). A greater loss of DMSP took place in the last days of the experiment at 15 °C. By day 13, 79 ± 3 % of the peak $DMSP_t$ concentration was lost in the 15 °C mesocosms, while 19 ± 4 % of the peak $DMSP_t$ concentration was lost at 10 °C. When averaged over the duration of the experiment, the mean $DMSP_t$ concentrations were not significantly affected by the pCO_2 gradient, the temperatures or the interaction between these two factors (Fig. 3b; Table 1).

Over the 13 days, the $DMSP_t$:Chl *a* ratio averaged 11.4 ± 0.4 nmol (µg Chl *a*)$^{-1}$ at 10 °C and was not affected by increasing pCO_2 (Fig. 5a; Table 1). Due to the aforementioned mismatch between the peaks in Chl *a* and $DMSP_t$, the average $DMSP_t$:Chl *a* ratios were significantly higher at 15 °C, averaging 19 ± 1 nmol (µg Chl *a*)$^{-1}$ over the experiment (Fig. 5a; Table 1). However, we found no significant relationship between $DMSP_t$:Chl *a* and the pCO_2 gradient.

Initial DMS concentrations were below the detection limit on day 0 (< 0.9 nmol L^{-1}) and slowly increased during the first 7 days, while most of the build-up took place after day 8 in all treatments (Fig. 3b). The net accumulation of DMS was faster at 15 °C than at 10 °C, with higher daily DMS concentrations at 15 °C compared to 10 °C from day 3 until day 13. At the end of the experiment, DMS concentrations averaged 21 ± 4 nmol L^{-1} at 10 °C and 74 ± 14 nmol L^{-1} at 15 °C. Over the full duration of the experiment, we found significant negative effects of increasing pCO_2 on mean DMS concentrations at the two temperatures tested (Fig. 3c; Table 1). At 10 °C, we measured a ~67 % reduction of mean DMS concentrations from the drifter relative to the most acidified treatment (~345 ppm vs ~3200 ppm), with values decreasing from 10 ± 2 nmol L^{-1} to 3.2 ± 0.8 nmol L^{-1}. At 15 °C, the mean DMS concentrations decreased by roughly the same percentage (~69 %) as pCO_2 increased from the drifter to the most acidified treatment (~130 ppm vs ~3130 ppm). Nevertheless, the mean DMS concentrations were higher at 15 °C, ranging from 34 ± 13 nmol L^{-1} to 11 ± 3 nmol L^{-1} (Fig. 3c; Table 1). Similarly, the peak DMS concentrations decreased linearly with increasing pCO_2 at both temperatures and concentrations were always higher at 15 than at 10 °C for any given pCO_2 (Fig. 4b; Table 2).

The DMS:Chl *a* ratios remained below 1 nmol (µg Chl *a*)$^{-1}$ during the first 8 days in all mesocosms as DMS concentrations were low, but increased exponentially at 15 °C in the following days. At 10 °C, the DMS:Chl *a* ratio averaged 0.43 ± 0.7 nmol (µg Chl *a*)$^{-1}$ over the 13 days and was not affected by the pCO_2 gradient. At 15 °C, the DMS:Chl *a* ratios were not significantly affected by the pCO_2 gradient, but were significantly higher in the warmer treatment (Fig. 5b; Table 1).

The DMS:$DMSP_t$ ratio exhibited the same general pattern as the DMS, i.e. low and stable values during the first 8 days, and increasing values between days 8–13 (Fig. 3e). The natural logarithm of the DMS:$DMSP_t$ ratio was not affected by the pCO_2 gradient at 10 °C when averaged over the 13 days experiment, but a significant decrease of the DMS:$DMSP_t$ ratios was observed with increasing pCO_2 at 15 compared to 10 °C (Fig. 3f; Table 1).

4. Discussion

4.1 General characteristics

As far as we know, this study is the first full factorial mesocosm experiment where all pCO_2 treatments (pH_T from 8.0 to 7.2) were replicated at two different temperatures (in situ and +5 °C), to assess the impact of ocean acidification and warming on the dynamics of DMSP and DMS concentrations during a phytoplankton bloom. A diatom bloom dominated by *Skeletonema costatum* developed in all mesocosms, regardless of the treatments. This chain-forming centric diatom is a cosmopolitan species in coastal and estuarine systems and a frequent bloomer in the Lower St. Lawrence Estuary (LSLE) (Kim et al., 2004; Starr et al., 2004; Annane et al., 2015). The 13 days where treatments were applied allowed us to capture the development and declining phases of the bloom. The impacts of the treatments on the dynamics of the bloom during these two phases are described in greater details in Bénard et al. (2018). Briefly, the acidification had no detectable effect on the development rate of the diatom bloom and on the maximum Chl *a* concentrations reached. However, increasing the water temperature by 5 °C increased the growth rate of the diatoms, shortening the development phase of the bloom, from 4–7 days at 10 °C to 1–4 days at 15 °C. However, these changes in the bloom timing did not alter the overall primary production throughout the experiment. Hereafter, we discuss how increasing pCO_2 (lowering the pH) affected DMSP and DMS concentrations and how a 5 °C increase in temperature altered the impacts of the pCO_2 gradient during the experiment.

4.2. DMSP dynamics

The buildup of the phytoplankton biomass during the bloom development was coupled with a rapid increase in $DMSP_t$ concentrations (Fig. 3a). Assuming that *S. costatum* was responsible for most of the DMSP production, our results indicate a low sensitivity of the DMSP synthesis pathway to acidification in this species. The net accumulation of $DMSP_t$ persisted several days after the peaks in Chl *a*, indicating a decoupling between DMSP synthesis, algal growth and nitrogen metabolism (Bénard et al., 2018).

4.2.1 Effects of acidification on DMSP

At in situ temperature, the averaged $DMSP_t$ concentrations were not affected by the increase in pCO_2 (Fig. 3b; Table 1). The lack of significant changes in the $DMSP_t$:Chl *a* ratio as a function of the pCO_2 gradient also supports this conclusion (Fig. 5a; Table 1). This result is consistent with those of previous studies that showed a relatively weak effect of an increase in pCO_2 on DMSP concentrations (Vogt et al., 2008; Lee et al., 2009; Avgoustidi et al., 2012; Archer et al., 2013; Webb et al., 2015). Furthermore, much like the patterns observed at 10 °C, there was no relationship between the concentrations of $DMSP_t$ and the pCO_2 gradient observable at 15 °C (Table 1).

4.2.2 Effects of warming on DMSP

In contrast to the absence of effects of acidification on DMSP, warming has been previously shown to affect DMSP concentrations in nature. Results from shipboard incubation experiments conducted in the North Atlantic have revealed an increase in particulate DMSP ($DMSP_p$) concentrations due to a 4 °C warming (Lee et al., 2009). During this last study, the higher $DMSP_p$ concentrations were attributed to a temperature-induced shift in community structure toward species with higher cellular DMSP content. During our study, the pCO_2 and temperature treatments did not alter the structure of the community (Bénard et al., 2018). Most of the DMSP synthesis was likely linked to the numerically dominant diatoms, as all other algal groups identified contributed to less than 10 % of the total algal abundance (see Fig. 6 in Bénard et al., 2018). Our results thus suggest that DMSP synthesis by *S. costatum* during the nitrate-replete growth phase was not significantly affected by warming. Rather, it is the accelerated growth rate of *S. costatum* that promoted the concurrent accumulation of biomass and $DMSP_t$, while the higher $DMSP_t$:Chl *a* ratio observable at 15 °C may be explained by the faster degradation of cells under warming. Several empty frustules were found during the last days of the experiment at 15 °C, suggesting a loss of integrity of the cells and potential increase of the release of intracellular dissolved organic matter, including DMSP. The increase in the abundance of bacteria and in bacterial production (Fig. 2c, e) during that period also suggest that more dissolved organic matter was produced during the decline of the bloom, as previously reported (Engel et al., 2004a, 2004b). During our experiment, transparent exopolymer particles (TEP) concentrations increased during this period (Gaaloul, 2017), adding to the evidence for heightened DOM production by the decaying bloom, with a potential increase in DMSP metabolization by heterotrophic bacteria under warming.

4.3 DMS dynamics

DMS concentrations remained very low during the development phase of the bloom (day 8) and increased in the latter days of the experiment. Most of the DMS accumulation in the mesocosms took place between days 8–13 and likely originated from DMSP that may have been released during cell lysis (Kwint and Kramer, 1995), or upon zooplankton grazing (Cantin et al., 1996). Unbalanced growth and photosynthesis of algal cells under nitrogen deficiency during that period may also be responsible for a greater production and active exudation of DMSP (Stefels et al., 2000; Kettles et al., 2014).

