

Research Article

Spectroscopic and Microscopic Characterization of Volcanic Ash from Puyehue-(Chile) Eruption: Preliminary Approach for the Application in the Arsenic Removal

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Volcanic ash from Puyehue Cordon Caulle Volcanic Complex (Chile), emitted on June 4, 2011, and deposited in Villa La Angostura at ~40 km of the source, was collected and analyzed by Raman spectroscopy, optical and scanning electron microscopy (SEM-EDS), X-ray diffraction (XRD), surface area (BET), and chemical analysis (ICP-AES-MS technique). The mineralogical and physicochemical study revealed that the pyroclastic mixture contains iron oxides in the form of magnetite and hematite as well as pyroxene and plagioclase mineral species and amorphous pumiceous shards. Carbonaceous material was also identified. Physicochemical techniques allow us to select two representative samples (average composition and Fe-rich materials) which were used to analyze their performances in the adsorption process to remove arsenic from water. Additional iron activation by means of ferric salts was performed under original sample. Results showed that the low-cost feedstock exhibited a good adsorption capacity to remove the contaminant, depending on the iron content and the water pH.

1. Introduction

The south-andes Cordillera is one of the world regions with intense volcanic and tectonic activity. The last eruption, started on June 4, 2011, was associated with the activity of Puyehue Cordon Caulle Volcanic Complex (PCCVC), Chile, located at 40°34'57"S-72°06'53"W, emitting more than 5×10^6 m³ of pyroclastic material [1]. The dominance of "west-erlies" regional winds was responsible for the dispersion and sedimentation of ash on a large region along eastern South America, particularly Patagonia and Argentine Pampean Plain.

Although the preliminary characterization of the ash showed a composition predominantly rhyolitic, the ejected material was composed of different types of particles. Eruption stages and distance from the volcanic source affected the chemical composition of the mixture, but, in general, the material was nearly amorphous to the XRD, with a texture characterized by the presence of vesicles. However, in

the early stages of the volcanic event a dark (brownish-black) and heavy particulate matter, with variable size and hardness, was emitted [2].

The application of the volcanic ash depends on several factors such as the mineralogy and chemical composition [3–5], so that the thorough understanding of physicochemical properties of the phases present in the feedstock can play a decisive role in the evaluation of their technological potential.

As for arsenic removal from contaminated aqueous medium, different iron-rich minerals can be used [6, 7]; the chemical affinity between Fe-O-H-arsenate groups can be applied to develop low-cost technologies, particularly useful in rural zones without any other water sources for human consumption [8, 9]. In fact, the presence of arsenate (V) is widespread in groundwater, reaching values exceeding those recommended by the WHO ($10 \mu\text{g L}^{-1}$), situation that leads to many chronic health problems [10]. In this sense, Argentina is one of the most affected countries of South America [11], where the contamination has become a serious problem.

The aim of this work is to study the physicochemical behavior of the ash and tephras deposited in Villa La Angostura ($40^{\circ}45'48''\text{S}$ - $71^{\circ}38'46''\text{W}$) and collected five months after the eruption and the need to approach the potential application (in the original form and activated with ferric phases) to remove arsenic from groundwater.

Scanning electron microscopy (SEM-EDS), optical measurements, surface (BET) and chemical analysis (by inductively coupled plasma (ICP) technique for major and trace elements), X-ray diffraction, and Raman spectroscopy were directed to the mineralogical and chemical characterization of the material.

2. Materials and Methods

The material from Villa La Angostura (about 40 km from the volcanic source), having particle sizes between 10 and $3000\ \mu\text{m}$, was dried at 100°C and identified as pyroclastic material (PM). No thermal transformations were observed in the 60 – 100°C range. The humidity elimination ($\sim 30\%$ of water) was necessary to perform the physicochemical characterization. The selection of the mixture components was done with the aid of the binocular microscope. Glass and light shards with vesicular microcavities were identified as pumiceous-type material (PT), whereas the dark particles, also vesiculated, were named as scoria-type (ST) material. Lithoclasts of size lower than the PT and ST materials, of intense black color and smooth uniform surfaces, were identified as rounded blocks (RB). The material densities followed the sequence $\text{RB} > \text{ST} > \text{PT}$. The former presents also a magnetic behavior. Additional optical measurements were useful to reveal the difference between crystalline and vitreous particles.

PT and ST materials were washed with ethanol by the ultrasonic technique to facilitate the separation of very thin adhered particles and then dried at 80°C .

