



Drivers of phytoplankton production and community structure in nutrient-poor estuaries receiving terrestrial organic inflow

Joanna Paczkowska^{a,b,c}, Owen F. Rowe^{a,b,d,e}, Daniela Figueroa^{a,b,f}, Agneta Andersson^{a,b,*}

^a Department of Ecology and Environmental Science, Umeå University, SE-901 87, Umeå, Sweden

^b Umeå Marine Sciences Centre, SE-905 71, Hörnefors, Sweden

^c Centro para el Estudio de Sistemas Marinos CESIMAR-CONICET, Blvd. Brown 2915, U9120ACD, Puerto Madryn, Chubut, Argentina

^d Guest researcher: Department of Food and Environmental Sciences, Division of Microbiology and Biotechnology, Viikki Biocenter 1, University of Helsinki, Helsinki, Finland

^e Helsinki Commission, HELCOM Secretariat, Baltic Marine Environment Protection Commission, Helsinki, Finland

^f Swedish Meteorological and Hydrological Institute, SMHI, Göteborg, Sweden

ARTICLE INFO

Keywords:

Phytoplankton
Size-structure
Primary production
Autotrophy
Mixotrophy
Taxonomic richness
Resource use efficiency
Coastal waters
Phosphorus-poor estuaries

ABSTRACT

The influence of nutrient availability and light conditions on phytoplankton size-structure, nutritional strategy and production was studied in a phosphorus-poor estuary in the northern Baltic Sea receiving humic-rich river water. The relative biomass of mixotrophic nanophytoplankton peaked in spring when heterotrophic bacterial production was high, while autotrophic microphytoplankton had their maximum in summer when primary production displayed highest values. Limiting substance (phosphorus) only showed small temporal variations, and the day light was at saturating levels all through the study period. We also investigated if the phytoplankton taxonomic richness influences the production. Structural equation modelling indicated that an increase of the taxonomic richness during the warm summer combined with slightly higher phosphorus concentration lead to increased resource use efficiency, which in turn caused higher phytoplankton biomass and primary production. Our results suggest that climate warming would lead to higher primary production in northerly shallow coastal areas, which are influenced by humic-rich river run-off from un-disturbed terrestrial systems.

1. Introduction

Phytoplankton communities are governed by many limiting and controlling factors, such as nutrient availability, light climate, temperature, salinity, competition, parasites and grazing (Andersson et al., 1996; Calbet, 2001; Dahlgren et al., 2010; Faithfull et al., 2011). In temperate aquatic systems the phytoplankton succession generally starts with a spring bloom dominated by relatively large autotrophic cells, which are favored by high light and nutrient concentrations. As nutrients are depleted and the water warms up, smaller plankton, i.e. autotrophic, mixotrophic and heterotrophic nano- and picoplankton are promoted (Sommer et al., 1986; Andersson et al., 1996; Legrand et al., 2015). These have a competitive advantage at low nutrient conditions due to their high surface to volume (S/V) ratio and thinner diffusion boundary layer (Raven, 1998). However, in systems influenced by terrestrial dissolved organic matter (tDOM) (e.g. lakes), the plankton succession pattern can be reversed. A profusion of tDOM during spring can decrease light availability and promote the growth of heterotrophic bacteria resulting in a weakened or absent phytoplankton bloom, with

maximum primary production rates occurring instead during the warmer summer months when river discharge is lower (Drakare et al., 2002; Figueroa et al., 2016). Under such conditions filamentous cyanobacteria could be promoted due to their capacity for phosphorus storage, atmospheric nitrogen fixation and buoyancy regulation (Paerl and Paul, 2012; Reynolds, 2006). However, the drivers of coastal phytoplankton communities are complex and may depend on the relative influence of river inflow and hydrodynamic interaction with offshore waters.

Many estuaries are highly productive, as phytoplankton growth is nurtured by river borne nutrients (Dorado et al., 2015; O'Boyle and Silke, 2010). However, rivers not only transport nutrients, but also tDOM, including coloured tDOM, to the sea. Studies have shown that tDOM can have both positive and negative effects on primary production (Andersson et al., 2013; Thrane et al., 2014; Seekell et al., 2015). High concentrations of tDOM can limit primary production by absorbing light and reducing the availability of phosphorus and iron, essential factors for phytoplankton growth (Jones, 1992; Carpenter et al., 1998). However, the promotion of phytoplankton growth due to the

* Corresponding author. Department of Ecology and Environmental Science, Umeå University, SE-901 87, Umeå, Sweden.
E-mail address: agneta.andersson@umu.se (A. Andersson).

<https://doi.org/10.1016/j.marenvres.2019.104778>

Received 2 April 2019; Received in revised form 16 August 2019; Accepted 19 August 2019

Available online 21 August 2019

0141-1136/© 2019 The Authors. Published by Elsevier Ltd. This is an open access article under the CC BY-NC-ND license (<http://creativecommons.org/licenses/by-nc-nd/4.0/>).

shielding effect from harmful ultraviolet (UV) radiation, and transportation of bioavailable nutrients, can also occur (Nielsen and Ekelund, 1993; Kissman et al., 2013; Seekell et al., 2015).

Since tDOM includes potentially bioavailable carbon, heterotrophic bacteria and the heterotrophic microbial food web can be favored (Tranvik, 1989; Jansson et al., 2007; Barrera-Alba et al., 2009; Hitchcock and Mitrovic, 2015). Under conditions favoring bacterial growth and production, a higher contribution of potentially mixotrophic flagellates has also been observed, which is explained by their ability to supplement photoautotrophic processes by ingesting bacteria (Jansson, 1996; Bergström et al., 2003; Stoecker et al., 2017). Consequently, the negative impacts of tDOM on phytoplankton growth and the concurrent promotion of heterotrophic bacteria can alter the ecosystem productivity and trophic balance (Sandberg et al., 2004; Andersson et al., 2013). Furthermore, this may affect the size-structure of the food web, which has implications for the number of trophic levels and the food web efficiency (Legendre and Rassoulzadegan, 1995; Havens, 1998; Dahlgren et al., 2010).

A less studied factor that may influence aquatic productivity is taxonomic richness. A few previous studies indicate that productivity and diversity display a unimodal or positive relationship (Irigoien et al., 2004; Korhonen et al., 2011). In unimodal relationship, at low productivity low resource availability would limit the number of species, while at high productivity, the phytoplankton community is dominated by few highly competitive species (Rosenzweig and Abramsky, 1993). An increase in productivity due to higher diversity is explained by higher possibility to contain more productive species (selection effect) as well as species which are complementary in the use of resources (complementarity effects) (Loreau et al., 2001; Loreau et al., 2002). The diversity-productivity relationship is likely to occur both on the geographical (e.g. local vs. regional) and ecological scale (e.g. within vs. between communities) (Waide et al., 1999; Gross et al., 2000). Highly diverse communities may occupy more niches than communities with lower diversity, which in turn might result in higher resource use efficiency (Loreau et al., 2001; Loreau et al., 2002). However, while some studies show a positive correlation between diversity and phytoplankton resource use efficiency (RUE) (Ptacnik et al., 2008), others do not (Hodapp et al., 2015). Thus, there is a need for more studies to get a general understanding of the relationship between taxonomic richness, resource use efficiency and productivity in different ecosystems.

The northern Baltic Sea is strongly influenced by phosphorus-poor riverine inflows with high tDOM concentrations (Kuklinski and Pempkowiak, 2011; Pettersson et al., 1997). In the north tDOM makes up ~80% of the dissolved organic matter pool (Alling et al., 2008). Seasonal variations in river discharge and characteristic differences in the properties of the catchment areas are major factors affecting the supply of organic matter to the sea (Skoog et al., 2011; Asmala et al., 2013). The inflow of tDOM to coastal areas is highest during spring when snowmelt in forest and peatland dominated areas takes place (Pettersson et al., 1997; Råike et al., 2012; Reader et al., 2014). Climate change scenarios indicate that precipitation and thus the inflow of tDOM to the northern Baltic Sea will increase in the future (Meier, 2006). This has the potential to modify the composition and size-structure of the phytoplankton community as well as alter the ratio between primary and bacteria production, with potential consequences for higher trophic levels.

Effects of increasing tDOM to coastal ecosystems, and thus surface water browning, due to climate change may also increase the release of greenhouse gases to the atmosphere (e.g., CO₂, N₂O) as a consequence of shift from net autotrophic to net heterotrophic ecosystem (Wikner and Andersson, 2012; Lapierre et al., 2013; Deininger and Frigstad, 2019). Additionally, heterotrophic bacteria as well as cyanobacteria are considered to be lower food quality for consumers because of a lack of polyunsaturated fatty acids (PUFAs) in their cells, which can decrease the food web efficiency of the system (Harwood and Russell, 1984; Berglund et al., 2007; Deininger and Frigstad, 2019). Higher tDOM

concentrations may also lead to decreased light reaching benthic environment, decreased oxygen production by autotrophs and increased coastal dead zones (Jones, 1992; Andersson et al., 2015a). Changes in tDOM input, stronger stratification and a decreased photic zone may promote filamentous cyanobacterial growth due to their capability for phosphorus storage, atmospheric nitrogen fixation and buoyancy regulation (Ibelings et al., 1991; Pettersson et al., 1993).

The aim of this study was to elucidate factors governing the production, size-structure and nutritional strategy in the phytoplankton community in a sub-arctic estuary with low nutrient concentrations and exposure to seasonal river discharge. We tested 1) what factors are influencing phytoplankton size-structure and production 2) if mixotrophic nanophytoplankton is promoted by heterotrophic bacterial production, and 3) if phytoplankton taxonomic richness is positively correlated to resource use efficiency (RUE) and phytoplankton production. Our results contribute to the understanding of the structure and function of phytoplankton communities in ecosystems heavily influenced by tDOM, and give insights into the potential ecological consequences of climate change in coastal environments.