4.3.1 Effects of acidification on DMS

At in-situ temperature, we observed a significant linear decrease in DMS concentrations (both averaged over the full duration of the experiment and peak concentrations) with increasing pCO_2 (Figs. 3c, 4b; Tables 1 and 2). Several earlier mesocosm experiments have shown similar decreasing trends of DMS concentrations with increasing pCO_2 (Hopkins et al., 2010; Archer et al., 2013; Park et al., 2014; Webb et al., 2015, 2016). In these studies, the pCO_2-induced decreases in DMS were generally attributed to changes in the microbial community speciation and structure, or to microzooplankton grazing, although decreases in bacterial DMSP-to-DMS conversion or increases in DMS consumption have also been suggested (Archer et al., 2013;

Hussherr et al., 2017). During our study, the decrease in DMS concentrations with increasing pCO_2 cannot be directly attributed to a decrease in $DMSP_t$ since this pool was not affected by the pCO_2 gradient (Figs. 3b, 4a; Tables 1 and 2). In Park et al. (2014), the increase in pCO_2 led to the reduction in the abundance of *Alexandrium* spp., an active DMSP and DMSP-Lyase (DLA) producer, and a concomitant reduction of the associated microzooplankton grazing. As *Alexandrium* spp. was less numerous, the associated attenuation of microzooplankton grazing resulted in a reduction of the mixing of DMSP and DLA, leading to lesser DMSP-to-DMS conversion. Given the strong contribution of *S. costatum* to the bloom, a species with no reported DLA, it can be assumed that most, if not all, of the DMS produced was driven by bacterial processes following DMSP release by the diatoms. Thus, the decrease in DMS concentrations in our study could have been the result of altered bacterial mediation; either through reduced bacterial production of DMS or heightened bacterial consumption of DMS. Whereas a reduction in bacterial uptake of DMSP is unlikely, given that the bacterial abundance and production were unaffected by the pCO_2 gradient (Table 1), the observed decrease in DMS concentrations could imply that at higher pCO_2 the bacterial yields of DMS are abated. The relative proportion of DMSP consumed by bacteria and further cleaved into DMS is closely tied to bacterial demand in carbon and sulfur as well as to the availability of DMSP relative to other sources of reduced sulfur in the environment (Levasseur et al., 1996; Kiene et al., 2000; Pinhassi et al., 2005). The absence of a significant pCO_2 effect on the concentrations of DMSP during this study may be interpreted as a pCO_2-related alteration of the microbially-mediated fate of consumed DMSP. Unfortunately, in the absence of detailed ^{35}S-$DMSP_d$ bioassays, it is impossible to confirm the outcome of the DMSP metabolic pathways including the DMSP-to-DMS conversion efficiency in relation to the pCO_2 gradient. A few studies (Grossart et al., 2006; Engel et al., 2014; Webb et al., 2015;) have reported enhanced bacterial abundance and production at high pCO_2, especially for attached bacteria as opposed to free-living (Grossart et al., 2006). However, regardless of the temperature treatment, neither the abundance nor the activity of bacteria seemed to be significantly impacted by pCO_2 in this study. A pCO_2-induced increase in bacterial DMS turnover could also explain the decrease in DMS concentrations, but several studies suggest that bacterial DMS consumption in natural systems is often tightly coupled to DMS production itself (Simó, 2001, 2004). Furthermore, while one laboratory study reported that non-limiting supplies of DMS may be used as a substrate by several members of Bacteroidetes (Green et al., 2011), another study showed that only a subset of the natural microbial population may turnover naturally-occurring levels of DMS (Vila-Costa et al., 2006b). Nevertheless, the sensitivity of these DMS-consuming bacteria to decreasing pH remains unknown. Likewise, whereas we cannot exclude a potential impact of pCO_2 on DMS turnover via bacterioplankton, it is plausible that the pCO_2 gradient may have affected a widespread physiological pathway among bacteria, specifically, the metabolic breakdown of DMSP.

4.3.2 Effects of warming on DMS

A warming by 5 °C increased DMS concentrations at all pCO_2 tested, resulting in an offset of the negative pCO_2 impact when compared to the in situ temperature. This result differs from the observation of Kim et al. (2010) and Park et al. (2014) in two ways. First, our results show an increase in DMS concentrations in the warmer treatment while the two previous studies reported a decrease. Second, our results confirm that a temperature effect may be measured over a large range of pCO_2. It is

noteworthy that the increase in DMS concentrations at the two temperatures tested varied from 110 % at pH 8.0 up to 370 % at pH 7.4. This highlights the scaling of the temperature effect over an extensive range of pCO_2 and the importance of simultaneously studying the impact of these two factors on DMS production. As observed at 10 °C, both the average and the peak DMS concentrations decreased linearly as pCO_2 increased in the warm treatment (Figs. 3d, 4b; Tables 1 and 2). Nevertheless, the pCO_2-induced decrease in DMS concentrations at 15 °C cannot be directly attributed to a decrease in $DMSP_t$ concentrations given that an increase in pCO_2 had no discernable effect on $DMSP_t$ concentrations. In contrast to our observations at the in situ temperature, where $DMSP_t$ continued to increase until day 12, $DMSP_t$ concentrations typically decreased from day 8 and onward (Fig. 3a). This loss in $DMSP_t$ suggests that microbial consumption of DMSP exceeded DMSP algal synthesis. In light of the dominance of *S. costatum*, a phytoplankton taxon not known to exhibit DLA, the bulk of microbial DMSP mediation was likely associated with heterotrophic bacteria. In support of this hypothesis, the bacterial production was ~2 times higher at 15 than at 10 °C between days 8–13 (19 ± 1 µg C L^{-1} d^{-1} vs 9.3 ± 0.9 µg C L^{-1} d^{-1}) (Fig. 2), and we observed a significant correlation between the quantity of $DMSP_t$ lost during Phase II (day of the $DMSP_t$ peak concentration to day 13) and the quantity of DMS produced during the same period (coefficient of determination, $r^2 = 0.60$, $p < 0.01$, $n = 11$). Assuming that all the $DMSP_t$ lost was transformed into DMS by bacteria, we calculated that DMS yields could have varied by 0.5 to 32 % across the pCO_2 gradient (mean = 13 ± 11 %). These very rough estimates of DMS yields are likely at the lower end since measured DMS concentrations also reflect losses of DMS through photo-oxidation and bacterial consumption. Nevertheless, we cannot exclude the possibility of some passive uptake of DMSP by the picocyanobacterial population in the mesocosms, although this pathway is not considered to be significant in natural systems (Malmstrom et al., 2005; Vila-Costa et al., 2006a) and does not lead to the production of DMS. Our 'minimum community' DMS yield estimates agree with an expected range of microbial DMS yields in natural environments, from 2 % to 45 % (see review table in Lizotte et al., 2017). These gross but realistic estimates of heterotrophic bacterial DMSP-to-DMS conversions could explain the bulk of the DMS present in our study, a hypothesis also supported by the strong positive correlation ($r^2 = 0.64$, $p < 0.001$, $n = 70$) between overall DMS concentrations and bacterial production rates. Combined, these findings reinforce the idea that bacterial metabolism, rather than bacterial stocks, may significantly affect the fate of DMSP (Malmstrom et al., 2004a, 2004b, 2005; Vila et al., 2004; Vila-Costa et al., 2007; Royer et al., 2010; Lizotte et al., 2017) and that drivers of environmental change, such as temperature and pH, that can alter bacterial activity and strongly impact the gross and net production of DMS. Specific measurements of bacterial DMSP uptake and DMS yields using ^{35}S-$DMSP_d$ should be conducted to assess the impacts of pCO_2 and temperature on the microbial fate of DMSP.

4.4 Limitations

During our study, the pCO_2 changes were applied abruptly, over a day, from in situ values to pCO_2 levels exceeding the most pessimistic pCO_2 scenarios for the end of the century. Compared to our manipulation, ocean acidification will proceed at a slower temporal scale, potentially allowing species to adapt and evolve to these changing conditions (Stillman and Paganini, 2015; Schlüter et al., 2016). However, in the LSLE, the upwelling of low oxygenated waters can rapidly reduce the pH_T to

~7.62, or even lower with contributions of low pH_T (7.12) freshwaters from the Saguenay River during the spring freshet (Mucci et al., 2017). Thus, the swift and extensive pCO_2 range deployed in our experiment may seem improbable for the open ocean on the short term, but may not be inconceivable for this coastal region. However, the warming of 5 °C used in this mesocosm study possibly exceeds the upper limit of temperature increase for the end of the century in this region. In the adjacent Gulf of St. Lawrence (GSL), surface waters temperature (SST) correlates strongly with air temperature, allowing the estimation of past SST. This relationship showed that SST has increased in the GSL by 0.9 °C per century since 1873 (Galbraith et al., 2012), although additional positive anomalies of 0.25–0.75 °C per decade have been shown between 1985 and 2013 (Galbraith et al., 2016). In the LSLE, the highest temperatures occur at the end of summer / early fall, and gradually dissipate by heating the subjacent cold intermediate layer through vertical mixing (Cyr et al., 2011). The extent of the projected warming in the LSLE is unknown, but will likely result from the multifaceted interactions between heat transfer from the air and physical factors controlling the water masses.

The results from our study could also be influenced by the absence of macrograzers in the mesocosms. An additional grazing pressure could limit the growth of the blooming species, reducing the amount of DMSP produced or could increase the release of $DMSP_d$ through sloppy feeding after the initial bloom (Lee et al., 2003). It is unclear how an increase in grazing pressure would have impacted the concentrations of DMS in our experiment. On the one hand, increased predation could have limited the net accumulation of $DMSP_p$, with a possible reduction in DMS production. On the other hand, increased grazing could have favoured the release of $DMSP_p$ as DMSPd, thus increasing the availability of this substrate for microbial uptake, mediation and possible conversion into DMS. Despite the absence of reported changes in community composition in our study, many OA mesocosm experiments have described changes in DMS concentrations associated with shifts in community structure in the past (Vogt et al., Hopkins et al.,2010; Kim et al., 2010, Park et al., 2014, Webb et al., 2015). Nonetheless, our results align with those of other OA studies (Archer et al., 2013; Hussherr et al., 2017), suggesting that the mediation of heterotrophic bacteria plays a major role in DMS cycling in the absence of reported phytoplanktonic DLA, such as in a diatom dominated bloom in the LSLE.