Scanning electron microscopy and electron diffraction spectroscopy (SEM-EDS) measurements were performed in an ESEM (FEI Quanta 200), with tungsten filament and an ETD (high vacuum secondary electron) detector. Microanalysis was carried out with an EDAX Detector Apollo 40. Chemical results were expressed as % oxides.

Chemical analysis was performed by ICP-AES for major elements (expressed as % oxides) and ICP-MS for trace Rare Earth Elements (REE) (in ppm) (ALS Chemex Lab., Canada). The geochemical behavior was analyzed from variance diagrams of REE normalized to chondrite igneous system.

The BET surface area was measured by N_2 adsorption using a Micromeritics ASAP 2020 Automated Brunauer-Emmett-Teller Sorptometer.

X-ray diffraction patterns for crystalline phase analysis were collected with a Philips PW 1710 diffractometer, $\text{Cu K}\alpha$ Ni-filtered radiation.

Raman spectroscopic analyses were carried out with inVia Renishaw micro-Raman spectrometer equipped with an air-cooled CCD detector and edge filters. A $785.0\ \text{nm}$ emission line from a diode laser was focused on the sample by a Leica DLML microscope, using 5x or 20x objectives.

The power of the incident beam is about 5 mW. Five 10 s accumulations were generally acquired for each sample. The resolution was $2\ \text{cm}^{-1}$, and the spectra were calibrated using the $520.5\ \text{cm}^{-1}$ line of a silicon wafer. Spectral analysis was done by background subtraction and curve fitting.

Preliminary tests for As removal by adsorption were carried out at room temperature (20°C) using PM material (up to $3000\ \mu\text{m}$) and ST material (about 1000 – $2000\ \mu\text{m}$) without grinding. The chemical activation of PM material by means of ferric chloride to give PMA sample was made following the procedure reported by Chen et al. [12]. The batch experiments to study the removal of As from solution were carried out by reacting 10 g solid with 100 mL contaminated water. The arsenic-containing solution of $134\ \mu\text{g L}^{-1}$, simulating conditions of affected groundwater in Buenos Aires province [13], was obtained from hydrated sodium arsenate ($\text{Na}_2\text{HAsO}_4 \cdot 7\text{H}_2\text{O}$) and deionized water. The mixture was kept at room temperature for 1 h, by stirring at 200 rpm to ensure complete homogenization. The slurries were kept at room temperature for 24 h. The supernatant was filtered through a $0.45\ \mu\text{m}$ membrane filter. The initial and final concentration of arsenic in the solution was determined by a Perkin Elmer Analyst 200 (equipped with a Perkin Elmer HGA 900 Graphite-Furnace). The As retained in the solid was calculated according to the expression:

$$\text{Re} = \frac{([C_0 - C_r] \times 100)}{C_0}, \quad (1)$$

where Re is the arsenate removal (%) and C_0 and C_r are the initial and residual (after 24 h) As concentrations, expressed in $\mu\text{g L}^{-1}$.

The pH of suspensions was adjusted between 3 and 9 by using 0.1 M solutions of HCl or NaOH. The values were monitored by means of the Denver Instrument Ultrabasic Benchtop pH meter. In order to maintain a relatively constant ionic strength, the arsenic solutions contain 0.01 M NaCl as background electrolyte. The experiments were done by triplicate.

3. Results and Discussion

3.1. Physicochemical Characterization. According to the reflection microscope, the volcanic material shows a predominance (70–80%) of pumice fragments (glass blabbing). To a lesser extent (30–20%) crystals and crystal fragments of metal oxide and silicate phases were observed.

Figure 1(a) shows the SEM images of the ejected volcanic pyroclastic mixture (PM), whereas Figure 1(b) corresponds to the major component named pumiceous type (PT). Typical SEM images of dark particles, classified as scoria type (ST), in the original form and after ultrasonic washing (ST_w) to remove fine particles (adhered at the surface and filling the vesicles), are shown in Figures 1(c) and 1(d) with the correlative images at different magnification. Black rounded blocks (RB), also at different magnification, are shown in Figure 1(e). It is possible to notice that the texture and morphology among PT, ST, and RB samples were different.