2. Material and methods

2.1. Field sampling

The study was performed in the sub-arctic Råne estuary, northern Baltic Sea, Sweden (Fig. 1). Monthly sampling was performed at 19 stations from May to August (2011) to encompass the river and its discharge area within the estuary. Henceforth May is referred to as spring while the remaining months are referred to as summer. Station 1 was located at the river mouth while 18 stations were evenly dispersed across the estuary region, with the most seaward station being circa 10 km from the river station sampled (Fig. 1). Water was collected at a depth of 1 m using a Ruttner sampler and transported to the laboratory in shaded 20 l bottles. Additional samples for primary and bacterial

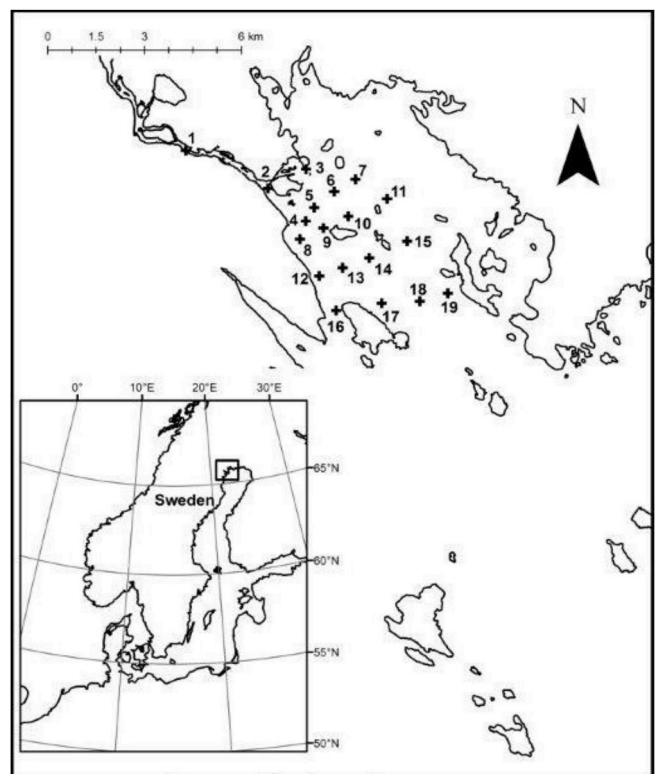


Fig. 1. Map of the study area, northern Baltic Sea, indicating the stations sampled (from Figueroa et al., 2016).

production were also taken at 3 and 5 m (where depth allowed) to determinate depth-integrated production. Temperature (Temp) and Photosynthetically active radiation (PAR) were measured *in situ*. Underwater PAR was recorded every 50 cm in the water column with a Licor LI-1400 connected to Spherical SPQ 1730 sensor, and surface incident PAR was monitored at the Umeå Centre for Marine Sciences (Licor LI-193 spherical quantum sensor). The light attenuation coefficient (Kd) was estimated from the slope of the linear regression of the natural logarithm of down-welling irradiance versus depth. Underwater PAR values recorded at each station (in total 19 stations), were used to calculate average PAR at 1 and 5 m depth for the specific sampling event each month.

Conductivity and pH were measured at 25 °C (Mettler Toledo probes) and *in situ* values were obtained according to Fofonoff and Millard (1983). Salinity was calculated from measurements of *in situ* conductivity. Samples for total nitrogen (Tot N), total phosphorus (Tot P), humic substances (HS), dissolved organic carbon (DOC), coloured dissolved organic matter (CDOM), chlorophyll *a* (Chl *a*), suspended particulate matter (SPM), phytoplankton species composition and biomass were preserved immediately on arrival to the laboratory. Data on river water discharge were obtained from the Swedish Meteorological and Hydrological Institute (SMHI).

2.2. Physicochemical analyses

Tot P and Tot N were measured in unfiltered water samples using a Braan and Luebbe TRAACS 800 autoanalyzer, according to standard analytical methods (Grasshoff et al., 1983). Tot P and Tot N were considered to reflect the nutrients available to phytoplankton. This assumption is based on results obtained in previous studies in the northern part of the Baltic Sea, during which a positive correlation with inorganic form was found. DOC analysis was performed on 0.22 µm filtered (Supor Membrane Syringe Filter, non-pyrogenic; Acrodisc®) and acidified water (18 mM HCl, final concentration). Samples were analyzed on a Shimadzu TOC-5000 analyzer. Measurements of Tot P, Tot N and DOC were performed at an accredited laboratory at Umeå Marine Sciences Center. Humic substances (HS) were determined from unfiltered water samples using a PerkinElmer LS 30 fluorometer at 350/450 excitation/emission wavelengths. Calibration standards were prepared from quinine dihydrogen sulfate dehydrate in 0.05 M sulfuric acid (Hoge et al., 1993; Wedborg et al., 1994). Sulfuric acid (0.05 M) was used as blank.

CDOM absorbance was measured in water samples filtered through a 0.22 µm polycarbonate membrane and stored in amber glass bottles in the dark at 4 °C until analysis. Absorbance values were recorded from 250 to 800 nm using Shimadzu UVPC-2501 scanning spectrophotometer, with Milli-Q water as the blank. The absorption coefficient at 440 nm was calculated by multiplying the absorbance at specific wavelengths with 2.303 and divided by the length of the cuvette (Kirk, 2011).

SPM was measured using the gravimetric method described by Strickland and Parsons (1972). Triplicate 1 l water samples were filtrated through pre-combusted (450 °C) and pre-weighed (W_0) Whatman GF/F filters. Post-sampling, filters were dried for 24 h at 60 °C and re-weighed (W_1). The final concentration of SPM was calculated as the average of triplicates ($W_1 - W_0$).

2.3. Chlorophyll *a* and primary production

Samples for Chl *a* (100 ml) were filtrated onto 25 mm GF/F filters under low pressure and stored at -80 °C until analysis. Chl *a* was extracted in 95% ethanol in the dark overnight at 4 °C. Samples were centrifuged for 10 min to separate ethanol containing chlorophyll *a* from solid material. The concentration of chlorophyll *a* was measured with a PerkinElmer LS 30 fluorometer (433 nm excitation and 674 nm emission wavelengths).

In situ photosynthetic rates of phytoplankton were measured using the ¹⁴C incorporation method. 5 ml of seawater were placed in four 20 ml bottles (three light and one dark) and incubated *in situ* with 7.2 µl ¹⁴C (¹⁴C Centralen Denmark, activity 100 µCi ml⁻¹) for a minimum of 3 h. Post incubation, 100 µl of 5 M hydrochloric acid were added to each tube and samples were ventilated for 12 h. Fifteen ml of scintillation cocktail were added to each sample and samples were measured on a Beckman 6500 scintillation counter. Dissolved inorganic carbon was calculated based on temperature, pH and salinity according to Gargas (1975). Daily net primary production (PP) was calculated using the “light factor method” as described in Gargas (1975) and Andersson et al. (1996).

2.4. Bacterial production

The ³H-thymidine incorporation method was used to measure bacterial production (BP) (Fuhrman and Azam, 1982). Triplicate 1 ml seawater samples (one control and two samples) were incubated with 2 µl of ³H-thymidine (84 Ci mmol⁻¹; PerkinElmer, Massachusetts, USA) (final concentration 24 nM) for 1 h at *in situ* temperature. This thymidine addition corresponded to the saturation level. The control sample was pre-killed by adding 100 µl of ice-cold 50% TCA and incubation at -20 °C for 5 min. Cell production was calculated using a conversion factor of 1.4×10^{18} cells mol⁻¹ of incorporated thymidine (Wikner and Hagström, 1999). Daily net production rates were calculated assuming stable uptake rates over the day and a bacterial carbon content of 20 fgC cell⁻¹ (Lee and Fuhrman, 1987). The assumptions are based on diel experiments and measurements of bacterial cell size in the study area (data not shown).

2.5. Plankton identification and enumeration

Samples for analysis of nano- and microplankton were fixed with 2% acidic Lugol's solution. 10–50 ml samples were settled for 12–48 h in sedimentation chambers. The cells were then counted with an inverted microscope using phase contrast imaging (Nikon Eclipse Ti) (Utermöhl, 1958). Microplankton (> 20 µm) and nanoplankton (2–20 µm) samples were counted at 100x and 400x magnification, respectively. For ciliates, 200x magnification was used. Different taxa and their nutritional characteristics were identified from the cell morphology, size and described trophic (Tikkanen and Willen, 1992; Hällfors, 2004; Olenina et al., 2006). Further, the coloration of the smallest cells was used to support the trophic classification as Lugol's solution stains chlorophyll *a* brown. Cell biovolume of autotrophic, heterotrophic and mixotrophic plankton and the ciliate *Mesodinium rubrum* were calculated according to Olenina et al. (2006) and carbon content was estimated following the Menden-Deuer and Lessard equations (Menden-Deuer and Lessard, 2000).

Picocyanobacteria were analyzed using epifluorescence microscopy, as described in Andersson et al. (1996). The samples were preserved with glutaraldehyde (2% final concentration), filtered (1 ml) onto 0.6 µm black polycarbonate filters and counted on an epifluorescence microscope (Nikon Eclipse TE 2000-U) at 1000x magnification, using green excitation light (510–560 nm, emission wavelength > 590 nm). Cells were counted in 20 randomly positioned fields of view, and a minimum of 300 cells were counted per sample. Biovolume and carbon biomass were estimated as described above.

Cells were grouped into three functional groups (AU: autotrophs, HT: heterotrophs, MX: mixotrophs), and three size categories (picoplankton: < 2 µm, nanoplankton: 2–20 µm, microplankton: > 20 µm), based on measurements of the longest cell axis. Total phytoplankton biomass (TB) was calculated as the sum of the carbon biomass of autotrophs (including *Mesodinium rubrum*) and mixotrophs. The relative biomass proportion of functional groups and size classes was calculated.

Phytoplankton taxonomic richness (S), defined as the number of

taxa found in a sample, was calculated as a proxy of diversity. Phosphorus has been shown to be the main limiting factor for phytoplankton in the studied coastal area in the northern Baltic Sea (Andersson et al., 1996), and therefore the phytoplankton resource use efficiency (RUE) was expressed as a natural logarithm of the ratio between TB and Tot P (RUE_p) (Ptacnik et al., 2008).