5. Conclusions

The objective of this study was to quantify the combined impact of increases in pCO_2 and temperature on the dynamics of DMS during a fall diatom bloom in the St. Lawrence Estuary. Our mesocosm experiment allowed us to capture the development and declining phases of a bloom strongly dominated by the diatom *Skeletonema costatum* and the related changes in bacterial abundance and production. As expected, warming accelerated the development of the bloom, but also its decline. Both $DMSP_t$ and DMS concentrations increased during the development phase of the bloom, but peak concentrations were reached as the bloom was declining. Increasing pCO_2 had no discernable effect on the total amount of $DMSP_t$ produced at both temperatures tested. In contrast, increasing the pCO_2 to the value forecasted for the end of this century resulted in a linear decrease in DMS concentrations by 33 % and by as much as 69 % over the full pCO_2 gradient tested. These results are

consistent with previous reports that acidification has a greater impact on the processes that control the conversion of DMSP to DMS than on the production of DMSP itself. The pCO_2-induced decrease in DMS concentrations observed in this study adds to the bulk of previous studies reporting a similar trend. In diatom dominated systems, such as the one under study in this experiment, heterotrophic processes underlying DMS production seem to be most sensitive to modifications in pCO_2. Whereas predatory grazing and its associated impacts on DMS production cannot be ruled out entirely, the decreases in DMS concentrations in response to heightened pCO_2 are likely related to reductions in bacterial-mediated DMS production, a hypothesis partly supported by the significant positive correlations found between bacterial production rates and DMS concentrations. Whereas the DMS concentrations decreased significantly with increasing pCO_2 at both 10 °C and 15 °C, warming the mesocosms by 5 °C translated into a positive offset in concentrations of DMS over the whole range of pCO_2 tested. Higher DMSP release and increased bacterial productivity in the warm treatment partially explain the stimulating effect of temperature on DMS net production. Overall, results from this full factorial mesocosm experiment suggest that warming could mitigate the expected reduction in DMS production due to ocean acidification, even increasing the net DMS production with the potential to curtail radiative forcing.

Data availability. The data have been submitted to be freely accessible via Pangaea or can be obtained by contacting the author (robin.benard.1@ulaval.ca).

Author contributions. R. Bénard was responsible for the experimental design elaboration, data sampling and processing, and the writing of this article. Several co-authors supplied specific data included in this article, and all co-authors contributed to this final version of the article.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements

The authors wish to thank the Station Aquicole-ISMER, particularly Nathalie Morin and her staff, for their support during the mesocosm experiment. We also wish to acknowledge Gilles Desmeules, Bruno Cayouette, Sylvain Blondeau, Claire Lix, Rachel Hussherr, Liliane St-Amand, Marjolaine Blais, Armelle Galine Simo Matchim and Marie-Amélie Blais for their precious help over the duration of the experiment. This study was funded by a Team grant from the Fonds de recherche du Québec – Nature et technologies (FRQNT-Équipe-165335), the Canada Foundation for Innovation, the Canada Research Chair on Ocean Biogeochemistry and Climate, and by Fisheries and Oceans Canada. This is a contribution to the research program of Québec-Océan.

References

Andreae, M. O.: Ocean-atmosphere interactions in the global biogeochemical sulfur cycle, Mar. Chem., 30, 1–3, 1990.

Andreae, M.O. and Crutzen, P. J.: Atmospheric aerosols: biogeochemical sources and role in atmospheric chemistry, Science, 276(5315), 1052–1058, doi:10.1126/science.276.5315.1052, 1997.

Annane, S., St-Amand, L., Starr, M., Pelletier, E., and Ferreyra, G. A.: Contribution of transparent exopolymeric particles (TEP) to estuarine particulate organic carbon pool, Mar. Ecol. Prog. Ser., 529, 17–34, doi:10.3354/meps11294, 2015.

Avgoustidi, V., Nightingale, P. D., Joint, I., Steinke, M., Turner, S. M., Hopkins, F. E., and Liss, P. S.: Decreased marine dimethyl sulfide production under elevated CO_2 levels in mesocosm and in vitro studies, Environ. Chem., 9(4), 399, doi:10.1071/EN11125, 2012.

Barnard, W.R., Andreae, M.O., Watkins, W.E., Bingemer, H., and Georgii, H.W.: The flux of dimethylsulfide from the oceans to the atmosphere. J. Geophys. Res., 87(C11), 8787–8793, 1982.

Bates, T. S., Lamb, B. K., Guenther, A., Dignon, J., and Stoiber, R. E.: Sulfur emissions to the atmosphere from natural sources, J. Atmos. Chem., 14(1–4), 315–337, doi:10.1007/BF00115242, 1992.

Bell, R. T.: Estimating production of heterotrophic bacterioplankton via incorporation of tritiated thymidine, in: Handbook of methods in aquatic microbial ecology, Eds: Kemp, P. F., Sherr, B. F., Sherr, E. B., and Cole, J., Lewis Publisher, Boca Raton, 495–503, 1993.

Belzile, C., Brugel, S., Nozais, C., Gratton, Y. and Demers, S.: Variations of the abundance and nucleic acid content of heterotrophic bacteria in Beaufort Shelf waters during winter and spring, J. Mar. Syst., 74(3–4), 946–956, doi:10.1016/j.jmarsys.2007.12.010, 2008.

Bénard, R., Levasseur, M., Scarratt, M. G., Blais, M.-A., Mucci, A., Ferreyra, G., Starr, M., Gosselin, M., Tremblay, J.-É., and Lizotte, M.: Experimental assessment of the sensitivity of an estuarine phytoplankton fall bloom to acidification and warming, Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-31, accepted, 2018.

Boyd, P. W., Lennartz, S. T., Glover, D. M. and Doney, S. C.: Biological ramifications of climate-change-mediated oceanic multi-stressors, Nat. Clim. Chang., 5(1), 71–79, doi:10.1038/nclimate2441, 2015.

Boyd, P. W., Collins, S., Dupont, S., Fabricius, K., Gattuso, J.-P., Havenhand, J., Hutchins, D. A., Riebesell, U., Rintoul, M. S., Vichi, M., Biswas, H., Ciotti, A., Gao, K., Gehlen, M., Hurd, C. L., Kurihara, H., McGraw, C. M., Navarro, J. M., Nilsson, G. E., Passow, U. and Pörtner, H.-O.: Experimental strategies to assess the biological ramifications of multiple drivers of global ocean change-A review, Glob. Chang. Biol., 24(6), 2239–2261, doi:10.1111/gcb.14102, 2018.

Brimblecombe, P., and Shooter, D.: Photo-oxidation of dimethylsulphide in aqueous solution, Mar. Chem., 19, 343–353, 1986.

Byrne, R. H.: Standardization of Standard Buffers by Visible Spectrometry, Anal. Chem, 59, 1479–1481, doi:10.1021/ac00137a025, 1987.

Caldeira, K., and Wickett, M. E.: Ocean model predictions of chemistry changes from carbon dioxide emissions to the atmosphere and ocean, J. Geophys. Res., 110(C9), 1–12, doi:10.1029/2004JC002671, 2005.

Cai, W. J., and Wang, Y.: The chemistry, fluxes, and sources of carbon dioxide in the estuarine waters of the Satilla and Altamaha Rivers, Georgia, Limnol. Oceanogr., 43(4), 657–668, doi:10.4319/lo.1998.43.4.0657, 1998.

Cantin, G., Levasseur, M., Gosselin, M., Michaud, S.: Role of zooplankton in the mesoscale distribution of surface dimethylsulfide concentrations in the Gulf of St. Lawrence, Canada, Mar. Ecol. Prog. Ser., 141, 103-117, 1996.

Cantoni, G. L. and Anderson, D. Enzymatic cleavage of dimethylpropiothetin by Polysiphonia Lanosa, J. Bio. Chem., 222, 171-177, 1956.

Carslaw, K. S., Boucher, O., Spracklen, D. V., Mann, G. W., Rae, J. G. L.,Woodward, S., and Kulmala, M.: A review of natural aerosol interactions and feedbacks within the Earth system, Atmos. Chem. Phys., 10, 1701–1737, doi:10.5194/acp-10-1701- 2010, 2010.

Charlson, R., Lovelock, J., Andreae, M., and Warren, S.: Oceanic phytoplankton, atmospheric sulphur, cloud albedo and climate, Nature, 326, 656–661, 1987.

Clayton, T. D., and Byrne, R. H.: Spectrophotometric seawater pH measurements: total hydrogen ion concentration scale calibration of m-cresol purple and at-sea results, Deep. Res. Part I, 40(10), 2115–2129, doi:10.1016/0967-0637(93)90048-8, 1993.

Cyr, F., Bourgault, D., and Galbraith, P. S.: Interior versus boundary mixing of a cold intermediate layer, J. Geophys. Res. Ocean., 116(12), 1–12, doi:10.1029/2011JC007359, 2011.

Dacey, J. W. H., and Wakeham, S. G.: Oceanic dimethylsulfide: production during zooplankton grazing, Science, 233, 1314–1316, 1986.

Dickson, A. G.: Standard potential of the reaction: AgCl(s) + 1 $2H_2$(g) = Ag(s) + HCl(aq) and the standard acidity constant of the ion HSO_4^- in synthetic sea water from 273.15 to 318.15 K, J. Chem. Thermodyn., 22(2), 113–127, doi:10.1016/0021-9614(90)90074-Z, 1990.

Doney, S. C., Fabry, V. J., Feely, R. A., and Kleypas, J. A.: Ocean acidification: The other CO_2 problem, Ann. Rev. Mar. Sci., 1(1), 169–192, doi:10.1146/annurev.marine.010908.163834, 2009.

Engel, A., Delille, B., Jacquet, S., Riebesell, U., Rochelle-Newall, E., Terbrüggen, A., and Zondervan, I.: Transparent exopolymer particles and dissolved organic carbon production by *Emiliania huxleyi* exposed to different CO_2 concentrations: A mesocosm experiment, Aquat. Microb. Ecol., 34(1), 93–104, doi:10.3354/ame034093, 2004a.

Engel, A., Thoms, S., Riebesell, U., Rochelle-Newall, E., and Zondervan, I.: Polysaccharide aggregation as a potential sink of marine dissolved organic carbon, Nature, 428(6986), 929–932, doi:10.1038/nature02453, 2004b.

Engel, A., Piontek, J., Grossart, H.-P., Riebesell, U., Schulz, K. G., and Sperling, M.: Impact of CO_2 enrichment on organic matter dynamics during nutrient induced coastal phytoplankton blooms, J. Plankton Res., 36(3), 641–657, doi:10.1093/plankt/fbt125, 2014.