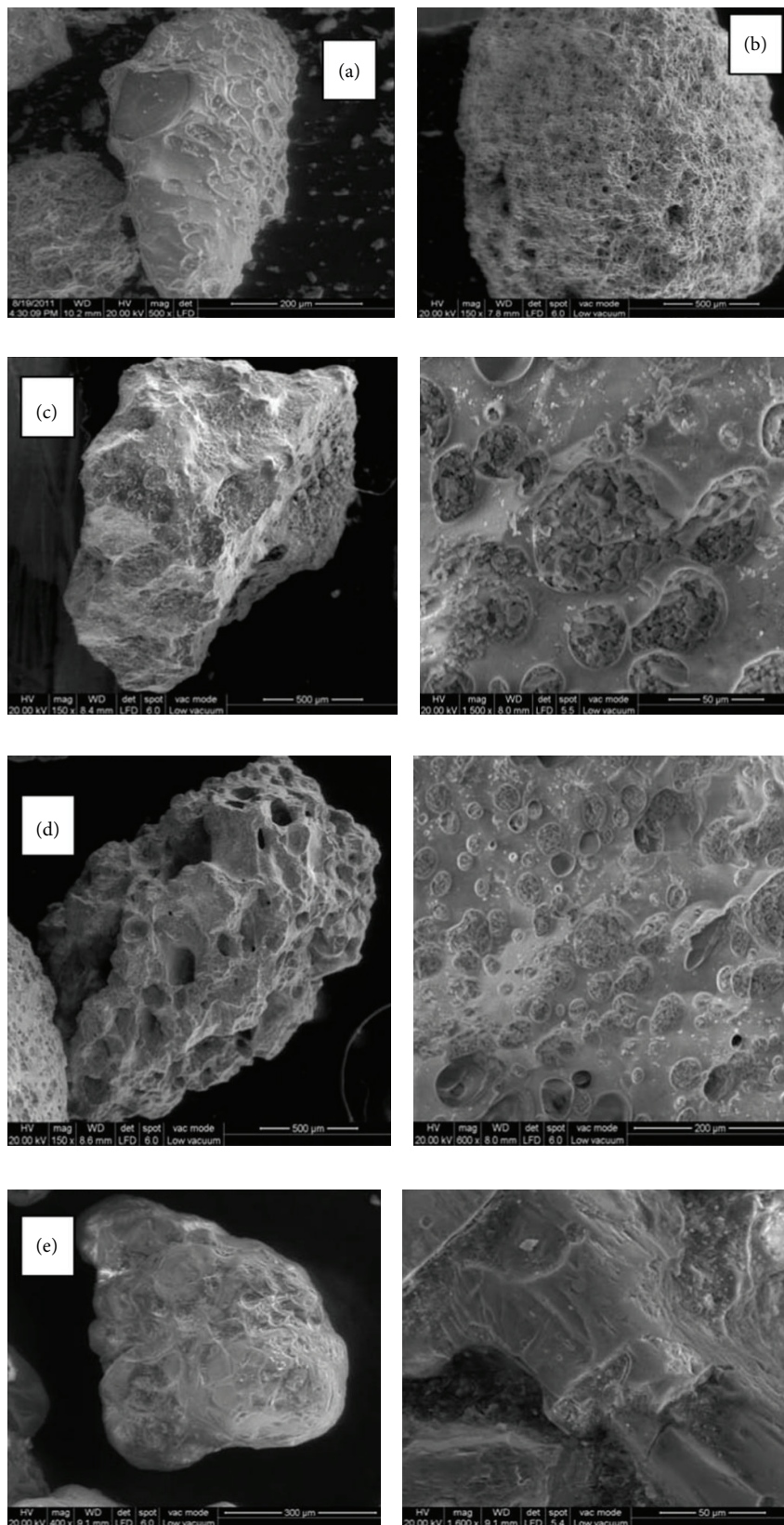


FIGURE 1: SEM micrographs of samples (a) PM, (b) PT, (c) ST, (d) ST_w, and (e) RB. PM: volcanic pyroclastic mixture; PT: pumiceous type; ST: scoria type; ST_w: washed scoria type; RB: rounded block.

TABLE 1: EDS data for volcanic pyroclastic mixture (PM) and selected typical components.

% oxide	PM	ST	ST _w	RB	PT	PT _w
CO ₂	4.12	7.30	6.15	12.56	3.64	2.44
Na ₂ O	4.65	4.05	3.44	Nd	5.58	5.37
MgO	2.08	2.82	3.02	10.07	1.41	0.57
Al ₂ O ₃	14.85	17.01	17.89	4.72	13.84	14.52
SiO ₂	61.25	54.56	51.60	37.48	62.76	68.28
K ₂ O	1.48	1.21	0.60	Nd	3.19	2.36
CaO	2.75	4.90	5.61	3.70	2.52	1.85
TiO ₂	1.01	0.88	1.57	4.47	1.07	1.06
Fe ₂ O ₃	7.81	7.27	10.12	27.00	5.99	3.55

PM: volcanic pyroclastic mixture; ST: scoria type; PT: pumiceous type; RB: rounded block; w: washed.