2.6. Statistical analyses

Generalized linear mixed-effects model (GLMM) was used to identify relationships between biological and physicochemical variables. Based on variance inflation factor (VIF) results, HS was not included in analysis due to high multicollinearity with other parameters e.g. DOC (VIF > 10). A backwards stepwise elimination process based on Akaike Information Criterion (AICs) was used to remove nonsignificant variables and obtain the final model. Additionally, Spearman's correlation coefficients were calculated between phytoplankton related variables and physicochemical variables. Changes in the phytoplankton composition between months and stations were visualized by non-metric multidimensional scaling (NMDS) based on Bray-Curtis similarity matrix, while analysis of similarity (ANOSIM) was performed to test differences in phytoplankton biomass composition between months. Phytoplankton abundance for both analyses was standardized by sample size. The redundancy analysis (RDA) was conducted to identify main physicochemical and biological variables influencing the size-structure of the phytoplankton. As the results of the variance inflation factor (VIF) indicated that HS was highly correlated with other variables, this variable was excluded from the RDA analysis. Pearson's correlation (r_p) between variables was estimated by RDAs. Forward selection (and a Monte-Carlo permutation test, $n = 999$ permutations) was used to estimate which variables had a significant influence on size-structure of the phytoplankton. Additionally, relationships between phytoplankton taxonomic richness (S), resource use efficiency (RUE_p), total biomass (TB) and primary production (PP) as endogenous variables and total phosphorus (Tot P) and temperature (Temp) as exogenous variable were examined by piecewise structural equation models (piecewiseSEMs). We used the d-separation (d-sep) test to investigate if all pathways in the model were included. Unstandardized path coefficients and R^2 values were calculated, while Fisher's test was used to investigate goodness of fit of the model. Data analyses were performed in R version 3.5.1 using the package 'MASS', 'piecewiseSEM', SPSS Statistics 22, Primer 6 and Canoco 5 softwares.

3. Results

3.1. Physicochemical variables

The river water discharge was highest during the May sampling, with flow rates of $\sim 100 \text{ m}^3 \text{ s}^{-1}$, after which it was lower, $\sim 30 \text{ m}^3 \text{ s}^{-1}$ for the remainder of the study period (Fig. S1, Table 1). Salinity increased from ~ 0.3 to 1 over time due to reduced river inflow (Table 1). Water temperature was $< 8^\circ \text{C}$ in May and increased to $> 15^\circ \text{C}$ in summer, reaching highest values in July (Table 1, Fig. 2D). Tot N and DOC concentrations were highest in May then decreased, and stayed at a similar level during summer (Table 1). A similar temporal trend was also observed for Kd (Table 1). Average Tot P was lowest in May and slightly increased during the remaining months (Table 1, Fig. 2E). The ratio between Tot N and Tot P was highest in May, ~ 90 , then decreased and stayed the same for the remaining months (Table 1). SPM and pH increased from May to August (Table 1). tDOM related variables, such as humic substances and CDOM showed highest values in May and July, concomitant to the river flush (Table 1). Generally, higher values were observed close to the river mouth and lesser at the more seaward locations (data not shown). The lowest average PAR at 1 m was observed in June and the highest in July while remaining at a level of $\sim 100 \mu\text{mol photon m}^{-2} \text{ s}^{-1}$ during the other months (Table 1, Fig. 2F). PAR at 5 m

Table 1

Monthly mean (\pm standard deviation) of physicochemical variables for all sampled stations during the study period.

| | May | June | July | August |
|------------------------------------------------------------|----------------|-----------------|-----------------|-----------------|
| Temp ($^\circ \text{C}$) | 6.7 \pm 0.7 | 15.7 \pm 0.7 | 21.4 \pm 1.0 | 16.5 \pm 0.5 |
| Salinity | 0.3 \pm 0.3 | 0.6 \pm 0.5 | 0.5 \pm 0.4 | 1.0 \pm 0.6 |
| Tot N ($\mu\text{mol l}^{-1}$) | 26.9 \pm 2.9 | 19.7 \pm 2.8 | 20.8 \pm 3.0 | 20.4 \pm 2.3 |
| Tot P ($\mu\text{mol l}^{-1}$) | 0.29 | 0.32 \pm 0.1 | 0.32 | 0.34 \pm 0.1 |
| N:P ratio | 92.0 \pm 9.4 | 63.3 \pm 11.4 | 64.6 \pm 10.4 | 62.4 \pm 9.6 |
| HS ($\mu\text{g l}^{-1}$) | 61.5 \pm 8.2 | 43.8 \pm 10.5 | 53.6 \pm 13.1 | 41.9 \pm 12.5 |
| DOC (mg l^{-1}) | 7.6 \pm 1.6 | 5.6 \pm 0.7 | 6.3 \pm 0.7 | 6.6 \pm 3.3 |
| CDOM (m^{-1}) | 3.0 \pm 0.6 | 2.8 \pm 0.9 | 2.9 \pm 0.8 | 2.1 \pm 0.8 |
| SPM (g m^{-3}) | 3.2 \pm 1.2 | 3.4 \pm 1.5 | 3.5 \pm 1.7 | 3.7 \pm 2.9 |
| Kd (m^{-1}) | 1.8 \pm 0.3 | 1.2 \pm 0.3 | 1.3 \pm 0.4 | 1.2 \pm 0.5 |
| PAR at 1m ($\mu\text{mol photon m}^{-2} \text{ s}^{-1}$) | 101 \pm 62 | 40 \pm 25 | 169 \pm 125 | 104 \pm 84 |
| PAR at 5m ($\mu\text{mol photon m}^{-2} \text{ s}^{-1}$) | 3 \pm 5 | 1.7 \pm 1.7 | 7 \pm 4 | 7 \pm 8 |
| River discharge ($\text{m}^3 \text{ s}^{-1}$) | 99.3 \pm 3.0 | 33.4 \pm 1.4 | 33.1 \pm 1.1 | 24.6 \pm 0.9 |
| pH | 6.9 \pm 0.1 | 7.2 \pm 0.3 | 7.2 \pm 0.2 | 7.4 \pm 0.2 |

was only about 5–10% of that at 1 m depth (Fig. S2A).

3.2. Primary production, total biomass and chlorophyll a

Lowest average primary production at the 1 m level was recorded in May ($\sim 14 \mu\text{g C l}^{-1} \text{ d}^{-1}$), with an increasing trend observed in the following months, reaching $\sim 51 \mu\text{g C l}^{-1} \text{ d}^{-1}$ in August (Fig. 2A). Primary production at 5 m was only about 1% of that at 1 m depth (Fig. S 2B). The total biomass of phytoplankton varied on average between 19.7 and 71.8 $\mu\text{g C l}^{-1}$, with lower values in May than later in the season (Fig. 2B). Temp, salinity and SPM were found to be the main factors influencing primary production, while total biomass was influenced by Tot P, salinity and Tot N (Table 2). The concentrations of Chl a were lowest in May ($\sim 1 \mu\text{g l}^{-1}$) and increased 3-fold in June to August (Fig. 2C). Temperature and DOC were indicated to be major drivers of Chl a (Table 2).

Phytoplankton primary production and heterotrophic bacterial production constituted the total "basal production". The total production was higher in May than during the summer, mainly due to the peak of heterotrophic bacterial production. The relative importance of primary production showed an increasing trend from spring to summer (Fig. S3). This pattern was observed both at 1 m depth level as well as on the depth integrated data (Fig. S3). Phytoplankton primary production constituted $< 10\%$ of the basal production in May, while in August it contributed to 40–60% of the production (Fig. S3).

3.3. Phytoplankton species composition, size-structure and nutritional strategy

The phytoplankton community composition varied between sampled months (ANOSIM global $R = 0.433$, $p < 0.001$), and was influenced by the river plume. NMDS ordination showed that the phytoplankton community structure was relatively similar in the estuary during the summer, while the community sampled in May was separated from other months. Additionally, the phytoplankton community in the river plume (station: 1, 2, 3 and 4) clearly differed from that at the more seaward stations in the summer months, clustering more closely with the spring/May samples (Fig. 3). Diatoms and dinoflagellates dominated the community in spring, while filamentous cyanobacteria dominated the biomass in the summer (Table 3, Fig. 4A).

The phytoplankton size-structure was dominated by micro-phytoplankton ($> 20 \mu\text{m}$) during the entire study period, with a biomass ranging from 1.8 to 133 $\mu\text{g C l}^{-1}$. The proportion of micro-phytoplankton was lowest in May, constituting $\sim 50\%$ of the phytoplankton biomass, later increasing to $\sim 70\%$ (Fig. 4B). Both