Feely, R. A., Doney, S. C., and Cooley, S. R.: Ocean Acidification: Present Conditions and Future Changes in a High-CO_2 World, Oceanography, 22(4), 36–47, doi:DOI 10.5670/oceanog.2009.95, 2009.

Fuhrman, J. A., and Azam, F.: Bacterioplankton secondary production estimates for coastal waters of British Columbia, Antarctica, and California, Appl. Environ. Microbiol., 39(6), 1085–1095, 1980.

Fuhrman, J. A., and Azam, F.: Thymidine incorporation as a measure of heterotrophic bacterioplankton production in marine surface waters: Evaluation and field results, Mar. Biol., 66(2), 109–120, doi:10.1007/BF00397184, 1982.

Gaaloul, H.: Effets du changement global sur les particules exopolymériques transparentes au sein de l'estuaire maritime du Saint-Laurent, M.Sc. thesis, Université du Québec à Rimouski, Canada, 133 pp., 2017.

Galbraith, P. S., Chassé, J., Gilbert, D., Larouche, P., Brickman, D., Pettigrew, B., Devine, L, Gosselin, A., Pettipas, R. G., and Lafleur, C.: Physical Oceanographic Conditions in the Gulf of St. Lawrence in 2011, DFO Can. Sci. Advis. Sec. Res. Doc., 2012/023, iii + 85 pp, 2012.

Galbraith, P. S., Chassé, J., Caverhill, C., Nicot, P., Gilbert, D., Pettigrew, B., Lefaivre, D., Brickman, D., Devine, L., and Lafleur, C.: Physical Oceanographic Conditions in the Gulf of St. Lawrence in 2015, DFO Can. Sci. Advis. Sec. Res. Doc., 2016/056, v + 90 pp, 2016.

Gattuso, J.-P., Magnan, A., Bille, R., Cheung, W. W. L., Howes, E. L., Joos, F., Allemand, D., Bopp, L., Cooley, S. R., Eakin, C. M., Hoegh-Guldberg, O., Kelly, R. P., Portner, H.-O., Rogers, a. D., Baxter, J. M., Laffoley, D., Osborn, D., Rankovic, A., Rochette, J., Sumaila, U. R., Treyer, S., and Turley, C.: Contrasting futures for ocean and society from different anthropogenic CO_2 emissions scenarios, Science, 349(6243), doi:10.1126/science.aac4722, 2015.

Green, D.H., Shenoy, D. M., Hart, M. C., and Hatton, A. D.: Coupling of dimethylsulfide oxidation to biomass production by a marine Flavobacterium, Appl. Environ. Microbiol., 77(9), 3137–3140, doi:10.1128/AEM.02675-10, 2011.

Grossart, H.-P., Allgaier, M., Passow, U., and Riebesell, U.: Testing the effect of CO_2 concentration on the dynamics of marine heterotrophic bacterioplankton, Limnol. Oceanogr., 51(1), 1–11, doi:10.4319/lo.2006.51.1.0001, 2006.

Gunderson, A. R., Armstrong, E. J., and Stillman, J. H.: Multiple Stressors in a Changing World: The Need for an Improved Perspective on Physiological Responses to the Dynamic Marine Environment, Ann. Rev. Mar. Sci., 8(1), 357–378, doi:10.1146/annurev-marine-122414-033953, 2016.

Hansen, H. P., and Koroleff, F.: Determination of nutrients, in: Methods of Seawater Analysis, 3, Eds: Grasshoff K., Kremling, K.., and Ehrhardt, M., Wiley-VCH Verlag GmbH, Weinheim, Germany, 159–228, doi:10.1002/9783527613984.ch10, 2007.

Hatton, A. D., Darroch, L., and Malin, G.: The role of dimethylsulphoxide in the marine biogeochemical cycle of dimethylsulphide, Oceanogr. Mar. Biol. An Annu. Rev., 42, 29–56, 2004.

Hopkins, F. E., and Archer, S. D.: Consistent increase in dimethyl sulfide (DMS) in response to high CO_2 in five shipboard bioassays from contrasting NW European waters, Biogeosciences, 11(18), 4925–4940, doi:10.5194/bg-11-4925-2014, 2014.

Hopkins, F. E., Turner, S. M., Nightingale, P. D., Steinke, M., Bakker, D., and Liss, P. S.: Ocean acidification and marine trace gas emissions, Proc. Natl. Acad. Sci., 107(2), 760–765, doi:10.1073/pnas.0907163107, 2010.

Hussherr, R., Levasseur, M., Lizotte, M., Tremblay, J. É., Mol, J., Thomas, H., Gosselin, M., Starr, M., Miller, L. A., Jarniková, T., Schuback, N., and Mucci, A.: Impact of ocean acidification on Arctic phytoplankton blooms and dimethyl sulfide concentration under simulated ice-free and under-ice conditions, Biogeosciences, 14(9), 2407–2427, doi:10.5194/bg-14-2407-2017, 2017.

606 IPCC: Working Group I Contribution to the Fifth Assessment Report Climate Change 2013: The Physical Science Basis,
607 Intergov. Panel Clim. Chang., 1535, doi:10.1017/CBO9781107415324., 2013.
608 Iverson, R. L., Nearhoof, F. L. and Andreae, M. O.: Production of dimethylsulfonium propionate and dimethylsulfide by
609 phytoplankton in estuarine and coastal waters, Limnol. Oceanogr., 34(1), 53–67, doi:10.4319/lo.1989.34.1.0053, 1989.
610 Karsten, U., Kück, K., Vogt, C., and Kirst, G. O.: Dimethylsulfoniopropionate production in phototrophic organisms and its
611 physiological functions as a cryoprotectant, in: Biological and environmental chemistry of DMSP and related sulfonium
612 compounds, Eds: Kiene, R. P., Visscher, P. T., Keller, M. D., and Kirst, G. O., Springer US, Boston, MA, 143–153,
613 doi:10.1007/978-1-4613-0377-0, 1996.
614 Keller, M. D.: Dimethyl sulfide production and marine phytoplankton: the importance of species composition and cell size,
615 Biol. Oceanogr., 6(5–6), 375–382, doi:10.1080/01965581.1988.10749540, 1989.
616 Kettle, A. J., Andreae, M. O., Amouroux, D., Andreae, T. W., Bates, T. S., Berresheim, H., Bingemer, H., Boniforti, R., Curran,
617 M. A. J., diTullio, G. R., Helas, G., Jones, G. B., Keller, I. M. D., Kiene, R. P., Leck, C., Levasseur, M., Maspero, M., Matrai,
618 P., McTaggart, A. R., Mihalopoulos, N., Nguyen, B. C., Novo, A., Putaud, J. P., Rapsomanikis, S., Roberts, G., Schebeske,
619 G.; Sharma, S., Simó, R., Staubes, R., Turner, S., and Uher, G.: A global database of sea surface dimethylsulfide (DMS)
620 measurements and a procedure to predict sea surface DMS as a function of latitude, longitude, and month, Global Biogeochem.
621 Cycles, 13, 399-444, 1999.
622 Kettles, N. L., Kopriva, S., and Malin, G.: Insights into the regulation of DMSP synthesis in the diatom *Thalassiosira*
623 *pseudonana* through APR activity, proteomics and gene expression analyses on cells acclimating to changes in salinity, light
624 and nitrogen, PLoS One, 9(4), doi:10.1371/journal.pone.0094795, 2014.
625 Kiene, R. P., Linn, L. J.: Distribution and turnover of dissolved DMSP and its relationship with bacterial production and
626 dimethylsulfide in the Gulf of Mexico, Limnol. Oceanogr., 45, 849–861, 2000.
627 Kiene, R. P., and Service, S. K.: Decomposition of dissolved DMSP and DMS in estuarine waters: dependence on temperature
628 and substrate concentration, Mar. Ecol. Prog. Ser., 76, 1–11, 1991.
629 Kiene, R. P., Linn, L. J., Gonzalez, J., Moran, M. A., and Bruton, J. A.: Dimethylsulfoniopropionate and methanethiol are
630 important precursors of methionine and protein-sulfur in marine bacterioplankton, Appl Environ. Microbiol., 65, 4549–4558,
631 1999.
632 Kiene, R. P., Linn, L. J., and Bruton, J. A. New and important roles for DMSP in marine microbial communities, J. Sea. Res.,
633 43, 209–224, 2000.
634 Kim, K. Y., Garbary, D. J., and Mclachlan, J. L.: Phytoplankton dynamics in Pomquet Harbour, Nova Scotia: a lagoon in the
635 southern Gulf of St Lawrence, Phycologica, 43(3), 311–328, 2004.
636 Kim, J. M., Lee, K., Yang, E. J., Shin, K., Noh, J. H., Park, K. T., Hyun, B., Jeong, H. J., Kim, J. H., Kim, K. Y., Kim, M.,
637 Kim, H. C., Jang, P. G. and Jang, M. C.: Enhanced production of oceanic dimethylsulfide resulting from CO_2-induced grazing
638 activity in a high CO_2 world, Environ. Sci. Technol., 44(21), 8140–8143, doi:10.1021/es102028k, 2010.

Kirst, G. O., Thiel, C., Wolff, H., Nothnagel, J., Wanzek M., and Ulmke, R.: Dimethylsulfoniopropionate (DMSP) in ice-algae and its possible biological role, Mar. Chem., 35, 381–388, 1991.

Kwint, R. L., and Kramer, K. J.: Dimethylsulphide production by plankton communities, Mar. Ecol. Prog. Ser., 121(1–3), 227–238, doi:10.3354/meps121227, 1995.

Hill, R. W., White, B. A., Cottrell, M. T., Dacey, J. W. H.: Virus-mediated total release of dimethylsulfoniopropionate from marine phytoplankton: a potential climate process, Aquat. Microb. Ecol., 14, 1–6, 1998.

Lana, A., Bell, T. G., Simó, R., Vallina, S. M., Ballabrera-Poy, J., Kettle, a. J., Dachs, J., Bopp, L., Saltzman, E. S., Stefels, J., Johnson, J. E., and Liss, P. S.: An updated climatology of surface dimethylsulfide concentrations and emission fluxes in the global ocean, Global Biogeochem. Cycles, 25(1), 1–17, doi:10.1029/2010GB003850, 2011.