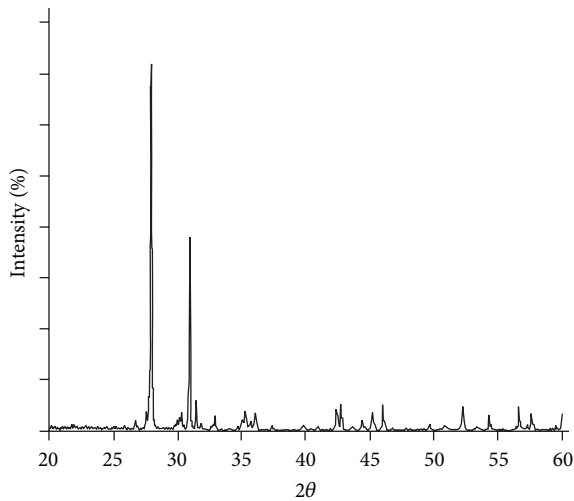


FIGURE 2: XRD pattern of ST sample (see text).

TABLE 2: Major and minor elements (expressed as % oxides) by ICP AES.

Sample	PM	PT	ST
SiO ₂	67.03	70.30	60.52
TiO ₂	0.89	0.65	1.17
Al ₂ O ₃	14.17	13.35	14.52
Fe ₂ O ₃	5.58	4.56	8.92
MnO	0.14	0.12	0.16
MgO	1.05	0.56	2.77
CaO	2.95	2.02	5.76
Na ₂ O	5.04	5.12	3.77
K ₂ O	2.37	2.52	1.49
P ₂ O ₅	0.17	0.13	0.33
Cr ₂ O ₃	0.01	0.01	0.01
BaO	0.08	0.08	0.05
SrO	0.02	0.02	0.04
LOI	0.50	0.56	0.49

PM: volcanic pyroclastic mixture; ST: scoria type; PT: pumiceous type.

Table 1 gives typical EDS results (average of ten determinations). The iron content of the samples increases in

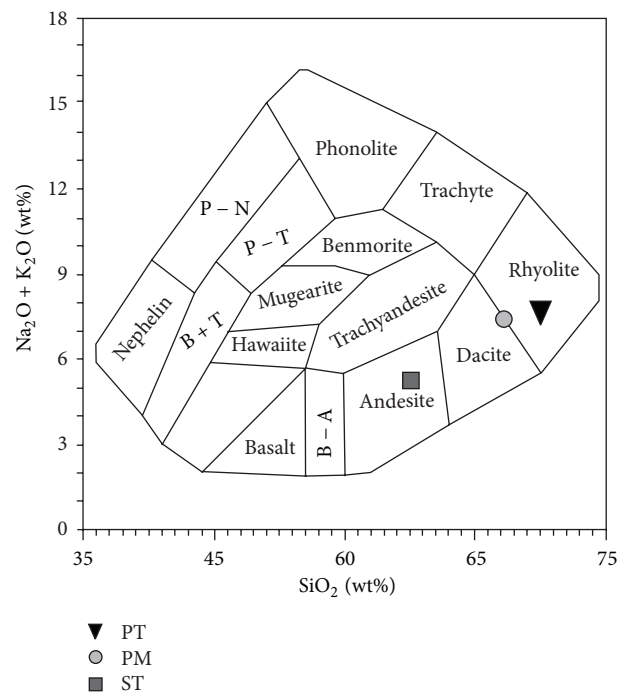


FIGURE 3: Total alkali-silica (TAS) plot for PT, PM and ST samples.

the order $PT < ST < RB$, whereas the PM pyroclastic material reveals the contribution of the mixture components. Rounded blocks (RB), of intense black color, are the Fe and Ti richest samples. Dark color can be associated with the Fe(II)-Fe(III) intervalence charge transfer process [14].

According to EDS results, the SiO_2/Al_2O_3 ratio in the ST particles is near 3, value that increases to ~ 5 in the PT samples. The diminution of Al content can be associated with the enrichment of Si species (crystalline or glassy). Likewise, the RB particles are characterized by the absence of alkaline elements (K and Na) and the presence of Mg and Ca. This fact suggests the existence of mixed valence iron oxides or related phases (magnetite-type) containing Fe, Mg, Al, and Ti. The presence of Ca can be associated with some silicate (probably plagioclase). It is evident that the finest particles are mobilized by the ultrasonic treatment, affecting the relative

TABLE 3: REE trace elements (expressed as ppm) by ICP MS.