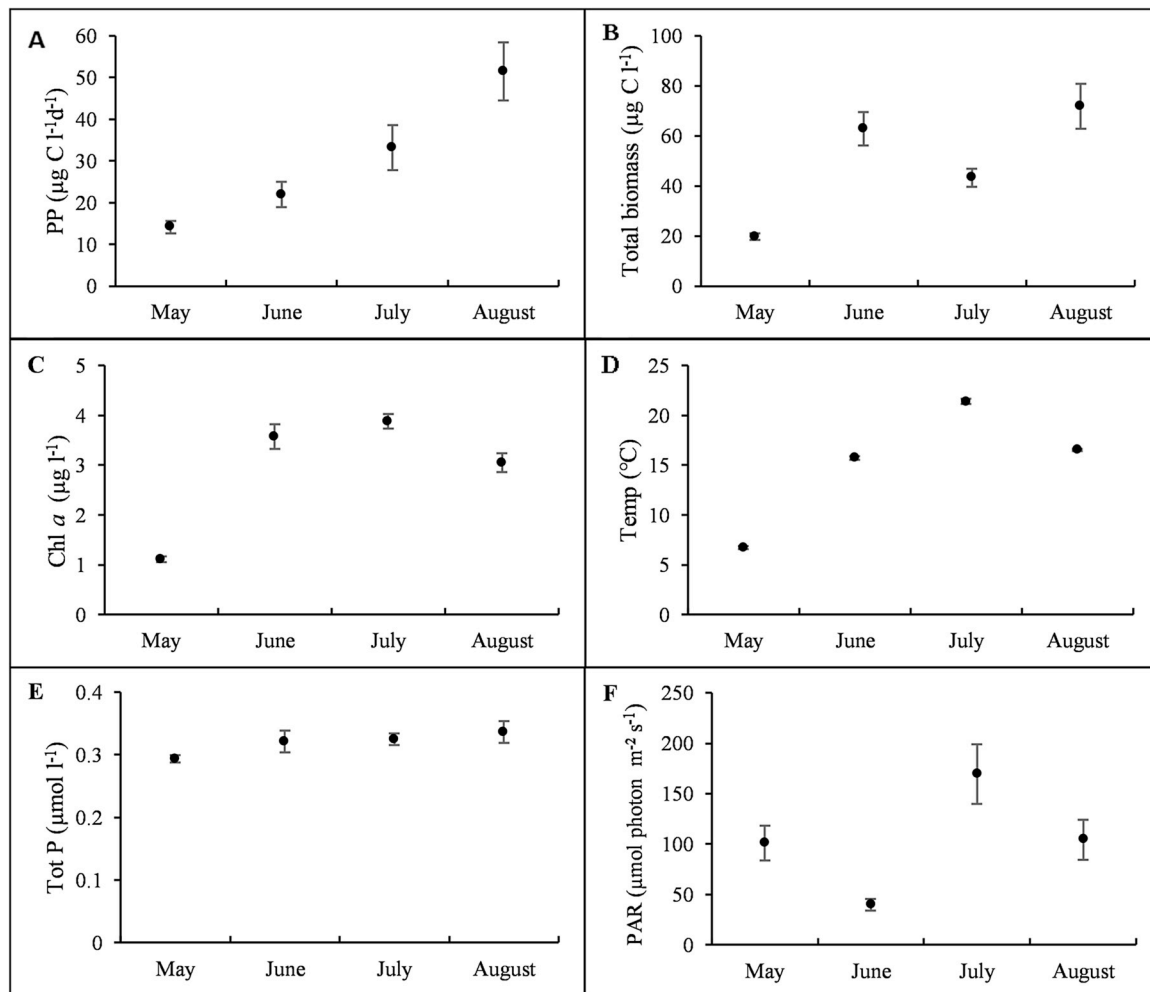


Fig. 2. Monthly average primary production (A), total phytoplankton biomass (B) chlorophyll a (C), temperature (D), total phosphorous (E) and average PAR at 1 m depth (F) in the study area, May–August 2011. Error bars denote standard error.

Table 2

Results of generalized linear models (GLMM) on the physicochemical variables influencing phytoplankton variables during the study period (AIC: Akaike Information Criterion). Phytoplankton variables: primary production (PP), total phytoplankton biomass (TB), chlorophyll a (Chl a), taxonomic richness (S) and resource use efficiency (RUE_p).

| | | Estimate | SD | t-value | p |
|------------------|----------|----------|-------|---------|---------|
| PP | Temp | 0.067 | 0.016 | 4.282 | < 0.001 |
| | Salinity | 1.223 | 0.213 | 5.746 | < 0.001 |
| | SPM | 0.124 | 0.038 | 3.267 | < 0.01 |
| TB | Tot P | 5.690 | 0.835 | 6.817 | < 0.001 |
| | Salinity | 0.399 | 0.126 | 3.183 | < 0.01 |
| | Tot N | -0.062 | 0.020 | -3.026 | < 0.01 |
| Chl a | Temp | 0.179 | 0.017 | 10.336 | < 0.001 |
| | DOC | -0.215 | 0.087 | -2.488 | < 0.05 |
| | Tot P | 1.012 | 0.260 | 3.886 | < 0.001 |
| S | Temp | 0.011 | 0.004 | 2.921 | < 0.01 |
| | Salinity | 0.152 | 0.058 | 2.609 | < 0.05 |
| | CDOM | 0.084 | 0.041 | 2.073 | < 0.05 |
| | Tot P | 2.410 | 0.340 | 7.106 | < 0.001 |
| RUE _p | Salinity | 0.183 | 0.046 | 3.948 | < 0.001 |
| | Tot N | -0.036 | 0.006 | -5.903 | < 0.001 |

biomass and the relative contribution of microphytoplankton correlated with most of the environmental variables, e.g. increased with increasing Temp, Tot P and decreasing with tDOM related variables, Tot N or Kd (Table S1). In May microphytoplankton was dominated by *Wolozynskia*

spp. and *Planktothrix* spp. Filamentous cyanobacteria became dominant members of the community at all post-May sampling events, with a peak of *Planktothrix* spp. and *Aphanizomenon* spp. in August (Table 4). Nanophytoplankton biomass (2–20 µm) varied between 3.6 and 28.1 µg C l⁻¹, and was on average 2-fold lower in May than in August. The highest proportion of nanophytoplankton (2–20 µm) was observed in May, constituting ~32% of the total biomass, with a fairly stable and slightly lower relative contribution recorded in later sampling months (Fig. 4B). Nanophytoplankton was dominated by the genus *Chrysochromulina* during spring and early summer (Table 4). Negative correlations were found between nanophytoplankton biomass and tDOM related variables, Tot N, Kd and water discharge, while a positive correlation was observed in relation to Temp, salinity and pH (Table S1). On the other hand, the proportion of nanoplankton in the total phytoplankton biomass increased with increasing Tot N and water discharge or decreasing Temp, Tot P and pH. Picoplankton (< 2 µm) biomass ranged from 1.5 to 8.9 µg C l⁻¹ with a higher contribution of small cells in May and July than June and August (Fig. 4B). The relationship between picocyanobacterial biomass and temperature and salinity was positive, while a negative correlation was found with water discharge. The relative contribution of picocyanobacteria correlated with most of the environmental variables, e.g. they increased with tDOM related variables and decreased with Tot P (Table S1).

The proportion of autotrophic biomass was high throughout the study period, constituting ~85% in May, increasing up to ~89% in July (Fig. 4C). The opposite trend was found for the mixotrophs. The

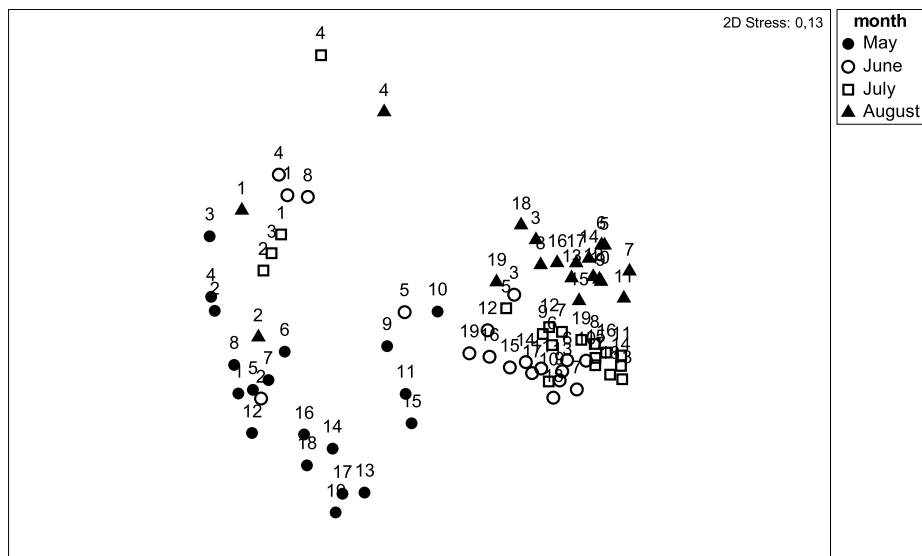


Fig. 3. Non-metric multidimensional scaling (NMDS) of the phytoplankton community in the study area, May–August 2011. Numbers represent different sampling stations, with station 1 being the river and 2 the river mouth (see Fig. 1).

Table 3
Monthly mean relative carbon biomass (%) of different phytoplankton groups in the study area (May–August).

| Class/group | May | June | July | August |
|---------------------------|------|------|------|--------|
| Cyanophyceae ^a | 6.7 | 46.5 | 52.7 | 60.1 |
| Picocyanobacteria | 15.8 | 3.8 | 11.1 | 6.8 |
| Dinophyceae | 16.6 | 3.9 | 1.2 | 0.6 |
| Diatomophyceae | 32.3 | 25.4 | 18.0 | 13.2 |
| Prymnesiophyceae | 11.2 | 9.7 | 5.0 | 3.7 |
| Others | 17.5 | 10.8 | 11.9 | 15.5 |

^a Colony-forming and filamentous cyanobacteria.

absolute biomass of autotrophs, mixotrophs and heterotrophs correlated with most of the physicochemical variables, while the relative biomass of autotrophs and mixotrophs was influenced by temperature, Tot P and PAR (Table S1). Mixotrophs were dominated by *Chrysochromulina* spp. during the whole study period and Dinophyceae was the most abundant class among heterotrophs.

In the RDA model, the first two axes explained 53.2% of the

variance in the size-structure of the phytoplankton ($p < 0.05$). The first RDA axis was strongly positively correlated with Temp and Tot P and negatively with the relative contribution of mixotrophs (0.56, 0.52, and -0.46 , respectively) (Fig. 5, Table S2). It explained the temporal variability of the size-structure of the phytoplankton during the study period. The second RDA axis explained the spatial variability of the size-structure of the phytoplankton and was strongly correlated with the relative contribution of autotrophs and mixotrophs (-0.38 and 0.34 , respectively). Forward selection indicated that temperature and Tot P were the variables statistically significantly shaping the size-structure of the phytoplankton and explained 22.4%, and 11%, respectively, of the total variance (Table 5).