Laroche, D., Vézina, A. F., Levasseur, M., Gosselin, M., Stefels, J., Keller, M. D., Matrai, P. A., and Kwint, R. L. J.: DMSP synthesis and exudation in phytoplankton: A modeling approach, Mar. Ecol. Prog. Ser., 180(May), 37–49, doi:10.3354/meps180037, 1999.

Lee, P. A., Saunders, P. A., De Mora, S. J., Deibel, D. and Levasseur, M.: Influence of copepod grazing on concentrations of dissolved dimethylsulfoxide and related sulfur compounds in the North Water, Northern Baffin Bay, Mar. Ecol. Prog. Ser., 255, 235–248, doi:10.3354/meps255235, 2003.

Lee, P. A., Rudisill, J. R., Neeley, A. R., Maucher, J. M., Hutchins, D. A., Feng, Y., Hare, C. E., Leblanc, K., Rose, J. M., Wilhelm, S. W., Rowe, J. M., and Giacomo, R.: Effects of increased pCO_2 and temperature on the North Atlantic spring bloom. III. Dimethylsulfoniopropionate, Mar. Ecol. Prog. Ser., 388, 41–49, doi:10.3354/meps08135, 2009.

Le Quéré, C., Andres, R. J., Boden, T., Conway, T., Houghton, R. A., House, J. I., Marland, G., Peters, G. P., Van Der Werf, G. R., Ahlström, A., Andrew, R. M., Bopp, L., Canadell, J. G., Ciais, P., Doney, S. C., Enright, C., Friedlingstein, P., Huntingford, C., Jain, A. K., Jourdain, C., Kato, E., Keeling, R. F., Klein Goldewijk, K., Levis, S., Levy, P., Lomas, M., Poulter, B., Raupach, M. R., Schwinger, J., Sitch, S., Stocker, B. D., Viovy, N., Zaehle, S., and Zeng, N.: The global carbon budget 1959-2011, Earth Syst. Sci. Data, 5(1), 165–185, doi:10.5194/essd-5-165-2013, 2013.

Le Quéré, C., Moriarty, R.., andrew, R. M., Canadell, J. G., Sitch, S., Korsbakken, J. I., Friedlingstein, P., Peters, G. P.., andres, R. J., Boden, T. A., Houghton, R. A., House, J. I., Keeling, R. F., Tans, P., Arneth, A., Bakker, D. C. E., Barbero, L., Bopp, L., Chang, J., Chevallier, F., Chini, L. P., Ciais, P., Fader, M., Feely, R. A., Gkritzalis, T., Harris, I., Hauck, J., Ilyina, T., Jain, A. K., Kato, E., Kitidis, V., Klein Goldewijk, K., Koven, C., Landschützer, P., Lauvset, S. K., Lefèvre, N., Lenton, A., Lima, I. D., Metzl, N., Millero, F., Munro, D. R., Murata, A., S. Nabel, J. E. M., Nakaoka, S., Nojiri, Y., O'Brien, K., Olsen, A., Ono, T., Pérez, F. F., Pfeil, B., Pierrot, D., Poulter, B., Rehder, G., Rödenbeck, C., Saito, S., Schuster, U., Schwinger, J., Séférian, R., Steinhoff, T., Stocker, B. D., Sutton, A. J., Takahashi, T., Tilbrook, B., Van Der Laan-Luijkx, I. T., Van Der Werf, G. R., Van Heuven, S., Vandemark, D., Viovy, N., Wiltshire, A., Zaehle, S., and Zeng, N.: Global Carbon Budget 2015, Earth Syst. Sci. Data, 7(2), 349–396, doi:10.5194/essd-7-349-2015, 2015.

671 Levasseur, M., Michaud, S., Egge, J., Cantin, G., Nejstgaard, J. C., Sanders, R., Fernandez, E., Solberg, P. T., Heimdal, B.,
672 and Gosselin, M.: Production of DMSP and DMS during a mesocosm study of an *Emiliania huxleyi* bloom: Influence of
673 bacteria and *Calanus finmarchicus* grazing, Mar. Biol., 126(4), 609–618, doi:10.1007/BF00351328, 1996.
674 Liss, P. S. and Lovelock, J. E.: Climate change: The effect of DMS emissions, Environ. Chem., 4(6), 377–378,
675 doi:10.1071/EN07072, 2007.
676 Lizotte, M., Levasseur, M., Michaud, S., Scarratt, M. G., Merzouk, A., Gosselin, M., Pommier, J., Rivkin, R. B., and Kiene,
677 R. P.: Macroscale patterns of the biological cycling of dimethylsulfoniopropionate (DMSP) and dimethylsulfide (DMS) in the
678 Northwest Atlantic, Biogeochemistry, 110(1–3), 183–200, doi:10.1007/s10533-011-9698-4, 2012.
679 Lizotte, M., Levasseur, M., Law, C. S., Walker, C. F., Safi, K. A., Marriner, A. and Kiene, R. P.: Dimethylsulfoniopropionate
680 (DMSP) and dimethyl sulfide (DMS) cycling across contrasting biological hotspots of the New Zealand subtropical front,
681 Ocean Sci., 13(6), 961–982, doi:10.5194/os-13-961-2017, 2017.
682 Lovelock, J. E., Maggs, R. J., and Rasmusse R. A.: Atmospheric dimethyl sulfide and natural sulfur cycle, Nature, 237(5356),
683 452-453, 1972.
684 Malin, G., and Kirst, G. O.: Algal production of dimethyl sulfide and its atmospheric role, J. Phycol., 33, 889-896, 1997.
685 Malin, G., Wilson, W. H., Bratbak, G., Liss, P. S., and Mann, N. H.: Elevated production of dimethylsulfide resulting from
686 viral infection of cultures of *Phaeocystis pouchetii*, Limnol. Oceanogr., 43(6), 1389–1393, doi:10.4319/lo.1998.43.6.1389,
687 1998.
688 Malmstrom, R. R., Kiene, R. P., Cottrell, M. T., and Kirchman, D. L.: Contribution of SAR11 bacteria to dissolved
689 dimethylsulfoniopropionate and amino acid uptake in the north Atlantic Ocean, Appl. Environ. Microbiol., 70(7), 4129–4135,
690 doi:10.1128/AEM.70.7.4129, 2004a.
691 Malmstrom, R. R., Kiene, R. P., and Kirchman, D. L.: Identification and enumeration of bacteria assimilating
692 dimethylsulfoniopropionate (DMSP) in the North Atlantic and Gulf of Mexico, Limnol. Oceanogr., 49, 597–606, doi:
693 10.4319/lo.2004.49.2.0597, 2004b.
694 Malmstrom, R. R., Kiene, R. P., Vila, M., and Kirchman, D. L.: Dimethylsulfoniopropionate (DMSP) assimilation by
695 *Synechococcus* in the Gulf of Mexico and northwest Atlantic Ocean, Limnol. Oceanogr., 50(6), 1924–1931,
696 doi:10.4319/lo.2005.50.6.1924, 2005.
697 Marie, D., Simon, N., and Vaulot, D.: Phytoplankton cell counting by flow cytometry, Algal Cult. Tech., 253–267,
698 doi:10.1016/B978-012088426-1/50018-4, 2005.
699 Mucci, A., Levasseur, M., Gratton, Y., Martias, C., Scarratt, M., Gilbert, D., Tremblay, J.-É., Ferreyra, G., and Lansard, B.:
700 Tidally-induced variations of pH at the head of the Laurentian Channel, Can. J. Fish. Aquat. Sci., doi:10.1139/cjfas-2017-
701 0007, 2017.
702 Nguyen, B. C., Belviso, S., Mihalopoulos, N., Gostan, J., Nival., P.: Dimethyl sulfide production during natural phytoplankton
703 blooms, Mar. Chem., 24, 133–141, 1988

Nightingale, P.D., Malin, G., Law, C.S., Watson, A.J., Liss, P.S., Liddicoat, M.I., Boutin, J., Upstill-Goddard, R.C.: In situ evaluation of air–sea gas exchange parameterizations using novel conservative and volatile tracers, Global Biogeochem. Cycles, 14 (1), 373–387, 2000.

Niki, T., Kunugi, M., and Otsuki, A.: DMSP-lyase activity in five marine phytoplankton species: Its potential importance in DMS production, Mar. Biol., 136(5), 759–764, doi:10.1007/s002279900235, 2000.

Park, K. T., Lee, K., Shin, K., Yang, E. J., Hyun, B., Kim, J. M., Noh, J. H., Kim, M., Kong, B., Choi, D. H., Choi, S. J., Jang, P. G., and Jeong, H. J.: Direct linkage between dimethyl sulfide production and microzooplankton grazing, resulting from prey composition change under high partial pressure of carbon dioxide conditions, Environ. Sci. Technol., 48(9), 4750–4756, doi:10.1021/es403351h, 2014.

Parsons, T. R., Maita, Y., and Lalli, C. M.: A manual of chemical and biological methods for seawater analysis, Permagon Press, New York, 1984.

Paul, C., Sommer, U., Garzke, J., Moustaka-Gouni, M., Paul, A., and Matthiessen, B.: Effects of increased CO_2 concentration on nutrient limited coastal summer plankton depend on temperature, Limnol. Oceanogr., 61(3), 853–868, doi:10.1002/lno.10256, 2016.

Pierrot, D. E., Lewis, E., and Wallace, D. W. R.: MS Excel program developed for CO2 system calculations, Carbon Dioxide Information Analysis Center, ONRL/CDIAC-105a, Oak Ridge National Laboratory, US Department of Energy, Oak Ridge, Tennessee, USA, 2006.

Pinhassi, J., Simó, R., González, J. M., Vila, M., Alonso-Sáez, L., Kiene, R. P., Moran, M. A., and Pedrós-Alió, C.: Dimethylsulfoniopropionate turnover is linked to the composition and dynamics of the bacterioplankton assemblage during a microcosm phytoplankton bloom, Appl. Environ. Microbiol., 71(12), 7650–7660, doi:10.1128/AEM.71.12.7650-7660.2005, 2005.

Quinn, P. K., and Bates, T. S.: The case against climate regulation via oceanic phytoplankton sulphur emissions., Nature, 480(7375), 51–6, doi:10.1038/nature10580, 2011.

Quinn, P. K., Coffman, D. J., Johnson, J. E., Upchurch, L. M., and Bates, T. S.: Small fraction of marine cloud condensation nuclei made up of sea spray aerosol, Nat. Geosci., 10(9), 674–679, doi:10.1038/ngeo3003, 2017.

R Core Team: R: A language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria. URL https://www.R-project.org/, 2016.

Riebesell, U. and Gattuso, J. P.: Lessons learned from ocean acidification research, Nat. Clim. Chang., 5(1), 12–14, doi:10.1038/nclimate2456, 2015.

Robert-Baldo, G., Morris, M., and Byrne, R.: Spectrophotometric determination of seawater pH using phenol red, Anal. Chem., 3(57), 2564–2567, doi:10.1021/ac00290a030, 1985.

Roemmich, D., Church, J., Gilson, J., Monselesan, D., Sutton, P. and Wijffels, S.: Unabated planetary warming and its ocean structure since 2006, Nat. Clim. Chang., 5(3), 240–245, doi:10.1038/nclimate2513, 2015.

737 Ruiz-González, C., Galí, M., Gasol, J. M. and Simó, R.: Sunlight effects on the DMSP-sulfur and leucine assimilation activities
738 of polar heterotrophic bacterioplankton, Biogeochemistry, 110(1–3), 57–74, doi:10.1007/s10533-012-9699-y, 2012.
739 Scarratt, M. G., Levasseur, M., Schultes, S., Michaud, S., Cantin, G., Vézina, A., Gosselin, M. and De Mora, S. J.: Production
740 and consumption of dimethylsulfide (DMS) in North Atlantic waters, Mar. Ecol. Prog. Ser., 204, 13–26,
741 doi:10.3354/meps204013, 2000.
742 Schlüter, L., Lohbeck, K. T., Gröger, J. P., Riebesell, U. and Reusch, T. B. H.: Long-term dynamics of adaptive evolution in
743 a globally important phytoplankton species to ocean acidification, Sci. Adv., 2(7), e1501660–e1501660,
744 doi:10.1126/sciadv.1501660, 2016.
745 Schwinger, J., Tjiputra, J., Goris, N., Six, K. D., Kirkevåg, A., Seland, Ø., Heinze, C. and Ilyina, T.: Amplification of global
746 warming through pH dependence of DMS production simulated with a fully coupled Earth system model, Biogeosciences,
747 14(15), 3633–3648, doi:10.5194/bg-14-3633-2017, 2017.
748 Simó, R.: Production of atmospheric sulfur by oceanic plankton: Biogeochemical, ecological and evolutionary links, Trends
749 Ecol. Evol., 16(6), 287–294, doi:10.1016/S0169-5347(01)02152-8, 2001.
750 Simó, R.: From cells to globe: approaching the dynamics of DMS(P) in the ocean at multiple scales, Can. J. Fish. Aquat. Sci.,
751 61(5), 673–684, doi:10.1139/f04-030, 2004.
752 Simó, R., and Pedrós-Alió, C.: Role of vertical mixing in controlling the oceanic production of dimethyl sulphide, Nature, 402,
753 396–399, doi:10.1038/46516, 1999.
754 Six, K. D., Kloster, S., Ilyina, T., Archer, S. D., Zhang, K., and Maier-Reimer, E.: Global warming amplified by reduced
755 sulphur fluxes as a result of ocean acidification, Nat. Clim. Chang., 3(11), 975–978, doi:10.1038/nclimate1981, 2013.
756 Starr, M., St-Amand, L., Devine, L., Bérard-Therriault, L. and Galbraith, P. S.: State of phytoplankton in the Estuary and Gulf
757 of St. Lawrence during 2003, CSAS Res. Doc., 2004/123, 35, 2004.
758 Stefels, J.: Physiological aspects of the production and conversion of DMSP in marine algae and higher plants, J. Sea Res.,
759 43(3–4), 183–197, doi:10.1016/S1385-1101(00)00030-7, 2000.
760 Stefels, J., and Van Boekel, W. H. M.: Production of DMS from dissolved DMSP in axenic cultures of the marine
761 phytoplankton species *Phaeocystis* sp., Mar. Ecol. Prog. Ser., 97(1), 11–18, doi:10.3354/meps097011, 1993.
762 Stefels, J., Steinke, M., Turner, S., Malin, G., and Belviso, S.: Environmental constraints on the production and removal of the
763 climatically active gas dimethylsulphide (DMS) and implications for ecosystem modelling, Biogeochemistry, 83(1–3), 245–
764 275, doi:10.1007/s10533-007-9091-5, 2007.
765 Steinke, M., Malin, G., Archer, S. D., Burkill, P. H. and Liss, P. S.: DMS production in a coccolithophorid bloom: Evidence
766 for the importance of dinoflagellate DMSP lyases, Aquat. Microb. Ecol., 26(3), 259–270, doi:10.3354/ame026259, 2002.
767 Stillman, J. H. and Paganini, A. W.: Biochemical adaptation to ocean acidification, J. Exp. Biol., 218(12), 1946–1955,
768 doi:10.1242/jeb.115584, 2015.
769 Sunda, W., Kieber, D. J., Kiene, R. P., and Huntsman, S.: An antioxidant function for DMSP and DMS in marine algae, Nature,
770 418(6895), 317-320, 2002.

Toole, D. A., and Siegel, D. A.: Light-driven cycling of dimethylsulfide (DMS) in the Sargasso Sea: closing the loop, Geophys. Res. Lett., 31(9), 1–4, doi:10.1029/2004GL019581, 2004.

Toole, D. A., Slezak, D., Kiene, R. P., Kieber, D. J., and Siegel, D. A.: Effects of solar radiation on dimethylsulfide cycling in the western Atlantic Ocean, Deep. Res. Part I Oceanogr. Res. Pap., 53(1), 136–153, doi:10.1016/j.dsr.2005.09.003, 2006.

Toole, D. A., Siegel, D. A. and Doney, S. C.: A light-driven, one-dimensional dimethylsulfide biogeochemical cycling model for the Sargasso Sea, J. Geophys. Res. Biogeosciences, 113(2), 1–20, doi:10.1029/2007JG000426, 2008.

Vallina, S. M., Simó, R., Anderson, T. R., Gabric, A., Cropp, R., and Pacheco, J. M.: A dynamic model of oceanic sulfur (DMOS) applied to the Sargasso Sea: Simulating the dimethylsulfide (DMS) summer paradox, J. Geophys. Res. Biogeosciences, 113(1), doi:10.1029/2007JG000415, 2008.

Vila, M., Simó, R., Kiene, R. P., Pinhassi, J., González, J. M., Moran, M. A., and Pedrós-Alió, C.: Use of microautoradiography combined with fluorescence in situ hybridization to determine dimethylsulfoniopropionate incorporation by marine bacterioplankton taxa, Appl. Environ. Microbiol., 70, 4648– 4657, https://doi.org/10.1128/AEM.70.8.4648-4657.2004, 2004.

Vila-Costa, M., Simó, R., Harada, H., Gasol, J. M., Slezak, D., and Kiene, R. P.: Dimethylsulfoniopropionate Uptake by Marine Phytoplankton, Science, 314, 652–654, 2006a.

Vila-Costa, M., Del Valle, D. A., González, J. M., Slezak, D., Kiene, R. P., Sánchez, O., and Simó, R.: Phylogenetic identification and metabolism of marine dimethylsulfide consuming bacteria, Environ. Microbiol., 8, 2189–2200, https://doi.org/10.1111/j.1462-2920.2006.01102.x, 2006b.

Vila-Costa, M., Pinhassi, J., Alonso, C., Pernthaler, J., and Simó, R.: An annual cycle of dimethylsulfoniopropionate sulfur and leucine assimilating bacterioplankton in the coastal NW Mediterranean, Environ. Microbiol., 9, 2451–2463, https://doi.org/10.1111/j.1462-2920.2007.01363.x, 2007.

Vogt, M., Steinke, M., Turner, S., Paulino, a., Meyerhöfer, M., Riebesell, U., Le Quéré, C. and Liss, P.: Dynamics of dimethylsulphoniopropionate and dimethylsulphide under different CO_2 concentrations during a mesocosm experiment, Biogeosciences, 5(2), 407–419, doi:10.5194/bg-5-407-2008, 2008.

Webb, A., Malin, G., Hopkins, F., Ho, K. L., Riebesell, U., Schulz, K., Larsen, A., and Liss, P.: Ocean acidification has different effects on the production of DMS and DMSP measured in cultures of *Emiliania huxleyi* and a mesocosm study: a comparison of laboratory monocultures and community interactions, Environ. Chem., 13(2), 314–329, doi:/10.1071/EN14268, 2015.

Webb, A. L., Leedham-Elvidge, E., Hughes, C., Hopkins, F. E., Malin, G., Bach, L. T., Schulz, K., Crawfurd, K., Brussaard, C. P. D., Stuhr, A., Riebesell, U., and Liss, P. S.: Effect of ocean acidification and elevated fCO_2 on trace gas production by a Baltic Sea summer phytoplankton community, Biogeosciences, 13(15), 4595–4613, doi:10.5194/bg-13-4595-2016, 2016.

Wolfe, G. V., and Steinke, M.: Grazing-activated production of dimethyl sulfide (DMS) by two clones of *Emiliania huxleyi*, Limnol. Oceanogr., 41(6), 1151–1160, doi:10.4319/lo.1996.41.6.1151, 1996.