3.4. Phytoplankton diversity and resource use efficiency (RUE)

Taxonomic richness was the lowest in May and the highest in July (Fig. 6A). Temp and Tot P were the most important factors influencing taxonomic richness (Table 2). The average RUEp was the lowest in May (~ 0.32) then increased and remained constant for the rest of the period (Fig. 6B). RUEp was shaped by Tot P, Tot N and salinity (Table 2).

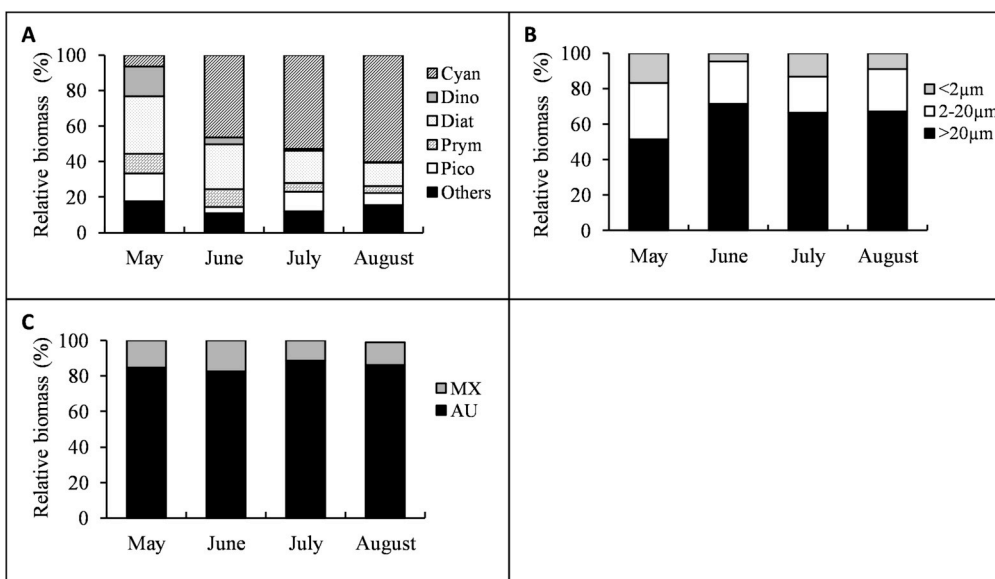


Fig. 4. Monthly average relative phytoplankton biomass (%) of phytoplankton groups: Cyanophyceae (Cyan), Dinophyceae (Dino), Diatomophyceae (Diat), Prymnesiophyceae (Prym), Picocyanobacteria (Pico) and Others (A), size groups pico- ($< 2\mu\text{m}$), nano- ($2\text{--}20\mu\text{m}$), micro- ($> 20\mu\text{m}$) phytoplankton (B), and AU (autotrophs) and MX (mixotrophs) (C) in the study area, May–August 2011.

Table 4
Dominant phytoplankton taxa in different size classes for 19 stations (constituting > 25% of the total carbon biomass), in the study area (May–August).

| Size fraction | Class/Phylum | May | June | July | August |
|---------------|------------------|------------------------------|------------------------------|------------------------------|-------------------------------------------------------|
| < 2 μm | Cyanophyceae | <i>Synechococcus</i> spp. | <i>Synechococcus</i> spp. | <i>Synechococcus</i> spp. | <i>Synechococcus</i> spp. |
| 2–20 μm | Prymnesiophyceae | <i>Chrysochromulina</i> spp. | <i>Chrysochromulina</i> spp. | <i>Chrysochromulina</i> spp. | |
| > 20 μm | Cyanophyceae | | <i>Planktothrix</i> spp. | <i>Planktothrix</i> spp. | <i>Planktothrix</i> spp. <i>Aphanizomenon</i> spp. |

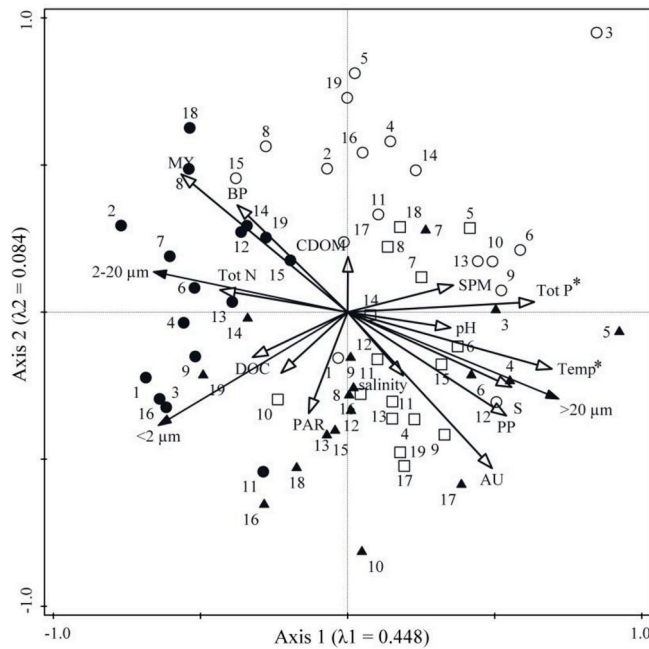


Fig. 5. Redundancy analysis (RDA) ordination plot of the relative contribution of phytoplankton size-structure: pico- (< 2 μm), nano- (2–20 μm), micro- (> 20 μm) phytoplankton, and physicochemical (DOC, CDOM, Tot P, Tot N, Temp, PAR at 1m, SPM, salinity, pH) and biological (primary production (PP), bacterial production (BP), the relative biomass of autotrophs and mixotrophs (AU, MX), taxonomic richness (S)) variables during the study period (● - May; ○ - June, □ - July, ▲ - August). Asterisks indicate statistical significance ($p < 0.05$) of physicochemical variables influencing the relative contribution of phytoplankton size-structure based on RDA, forward selection method.

Table 5
Results of the forward selection of physicochemical and biological variables that significantly influenced the size-structure of phytoplankton during the study period.

| Variables | % explained | p-value | F-value |
|-----------|-------------|---------|---------|
| Temp | 22.4 | 0.015 | 18.5 |
| Tot P | 11.0 | 0.015 | 10.4 |

Furthermore, higher phytoplankton diversity led to more efficient use of phosphorus (Spearman rho = 0.44, $p < 0.001$) (Fig. 6C). SEM analyses showed that both Temp and Tot P had a strong significant impact on richness (respectively, 0.49, $p < 0.001$, 28.37, $p < 0.01$). Moreover RUEP was directly influenced by Temp (0.02, $p < 0.01$) and indirectly by Tot P through richness mediator (0.02, $p < 0.01$) (Fig. 7). Higher Temp and Tot P concentration lead to higher phytoplankton biomass (TB) and primary production (Fig. 7). Overall, SEM models showed a similar goodness of fit to the data (Tot P model: Fisher's $C = 7.0$, $df = 6$, $p = 0.33$, Temp model: Fisher's $C = 7.4$, $df = 6$, $p = 0.29$).

4. Discussion

Our results indicate that the phytoplankton production, size-structure and nutritional strategy were affected in a complex way by the concurrent effects of factors like temperature, Tot P and tDOM variables. The proportion of small cells picocyanobacteria and nanophytoplankton to the total phytoplankton biomass was negatively correlated with Tot P which can be explained by the increased importance of smaller cells under lower nutrient concentrations, due to higher surface-to-volume ratio (Bell and Kalf, 2001; Callieri et al., 2007). It is likely that re-mineralization of phosphorus was higher in the warm summer than in the cold spring, leading to higher P availability in summer. The absolute biomass of pico- and nanophytoplankton, on the contrary, related positively to temperature, which is in agreement with earlier studies (e.g. Andersson et al., 1994; Moran et al., 2010). Species with small cell size, such as *Synechococcus* spp., in general have higher specific growth rates at high temperature (Jöhnk et al., 2008; Paerl and Huisman, 2009). A previous study performed at a coastal location in the northern Baltic Sea, estimated the generation time of picocyanobacteria to be a few days under summer conditions, while it increased to ~120 days during winter (Andersson et al., 1994). Temperature can thus be a direct driver of phytoplankton community composition, however due to a strong covariance with nutrient concentrations, individual effect can often be difficult to distinguish (Li, 1998; Agawin et al., 2000; Mousing et al., 2014). During summer samplings (June–August), lower river discharge impacted only minimally the majority of estuarine stations, and the effective light climate and temperature were relatively high across the majority of the estuarine stations. This is in agreement with studies from shallow humic lakes where temperature and light climate were determined to be the main factors limiting small cells abundance (Jasser and Arvola, 2003).

The proportion of nanophytoplankton correlated positively with water discharge and Tot N. The nanophytoplankton fraction was dominated by the mixotrophic flagellate *Chrysochromulina* spp. during most of the study period. Mixotrophs combine photosynthesis and phagotrophy, which comes with associated metabolic costs, leading to lower reproductive rates compared to single nutritional mode organisms (Rothhaupt, 1996). This means that mixotrophs have a competitive advantage in environments where nutrient concentrations are low, light availability limited and where they can gain nutrients via consumption of bacteria (Hajdu et al., 1996; Jansson et al., 1996; Dahl et al., 2005). Heterotrophic bacteria were promoted by the spring flush and inputs of bioavailable carbon (Figuera et al., 2016), which in turn promoted the mixotrophs, potentially important mediators of bacterivory in coastal waters (Havskum and Riemann, 1996), likely feeding mixotrophically to supplement the constrained availability of nutrients such as P (Nygaard and Tobiesen, 1993; Jansson et al., 1996). The important role of mixotrophy in our study system was further cemented by the positive relationship between mixotrophs and bacterial production, which in turn was influenced by tDOM. Similar relationships have been observed in humic lakes (Drakare et al., 2002; Bergström et al., 2003), indicating that phytoplankton communities in subarctic estuaries are regulated in a similar way as unproductive humic lakes.