Woodhouse, M. T., Mann, G. W., Carslaw, K. S., and Boucher, O.: Sensitivity of cloud condensation nuclei to regional changes in dimethyl-sulphide emissions, Atmos. Chem. Phys., 13(5), 2723–2733, doi:10.5194/acp-13-2723-2013, 2013.

805 Yoch, D. C.: Dimethylsulfoniopropionate: Its sources, role in the marine food web, and biological degradation to
806 dimethylsulfide, Appl. Environ. Microbiol., 68(12), 5804–5815, doi:10.1128/AEM.68.12.5804-5815.2002, 2002.
807

Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

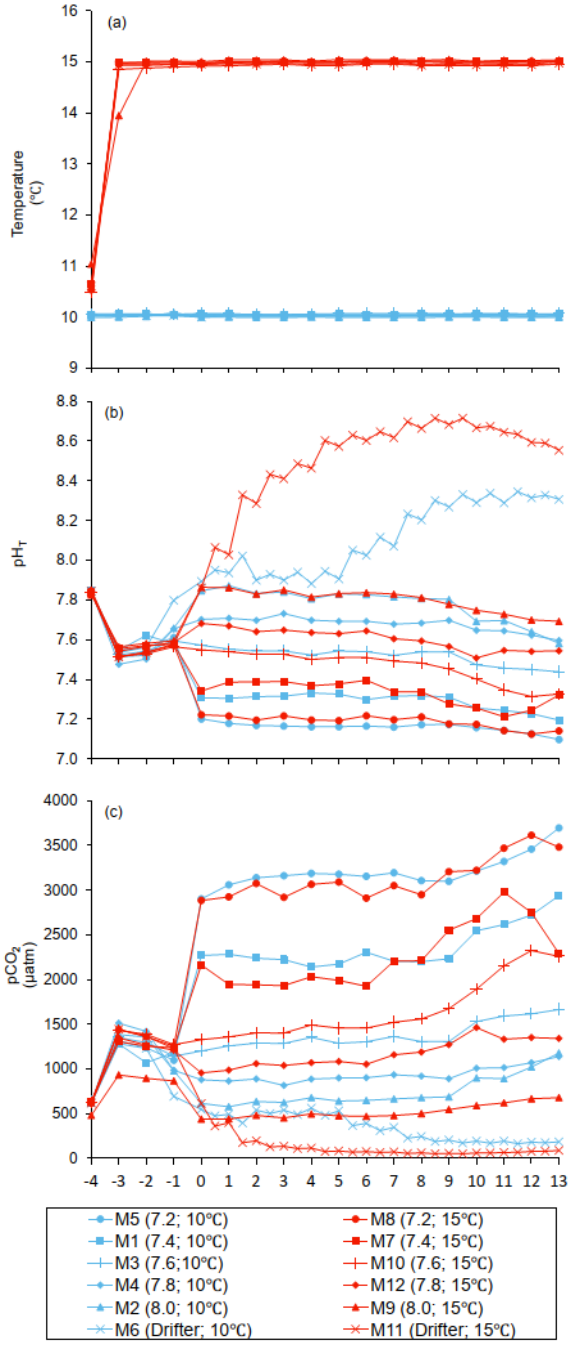

Figure 1. Temporal variations over the course of the experiment for: (a) temperature, (b) pH_T, (c) pCO_2. For symbol attribution to treatments, see legend. Adapted from Bénard et al. (2018).

Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

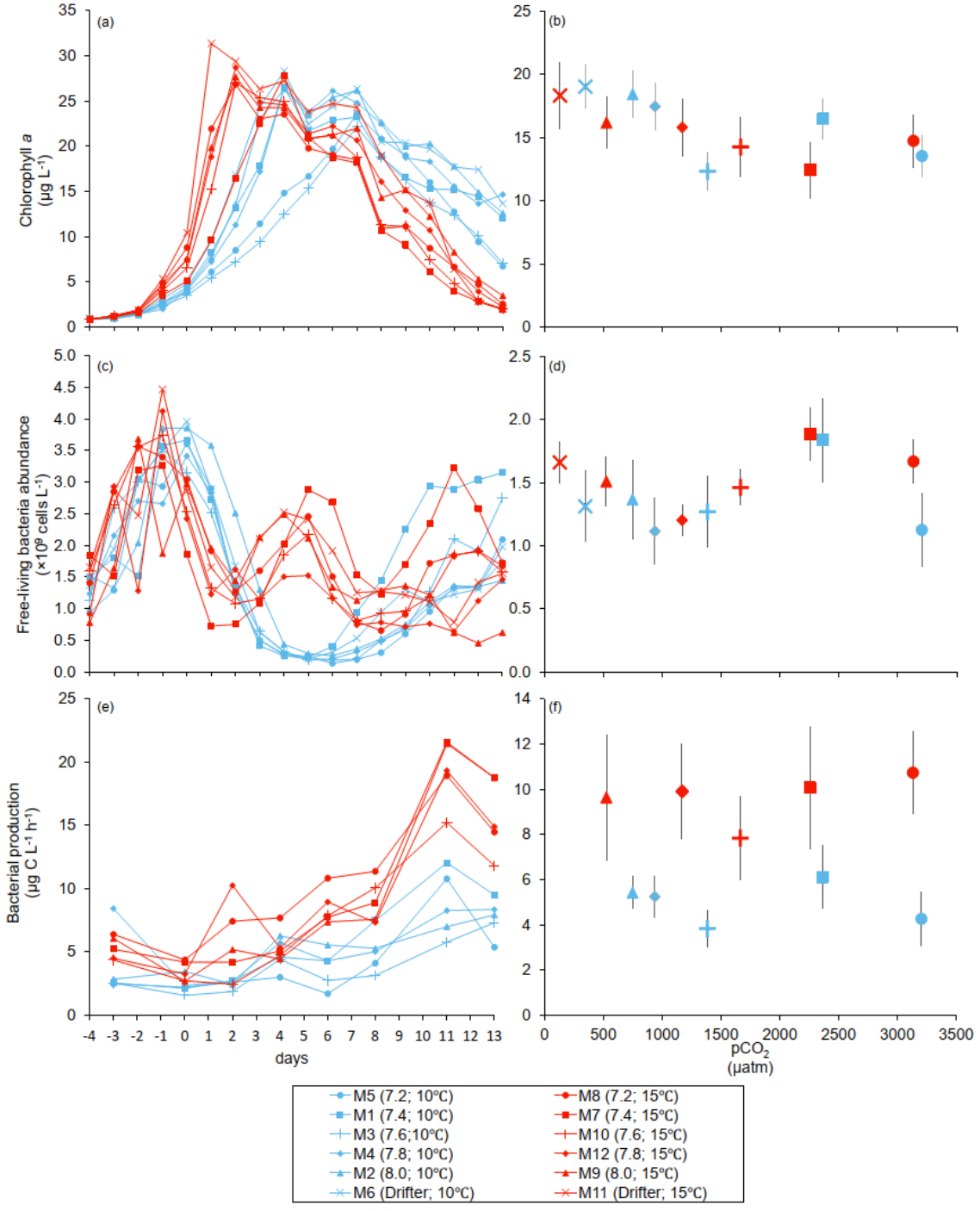

Figure 2. Temporal variations, and averages over the course of the experiment (day 0 to day 13) for: (a–b) chlorophyll *a* (adapted from Bénard et al., 2018), (c–d) free-living bacteria abundance, (e–f) bacterial production. For symbol attribution to treatments, see legend.

Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

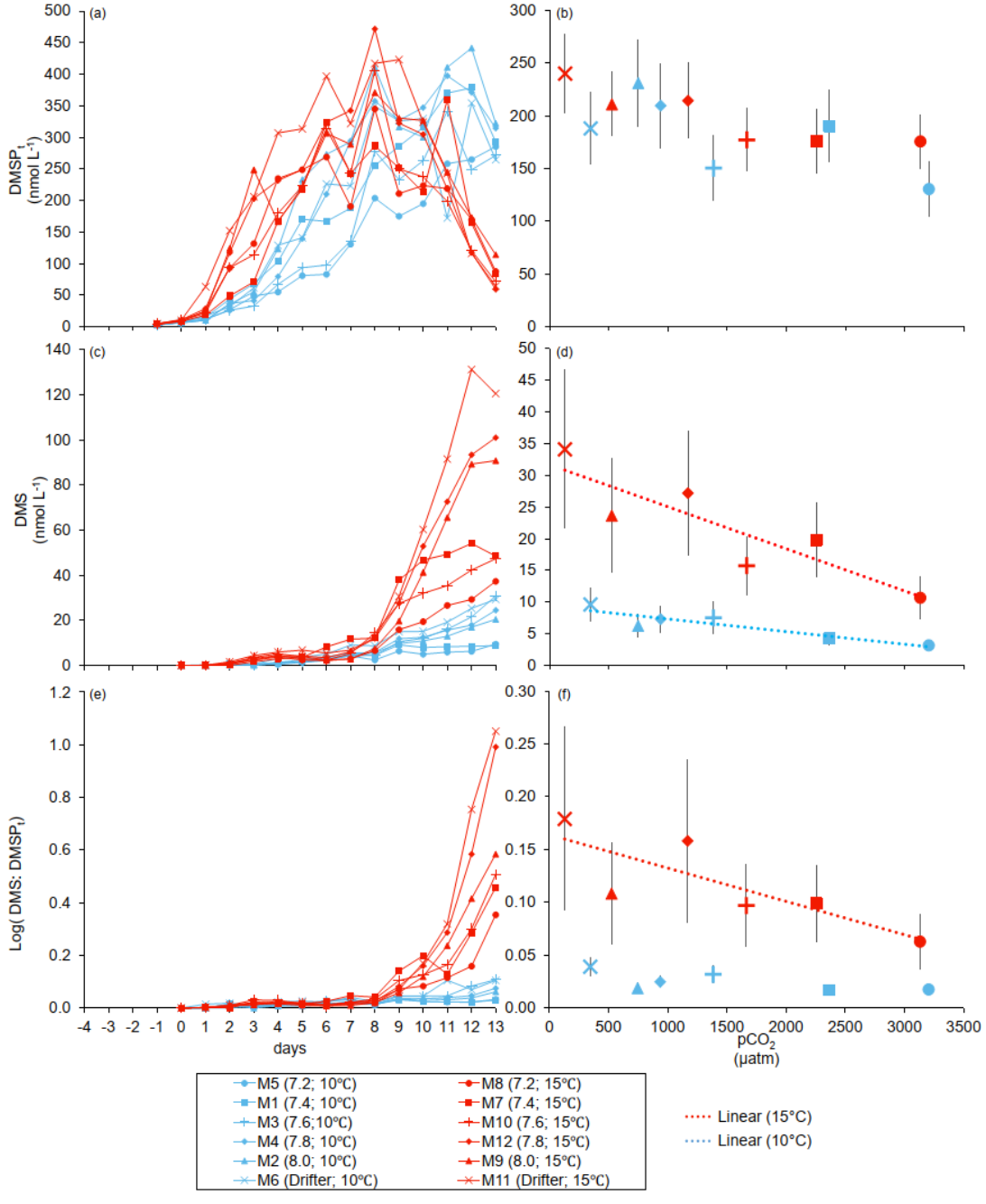

Figure 3. Temporal variations, and averages over the course of the experiment (day 0 to day 13) for: (a–b) DMSP$_t$, (c–d) DMS, (e–f) the natural logarithm of the DMS:DMSP$_t$ ratio. For symbol attribution to treatments, see legend.

Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

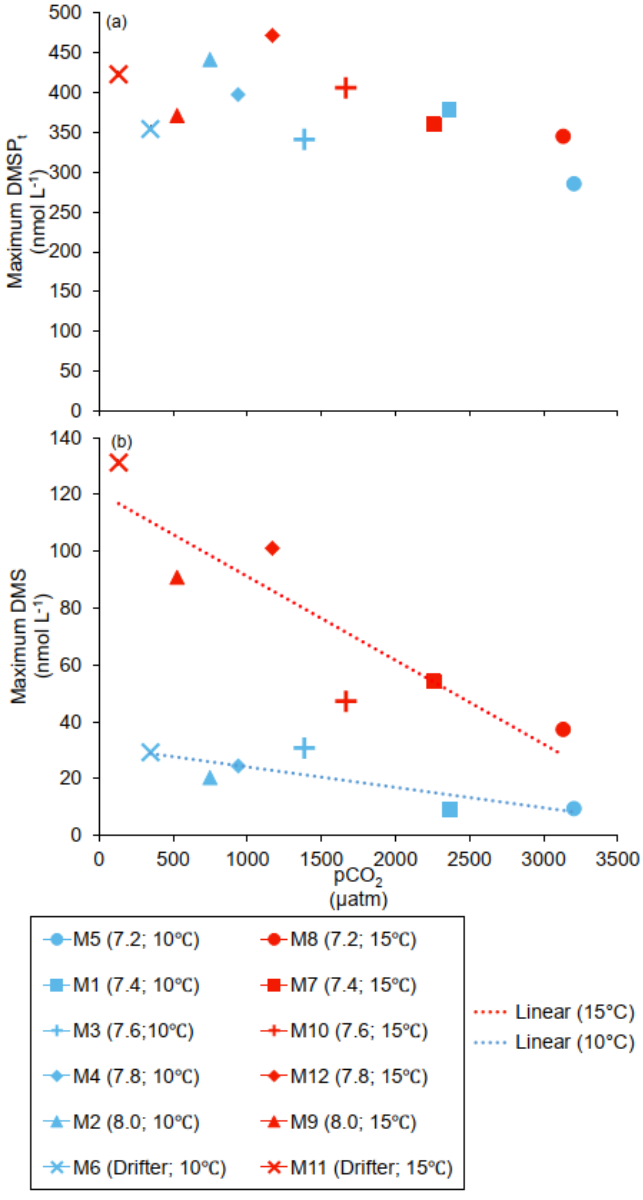

Figure 4. (a) Maximum $DMSP_t$ concentrations, (b) maximum DMS concentrations reached over the full course of the experiment (day 0 to day 13). For symbol attribution to treatments, see legend

Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

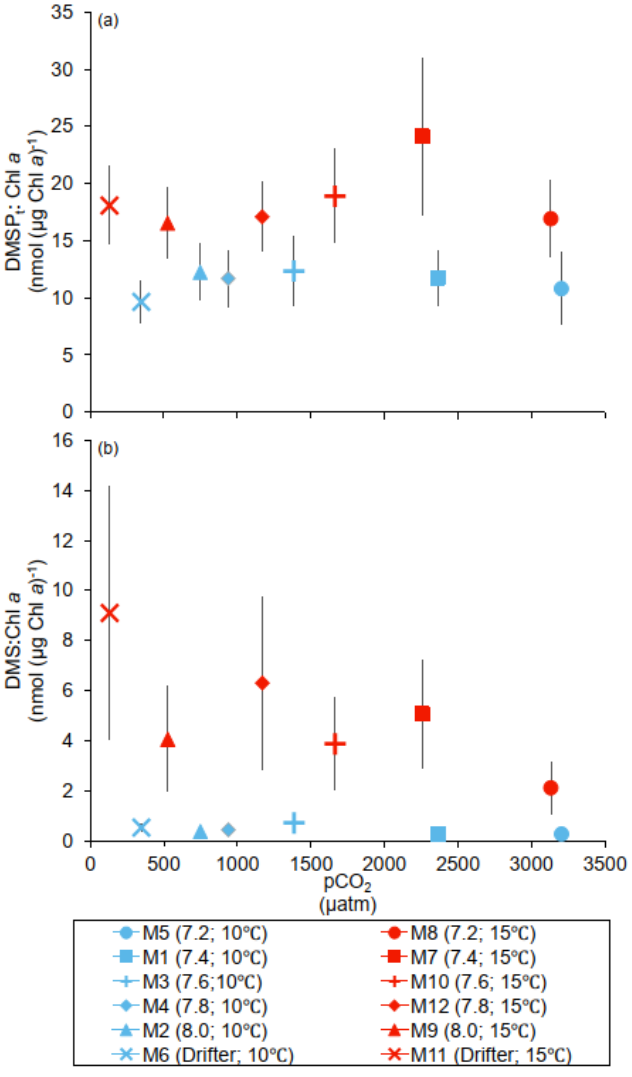

Figure 5. Averages over the course of the experiment (day 0 to day 13) for: (a) $DMSP_t$:Chl *a* ratio, (b) DMS:Chl *a* ratio. For symbol attribution to treatments, see legend.

Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

Table 1. Results of the generalized least squares models (gls) tests for the effects of temperature, pCO_2, and their interaction over the duration of the experiment (day 0 to day 13). Separate analyses with pCO_2 as a continuous factor were performed when temperature had a significant effect. Averages of bacterial abundance and production, $DMSP_t$, DMS, Chl *a*-normalized $DMSP_t$ and DMS concentrations, and $DMS:DMSP_t$ ratios are presented. Natural logarithm transformation is indicated when necessary. Significant results are in bold. *p<0.05, **p<0.01, ***p<0.001.

Response Variable	Factor	df	t-value	p-value
Free-living bacterial abundance ($\times 10^9$ cells L^{-1})	Temperature	8	0.635	0.543
	pCO_2	8	-0.083	0.936
	pCO_2 x Temperature	8	0.221	0.830
Bacterial production (µg C L^{-1} d^{-1})	Temperature	6	**2.454**	**0.050***
	pCO_2 (10°C)	3	-0.272	0.803
	pCO_2 (15°C)	3	0.746	0.510
$DMSP_t$ (nmol L^{-1})	Temperature	8	0.509	0.625
	pCO_2	8	-0.767	0.465
	pCO_2 x Temperature	8	0.134	0.897
DMS (nmol L^{-1})	Temperature	8	**6.822**	**<0.001****
	pCO_2 (10°C)	4	**-4.483**	**0.011***
	pCO_2 (15°C)	4	**-3.799**	**0.019***
$DMSP_t$:Chl *a* ratio (nmol (µg Chl *a*)$^{-1}$)	Temperature	8	**2.627**	**0.030***
	pCO_2	8	0.123	0.908
	pCO_2 x Temperature	8	0.621	0.568
DMS:Chl *a* ratio (nmol (µg Chl *a*)$^{-1}$)	Temperature	8	**5.225**	**<0.001****
	pCO_2 (10°C)	4	-1.373	0.242
	pCO_2 (15°C)	4	-2.227	0.090
Log($DMS:DMSP_t$)	Temperature	8	**5.131**	**<0.001****
	pCO_2 (10°C)	4	-1.844	0.139
	pCO_2 (15°C)	4	**-3.138**	**0.035***

Biogeosciences Discuss., https://doi.org/10.5194/bg-2018-338
Manuscript under review for journal Biogeosciences
Discussion started: 22 August 2018

833 **Table 2. Results of the generalized least squares models (gls) tests for the effects of temperature, pCO_2, and their interaction on the**
834 **maximum values of the parameters measured during the experiment. Separate analyses with pCO_2 as a continuous factor were**
835 **performed when temperature had a significant effect. Maxima of $DMSP_t$, and DMS concentrations are presented. Significant results**
836 **are in bold. *p<0.05, **p<0.01, ***p<0.001.**

837

Response Variable	Factor	df	t-value	p-value
$DMSP_t$ ($nmol\ L^{-1}$)	Temperature	8	0.384	0.711
	pCO_2	8	-0.713	0.496
	pCO_2 x Temperature	8	0.300	0.772
DMS ($nmol\ L^{-1}$)	Temperature	8	**6.403**	**<0.001*** **
	pCO_2 (10°C)	4	**-2.868**	**0.046***
	pCO_2 (15°C)	4	**-4.061**	**0.015***

838