The proportion of microphytoplankton increased with higher Tot P concentrations and higher temperature. Diatoms dominated the

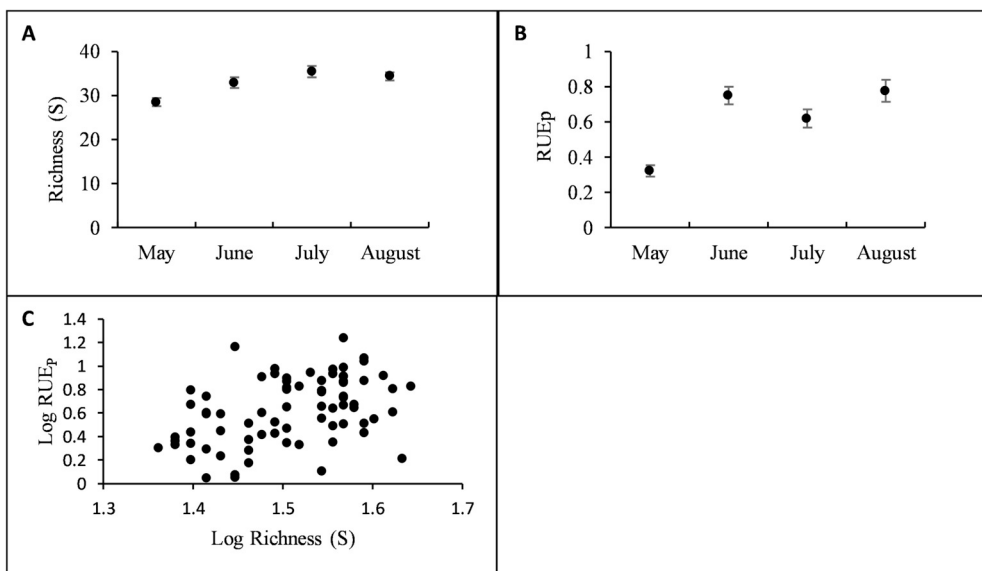


Fig. 6. Monthly average taxonomic richness (A), resource use efficiency (B), and the relationship resource use efficiency vs. taxonomic richness (C) in the study area, May–August 2011. Error bars denote standard error.

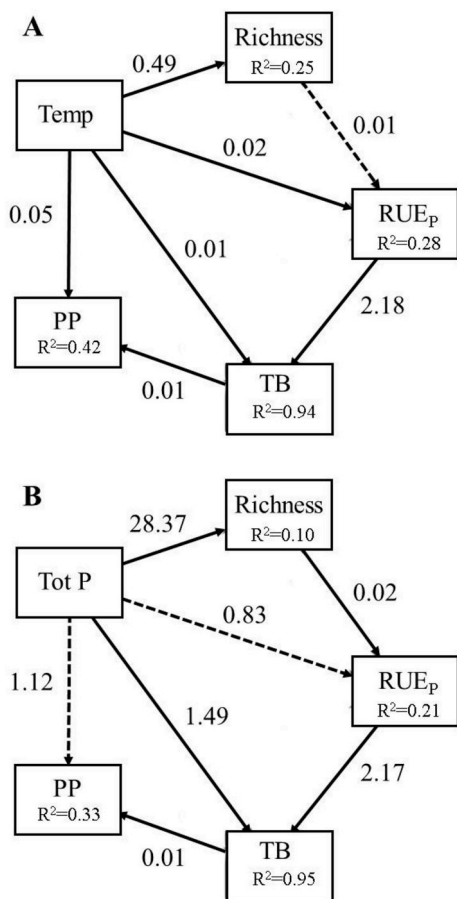


Fig. 7. Path diagram for structural equation model relating temperature, Temp (A), and total phosphorus, Tot P (B), to taxonomic richness, resource use efficiency (RUEp), total phytoplankton biomass (TB) and primary production (PP). The numbers next to each arrow are unstandardized regression coefficients of the SEM. Solid black arrows represent significant paths ($p < 0.05$) while dash arrows non-significant paths ($p > 0.05$).

phytoplankton community during the spring, likely due to water turbulence, high suspended matter load and the high Tot N:P ratio (Margalef, 1978; Kiørboe, 1993). However, we did not observe a spring phytoplankton bloom, probably due to high freshwater inflows which have been shown to counteract spring blooms in shallow coastal systems (Gasiūnaitė et al., 2005). A clear shift from diatoms to filamentous cyanobacteria was observed during the summer, as a consequence of changes in Tot N:P ratio, tDOM related variables and increasing temperature and salinity. This is likely explained by adaptation of large filamentous cyanobacteria to high temperature in combination with a capability to store phosphorous within the cells, nitrogen fixation and buoyant regulation via gas vacuoles. Cyanobacteria generally have growth optima at relatively high temperature, giving them a competitive advantage over diatoms during warmer summer months (Jöhnk et al., 2008). From May to July, conditions favored Oscillatoriales, which in the Baltic Sea comprises non-nitrogen fixing filamentous forms. In August, their contribution decreased while the abundance of Nostocales increased, reaching ~50%. This can be explained by the lower N:P ratio which promoted filamentous cyanobacteria capable of atmospheric nitrogen fixation (*Aphanizomenon* spp. and *Dolichospermum* spp.). Additionally, a positive relationship between cyanobacteria and salinity was found, supporting the idea that salinity can play an important role in shaping the cyanobacterial community (Andersson et al., 2015b).

In estuaries, freshwater inflows can play an important role in regulating the balance between bacterial and primary production due to the transport of nutrients and carbon, and the influence on light availability (Hoch and Kirchman, 1993). The primary production to bacterial production ratio was < 1 in May which indicates that the ecosystem was net heterotrophic, while it switched to net autotrophy by late summer (August). Bacterial production was positively correlated with tDOM related parameters, suggesting that allochthonous carbon carried by water discharge promoted bacterial growth and decoupled them from primary production in the spring while autochthonous produced carbon was the main source of carbon during summer.

The increased primary production from spring to summer seems to have been mediated by higher temperatures and higher Tot P concentrations. tDOM concentration was probably not a major influence on primary production, since average PAR values at 1 m depth were always at saturating levels for primary production, i.e. $> 70 \mu\text{mol quanta m}^{-2} \text{s}^{-1}$ (Andersson et al., 1994). At 5 m depth however, the PAR

values were low ($< 10 \mu\text{mol quanta m}^{-2} \text{ s}^{-1}$) and therefore also the primary production rates were minor, indicating that efficient primary production only occurred in a relatively small and distinct portion of the water column.

Our study indicates that higher temperature and phosphorus concentrations led to increased taxonomic richness, which in turn promoted resource use efficiency and high primary production. This is consistent with previous studies, performed on relatively small area (e.g. less than 10 km) and within a short period of time, where a positive or unimodal diversity-productivity relationship has been found (Chase and Leibold, 2002; Korhonen et al., 2011). The observed trends in community composition across the sampling period and the complex interactions between physicochemical and biological factors are likely indicative of selection processes and complementarity effects (Loreau and Hector, 2001). Additionally, higher phytoplankton diversity led to more efficient use of phosphorus. It confirmed that a more diverse community is able to capture limiting nutrient more efficiently, and higher overall productivity is the result.

In conclusion, our study shows that primary production, the phytoplankton size-structure, and the phytoplankton nutritional strategy in phosphorus-poor estuaries receiving terrestrial input follow patterns more reminiscent of humic lakes than those observed in the open sea. Elevated levels of tDOM related variables and low concentrations of phosphorus favored smaller size cells due to their higher surface to volume ratio and higher light harvesting efficiency. Furthermore, the relative contribution of mixotrophs was higher when basal production was dominated by bacteria, supporting observations that grazing of bacterioplankton can be an important nutrient source under environmental conditions generally perceived as unfavourable or limiting. The decreasing Tot N:P ratio was found to be a main factor shaping changes in the community composition of filamentous cyanobacteria, shifting the community towards nitrogen-fixing species during summer. A positive relationship between phytoplankton taxonomic richness, resource use efficiency and productivity were found. Furthermore, temperature seems to be a dominant factor, as can be expected in regions with high seasonality. Climate change induced increased temperature might therefore lead to increased resource use efficiency and in turn high primary production in shallow coastal areas. On the other hand, browning of the water points towards decreased primary production and a stronger reliance on heterotrophic processes. Both such processes can have potential ecosystem impacts, such as oxygen depletion, an increase in release of greenhouse gases to the atmosphere or a decrease in food quality transfer to higher trophic levels. However, the net effect on primary production and wider ecosystem function is presently difficult to interpret and would need modelling studies.

Acknowledgement

This study was supported by the marine Strategic Research Environment EcoChange (the Swedish Research Council Formas) and the research program WATERS (the Swedish Agency for Marine and Water Management and the Swedish Environmental Protection Agency). We are grateful to the staff at the Umeå Marine Sciences Centre for their expert assistance in the field and laboratory, and for chemical analysis. Jonas Forsberg is gratefully acknowledged for phytoplankton analysis.

Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.marenvres.2019.104778>.

Author contribution

Joanna Paczkowska (JP), Owen Rowe (OR), Daniela Figueroa (DF) and Agneta Andersson (AA) jointly designed the study, performed the

field work and analyzed the samples. JP performed the statistical analyses and wrote the article together with AA. OR and DF commented on the written text.

References

- Agawin, N.S.R., Duarte, C.M., Agusti, S., 2000. Nutrient and temperature control of the contribution of picoplankton to phytoplankton biomass and production. *Limnol. Oceanogr.* 45, 591–600.
- Alling, V., Humborg, C., Morth, C.-M., Rahm, L., Pollehne, F., 2008. Tracing terrestrial organic matter by delta(³⁴S) and delta(¹³C) signatures in a subarctic estuary. *Limnol. Oceanogr.* 53, 2594–2602.
- Andersson, A., Haecy, P., Hagstrom, A., 1994. Effect of temperature and light on the growth of micro- nano- and pico-plankton: impact on algal succession. *Mar. Biol.* 120, 511–520.
- Andersson, A., Hajdu, S., Haecy, P., Kuparinen, J., Wikner, J., 1996. Succession and growth limitation of phytoplankton in the gulf of bothnia (Baltic Sea). *Mar. Biol.* 126, 791–801.
- Andersson, A., Jurgensone, I., Rowe, O.F., Simonelli, P., Bignert, A., Lundberg, E., Karlsson, J., 2013. Can humic water discharge counteract eutrophication in coastal waters? *PLoS One* 8 (4), e61293. <https://doi.org/10.1371/journal.pone.0061293>.
- Andersson, A., Meier, H.E.M., Rippsam, M., Rowe, O.F., Wikner, J., Haglund, P., Eilola, K., Legrand, Figueroa, D., Paczkowska, J., Lindehoff, E., Tysklind, M., Elmgren, R., 2015a. Projected future climate change and Baltic Sea ecosystem management. *Ambio* 44 (3), S345–S356. <https://doi.org/10.1007/s13280-015-0654-8>.
- Andersson, A., Hoglander, H., Karlsson, C., Huseby, S., 2015b. Key role of phosphorus and nitrogen in regulating cyanobacterial community composition in the northern Baltic Sea. *Estuar. Coast Shelf Sci.* 164, 161–171.
- Asmala, E., Autio, R., Kaartokallio, H., Pitkanen, L., Stedmon, C.A., Thomas, D.N., 2013. Bioavailability of riverine dissolved organic matter in three Baltic Sea estuaries and the effect of catchment land use. *Biogeosciences* 10, 6969–6986.
- Barrera-Alba, J.J., Flores Ganesella, S.M., Olivera Moser, G.A., Prado Saldanha-Correa, F.M., 2009. Influence of allochthonous organic matter on bacterioplankton biomass and activity in a eutrophic, sub-tropical estuary. *Estuar. Coast Shelf Sci.* 82, 84–94.
- Bell, T., Kalf, J., 2001. The contribution of picophytoplankton in marine and freshwater systems of different trophic status and depth. *Limnol. Oceanogr.* 46, 1243–1248.
- Berglund, J., Muren, U., Båmstedt, U., Andersson, A., 2007. Efficiency of a phytoplankton-based and a bacteria-based food web in a pelagic marine system. *Limnol. Oceanogr.* 52, 121–131.
- Bergström, A.K., Jansson, M., Drakare, S., Blomqvist, P., 2003. Occurrence of mixotrophic flagellates in relation to bacterioplankton production, light regime and availability of inorganic nutrients in unproductive lakes with differing humic contents. *Freshw. Biol.* 48, 868–877.
- Calbet, A., 2001. Mesozooplankton grazing effect on primary production: a global comparative analysis in marine ecosystems. *Limnol. Oceanogr.* 46, 1824–1830.
- Callieri, C., Modenutti, B., Queimalinos, C., Bertoni, R., Balseiro, E., 2007. Production and biomass of picophytoplankton and larger autotrophs in Andean ultraoligotrophic lakes: differences in light harvesting efficiency in deep layers. *Aquat. Ecol.* 41, 511–523.
- Carpenter, S.R., Cole, J.J., Kitchell, J.F., Pace, M.L., 1998. Impact of dissolved organic carbon, phosphorus, and grazing on phytoplankton biomass and production in experimental lakes. *Limnol. Oceanogr.* 43, 73–80.
- Chase, J., Leibold, M., 2002. Spatial scale dictates the productivity-biodiversity relationship. *Nature* 416, 427–430.
- Dahl, E., Bagoien, E., Edvardsen, B., Stenseth, N.C., 2005. The dynamics of *Chrysochromulina* species in the Skagerrak in relation to environmental conditions. *J. Sea Res.* 54, 15–24.
- Dahlgren, K., Andersson, A., Larsson, U., Hajdu, S., Bamstedt, U., 2010. Planktonic production and carbon transfer efficiency along a north-south gradient in the Baltic Sea. *Mar. Ecol. Prog. Ser.* 409, 77–94.
- Deininger, A., Frigstad, H., 2019. Reevaluating the Role of Organic Matter Sources for Coastal Eutrophication, Oligotrophication, and Ecosystem Health. *Front. Mar. Sci.* 6 <https://doi.org/10.3389/fmars.2019.00210>. UNSP 210.
- Dorado, S., Booe, T., Steichen, J., McInnes, A.S., Windham, R., Shepard, A., Lucchese, A.E.B., Preisichel, H., Pinckney, J.L., Davis, S.E., Roelke, D.L., Quigg, A., 2015. Towards an understanding of the interactions between freshwater inflows and phytoplankton communities in a subtropical estuary in the gulf of Mexico. *PLoS One* 10 (7), e0130931. <https://doi.org/10.1371/journal.pone.0130931>.
- Drakare, S., Blomqvist, P., Bergstrom, A.K., Jansson, M., 2002. Primary production and phytoplankton composition in relation to DOC input and bacterioplankton production in humic Lake Ortrasket. *Freshw. Biol.* 47, 41–52.
- Faithfull, C.L., Bergstrom, A.K., Vrede, T., 2011. Effects of nutrients and physical lake characteristics on bacterial and phytoplankton production: a meta-analysis. *Limnol. Oceanogr.* 56, 1703–1713.
- Figueroa, D., Rowe, O.F., Paczkowska, J., Legrand, C., Andersson, A., 2016. Allochthonous carbon-a major driver of bacterioplankton production in the subarctic northern Baltic Sea. *Microb. Ecol.* 71, 789–801.
- Fofonoff, N.P., Millard, R.C., 1983. Algorithms for computational properties of seawater. UNESCO Tech. Pap. Mar. Sci. 44.
- Fuhrman, J.A., Azam, F., 1982. Thymidine incorporation as a measure of heterotrophic bacterioplankton production in marine surface waters: evaluation and field results. *Mar. Biol.* 66, 109–120.
- Gargas, E., 1975. In: *A Manual for Phytoplankton Primary Production Studies in the Baltic*. Water Quality Institute, Horsholm, Denmark Baltic Marine Biologists. Publ. No. 2.
- Gasiunaitė, Z.R., Cardoso, A.C., Heiskanen, A.S., Henriksen, P., Kauppila, P., Olenina, I., Pilkaitienė, R., Purina, I., Razinkovas, A., Sagert, S., Schubert, H., Wasmund, N., 2005. Seasonality of coastal phytoplankton in the Baltic Sea: influence of salinity and

- eutrophication. *Estuar. Coast Shelf Sci.* 65, 239–252.
- Grasshoff, K., Ehrhardt, M., Kremling, K., 1983. *Methods of Seawater Analysis*, second ed. Verlag Chemie, Weinheim, Germany.
- Gross, K.L., Willig, M.R., Gough, L., Inouye, R., Cox, S.B., 2000. Patterns of species density and productivity at different spatial scales in herbaceous plant communities. *Oikos* 89, 417–427.
- Hajdu, S., Larsson, U., Moestrup, O., 1996. Seasonal dynamics of *Chrysochromulina* species (Prymnesiophyceae) in a coastal area and a nutrient-enriched inlet of the northern Baltic proper. *Bot. Mar.* 39, 281–295.
- Havens, K.E., 1998. Size structure and energetics in a plankton food web. *Oikos* 81, 346–358.
- Harwood, J.L., Russell, N.J., 1984. *Lipids in Plants and Microbes*. George Allen and Unwin, London.
- Havskum, H., Riemann, B., 1996. Ecological importance of bacterivorous, pigmented flagellates (mixotrophs) in the Bay of Aarhus, Denmark. *Mar. Ecol. Prog. Ser.* 137, 251–263.
- Hoch, M.P., Kirchman, D.L., 1993. Seasonal and inter-annual variability in bacterial production and biomass in a temperate estuary. *Mar. Ecol. Prog. Ser.* 98, 283–295.
- Hodapp, D., Meler, S., Muijers, F., Badewien, T.H., Hillebrand, H., 2015. Structural equation modelling approach to the diversity-productivity relationship of Wadden Sea phytoplankton. *Mar. Ecol. Prog. Ser.* 523, 31–40.
- Hoge, F.E., Vodacek, A., Blough, N.V., 1993. Inherent optical properties of the ocean: retrieval of the absorption coefficient of chromophoric dissolved organic matter from airborne laser spectrofluorescence measurements. *Limnol. Oceanogr.* 38, 1394–1402.
- Hitchcock, J., Mitrovic, S., 2015. After the flood: changing dissolved organic carbon bioavailability and bacterial growth following inflows to estuaries. *Biogeochemistry* 124, 219–233.
- Hällfors, G., 2004. Checklist of Baltic Sea phytoplankton species. *Baltic Sea Environ. Proc.* 95.
- Irigoiien, X., Huisman, J., Harris, R.P., 2004. Global biodiversity patterns of marine phytoplankton and zooplankton. *Nature* 429, 863–867.
- Ibelings, B.W., Mur, L.R., Walsby, A.E., 1991. Diurnal changes in buoyancy and vertical distribution in populations of *Microcystis* in two shallow lakes. *J. Plankton Res.* 13, 419–436.
- Jansson, M., Blomqvist, P., Jonsson, A., Bergstrom, A.K., 1996. Nutrient limitation of bacterioplankton, autotrophic and mixotrophic phytoplankton, and heterotrophic nanoflagellates in Lake Ortrasket. *Limnol. Oceanogr.* 41, 1552–1559.
- Jansson, M., Persson, L., De Roos, A.M., Jones, R.I., Tranvik, L.J., 2007. Terrestrial carbon and intraspecific size-variation shape lake ecosystems. *Trends Ecol. Evol.* 22, 316–322.
- Jasser, I., Arvola, L., 2003. Potential effects of abiotic factors on the abundance of autotrophic picoplankton in four boreal lakes. *J. Plankton Res.* 25, 873–883.
- Jöhnk, K.D., Huisman, J., Sharples, J., Sommeijer, B., Visser, P.M., Stroom, J.M., 2008. Summer heatwaves promote blooms of harmful cyanobacteria. *Glob. Chang. Biol.* 14, 495–512.
- Jones, R.I., 1992. The influence of humic substances on lacustrine planktonic food chains. *Hydrobiologia* 229, 73–91.
- Kiorboe, T., 1993. Turbulence, phytoplankton cell-size, and the structure of pelagic food webs. *Adv. Mar. Biol.* 29, 29, 1–72.
- Kirk, J.T.O., 2011. *Light and Photosynthesis in Aquatic Ecosystems*, third ed. Cambridge University Press, Cambridge.
- Kissman, C.E.H., Williamson, C.E., Rose, K.C., Saros, J.E., 2013. Response of phytoplankton in an alpine lake to inputs of dissolved organic matter through nutrient enrichment and trophic forcing. *Limnol. Oceanogr.* 58, 867–880.
- Korhonen, J.J., Wang, J., Soinenen, J., 2011. Productivity-diversity relationships in lake plankton communities. *PLoS One* 6 (8), e22041 0022041. <https://doi.org/10.1371/journal.pone>.
- Kulinski, K., Pempkowiak, J., 2011. The carbon budget of the Baltic Sea. *Biogeochemistry* 8, 3219–3230.
- Lapierre, J.F., Guillemette, F., Berggren, M., del Giorgio, P.A., 2013. Increases in terrestrially derived carbon stimulate organic carbon processing and CO₂ emissions in boreal aquatic ecosystems. *Nat. Commun.* 4, 2972.
- Lee, S., Fuhrman, J., 1987. Relationship between biovolume and biomass of naturally derived bacterioplankton. *Appl. Environ. Microbiol.* 53, 1298–1303.
- Legendre, L., Rassoulzadegan, F., 1995. Plankton and nutrient dynamics in marine waters. *Ophelia* 41, 153–172.
- Légrand, C., Fridolfsson, E., Bertos-Fortis, M., Lindehoff, E., Larsson, P., Pinhassi, J., Andersson, A., 2015. Interannual variability of phyto-bacterioplankton biomass and production in coastal and offshore waters of the Baltic Sea. *Ambio* 44, S427–S438.
- Li, W.K.W., 1998. Annual average abundance of heterotrophic bacteria and *Synechococcus* in surface ocean waters. *Limnol. Oceanogr.* 43, 1746–1753.
- Loreau, M., Hector, A., 2001. Partitioning selection and complementarity in biodiversity experiments. *Nature* 412, 72–76.
- Loreau, M., Naeem, S., Inchausti, P., Bengtsson, J., Grime, J.P., et al., 2001. Biodiversity and ecosystem functioning: current knowledge and future challenges. *Science* 294, 804–808.
- Loreau, M., Naeem, S., Inchausti, P. (Eds.), 2002. *Biodiversity and Ecosystem Functioning — Synthesis and Perspectives*. Oxford University Press.
- Margalef, R., 1978. Life-forms of phytoplankton as survival alternatives in an unstable environment. *Oceanol. Acta* 1, 493–509.
- Meier, H.E.M., 2006. Baltic Sea climate in the late twenty-first century: a dynamical downscaling approach using two global models and two emission scenarios. *Clim. Dyn.* 27, 39–68.
- Menden-Deuer, S., Lessard, E.J., 2000. Carbon to volume relationships for dinoflagellates, diatoms, and other protist plankton. *Limnol. Oceanogr.* 45, 569–579.
- Moran, X.A.G., Lopez-Urrutia, A., Calvo-Diaz, A., Li, W.K.W., 2010. Increasing importance of small phytoplankton in a warmer ocean. *Glob. Chang. Biol.* 16, 1137–1144.
- Mousing, E.A., Ellegaard, M., Richardson, K., 2014. Global patterns in phytoplankton community size structure-evidence for a direct temperature effect. *Mar. Ecol. Prog. Ser.* 497, 25–38.
- Nielsen, T., Ekelund, N.G.A., 1993. Effect of UV-B Radiation and Humic Substances on Growth and Motility of Gyrodinium Aureolum. *Limnology and Oceanography*, vol. 38, pp. 1570–1575.
- Nygaard, K., Tobiesen, A., 1993. Bacterivory in algae: a survival strategy during nutrient limitation. *Limnol. Oceanogr.* 38, 273–279.
- O'Boyle, S., Silke, J., 2010. A review of phytoplankton ecology in estuarine and coastal waters around Ireland. *J. Plankton Res.* 32, 99–118.
- Olenina, I., Hajdu, S., Edler, L., Andersson, A., Wasmund, N., Busch, S., Göbel, J., Gromisz, S., Huseby, S., Huttunen, M., Jaanus, A., Kokkonen, P., Ledaine, I., Niemkiewicz, E., 2006. Biovolumes and size-classes of phytoplankton in the Baltic Sea. *HELCOM Baltic Sea Environ. Proc.* 106, 144.
- Paerl, H.W., Huisman, J., 2009. Climate change: a catalyst for global expansion of harmful cyanobacterial blooms. *Environ. Microbiol. Rep.* 1, 27–37.
- Paerl, H.W., Paul, V.J., 2012. Climate change: links to global expansion of harmful cyanobacteria. *Water Res.* 46, 1349–1363. <https://doi.org/10.1016/j.watres.2011.08.002>.
- Pettersson, K., Herlitz, E., Istvanovics, V., 1993. In: The Role of Gloeotrichia Echinulata in the Transfer of Phosphorus from Sediments to Water in Lake Erken. *Hydrobiologia* 253, pp. 123–129.
- Pettersson, C., Allard, B., Boren, H., 1997. River discharge of humic substances and humic-bound metals to the Gulf of Bothnia. *Estuar. Coast Shelf Sci.* 44, 533–541.
- Ptacinik, R., Solimini, A.G., Andersen, T., Tamminen, T., Brettum, P., Lepistö, L., Willen, E., Rekolainen, S., 2008. Diversity predicts stability and resource use efficiency in natural phytoplankton communities. *Proc. Natl. Acad. Sci. USA* 105, 5134–5138.
- Räike, A., Kortelainen, P., Mattsson, T., Thomas, D.N., 2012. 36 year trends in dissolved organic carbon export from Finnish rivers to the Baltic Sea. *Sci. Total Environ.* 435, 188–201.
- Raven, J.A., 1998. The twelfth Tansley Lecture. Small is beautiful: the picophytoplankton. *Funct. Ecol.* 12, 503–513.
- Reader, H.E., Stedmon, C.A., Kritzberg, E.S., 2014. Seasonal contribution of terrestrial organic matter and biological oxygen demand to the Baltic Sea from three contrasting river catchments. *Biogeochemistry* 11, 3409–3419.
- Reynolds, C.S., 2006. *Ecology of Phytoplankton (Ecology, Biodiversity and Conservation)*. Cambridge University Press, Cambridge, UK.
- Rothhaupt, K.O., 1996. Utilization of substitutable carbon and phosphorus sources by the mixotrophic chrysophyte *Ochromonas* sp. *Ecology* 77, 706–715.
- Rosenzweig, M.L., Abramsky, Z., 1993. How are diversity and productivity related? In: *Species Diversity in Ecological Communities: Historical and Geographical Perspectives*. University of Chicago Press, Chicago, pp. 52–65.
- Sandberg, J., Andersson, A., Johansson, S., Wikner, J., 2004. Pelagic food web structure and carbon budget in the northern Baltic Sea: potential importance of terrigenous carbon. *Mar. Ecol. Prog. Ser.* 268, 13–29.
- Seekell, D.A., Lapierre, J.F., Ask, J., Bergstrom, A.K., Deininger, A., Rodriguez, P., Karlsson, J., 2015. The influence of dissolved organic carbon on primary production in northern lakes. *Limnol. Oceanogr.* 60, 1276–1285.
- Skoog, A., Wedborg, M., Fogelqvist, E., 2011. Decoupling of total organic carbon concentrations and humic substance fluorescence in an extended temperate estuary. *Mar. Chem.* 124, 68–77.
- Sommer, U., Gliwicz, Z.M., Lampert, W., Duncan, A., 1986. The PEG-model of seasonal succession of planktonic events in fresh waters. *Archiv Fur Hydrobiologie* 106, 433–471.
- Stoecker, D.K., Hansen, P.J., Caron, D., Mitra, A., 2017. Mixotrophy in the marine plankton. In: *Annual Review of Marine Science*, vol. 9, pp. 311–335.
- Strickland, J.D.H., Parsons, T.R., 1972. second ed. *A Practical Handbook of Seawater Analysis Bulletin 167 Fisheries Research Board of Canada*, Ottawa.
- Thrane, J.E., Hessen, D.O., Andersen, T., 2014. The absorption of light in lakes: negative impact of dissolved organic carbon on primary productivity. *Ecosystems* 17, 1040–1052.
- Tranvik, L.J., 1989. Bacterioplankton growth, grazing mortality and quantitative relationship to primary production in a humic and a clearwater lake. *J. Plankton Res.* 11, 985–1000.
- Tikkanen, T., Wille'n, T., 1992. *Phytoplankton Flora*. The Swedish Environmental Protection Agency, Solna (in Swedish).
- Utermöhl, H., 1958. Zur Vervollkommnung der quantitativen Phytoplankton-Methodik. *Verhandlungen der Internationalen Vereinigung für Theoretische und Angewandte Limnologie* 9, 1–38.
- Waide, R.B., Willig, M.R., Steiner, C.F., Mittelbach, G., Gough, L., Dodson, S.I., et al., 1999. The relationship between productivity and species richness. *Annu. Rev. Ecol. Systemat.* 30, 257–300.
- Wedborg, M., Skoog, A., Fogelqvist, E., 1994. Organic carbon and humic substances in the Baltic Sea, kattegat and skagerrak. In: Sense, N., Miano, T.M. (Eds.), *Humic Substances in the Global Environment and Implications on Human Health*. Elsevier Science, Amsterdam, pp. 917–924.
- Wikner, J., Andersson, A., 2012. Increased freshwater discharge shifts the trophic balance in the coastal zone of the northern Baltic Sea. *Glob. Chang. Biol.* 18, 2509–2519. <https://doi.org/10.1111/j.1365-2486.2012.02718.x>.
- Wikner, J., Hagstrom, A., 1999. Bacterioplankton intra-annual variability: importance of hydrography and competition. *Aquat. Microb. Ecol.* 20, 245–260.