Sample	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu
PM	29	65.7	9	40	8.13	2	8	1	8	1.71	5.35	0.82	5.4	0.87
PT	31	69.6	9	37	8.57	2	8	1	8	1.86	5.45	0.88	5.8	0.92
ST	20	45.5	6	26	6.51	2	6	1	6	1.35	3.98	0.60	3.8	0.59

PM: volcanic pyroclastic mixture; ST: scoria type; PT: pumiceous type.

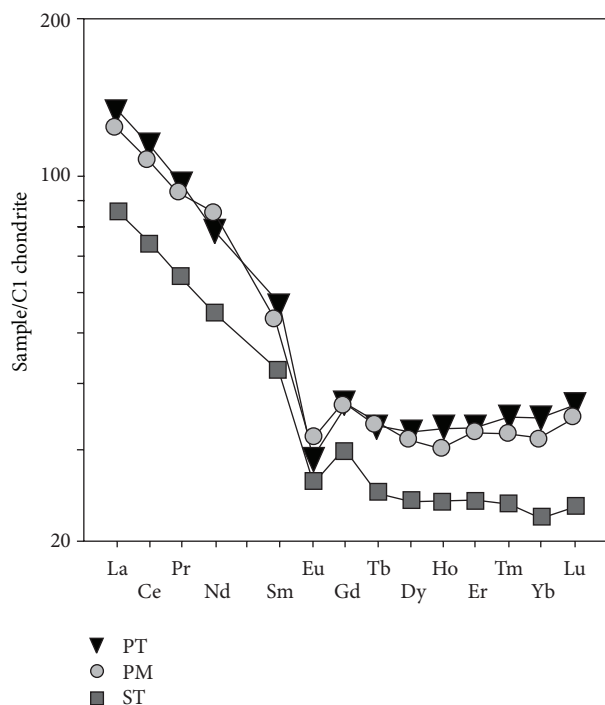


FIGURE 4: Spider diagram of REE for PT, PM, and ST samples.

Si/Fe composition. The washing of ST samples facilitates the exposition of iron species by elimination of Si-rich particles adhered in vesicles. Finally, the presence of carbonaceous materials is observed particularly in the dark particles.

Moreover, Table 2 presents the bulk chemical analysis by ICP-AES (major and minor elements), whereas Table 3 gives the ICP-MS data for REE trace elements. Comparison between bulk and EDS chemical data suggests that iron phases are concentrated in the surface of siliceous materials. On the other hand, round blocks of iron phases can be isolated and also viewed immersed in the glassy matrix, as it is observed in Figure 1(a).

The X-ray diffraction pattern of the volcanic pyroclastic mixture (PM) indicates the presence of an absolute majority of a vitreous amorphous phase, whereas the PT samples is also characterized by a very low resolution. However, the XRD pattern of the dark-ST material, shown in Figure 2, reveals the presence of crystalline phases. X-ray diffraction lines can be correlated to siliceous phases in the form of substituted calcium plagioclase ($\text{CaAl}_2\text{Si}_2\text{O}_8$) and pyroxene ($\text{Mg}_2\text{Si}_2\text{O}_6$) (PDF 89-1463, 85-1740, 88-2377), both showing the stronger peaks in the zone of $\sim 28^\circ$, $\sim 30^\circ$ of 2θ angles. Likewise,

iron oxides, observed as weak reflections, can be assigned to binary and/or mixed oxide, for example, magnetite-type structures PDF 80-0390 (more intense lines at $\sim 30^\circ$, $\sim 35^\circ$ of 2θ) as well as hematite or related phases PDF 89-2810 (stronger reflections at $\sim 33^\circ$, $\sim 35^\circ$ of 2θ). For the ST sample, containing the majority of crystalline phases, it is possible to estimate the presence of plagioclase as major phase (~ 50 – 60%) with ~ 10 – 14% of pyroxene, $\sim 10\%$ of iron oxides, and a low proportion ($<5\%$) of K-feldspar, not detected clearly by XRD. The remainder of material is vitreous.

According to the total alkali versus silica diagram (TAS) for the volcanic rocks, shown in Figure 3, obvious signals of fractionation process are observed, whereas ST materials are located in the andesite-type stability field, the PT fragments belong to rhyolite type. Alkali depletion observed in ST samples can be related to the concentration thereof in the melted phase.

Figure 4 shows the spider diagram for the samples obtained from REE trace elements. The REE depletion observed for ST samples (more pronounced for the heavy elements HREE) can be attributed to dissimilar degree of fractionation. This can be associated not only with the chemistry of the source but also with the crystal-melt equilibrium during the magmatic evolution. Partition coefficients are higher with increasing silica content of the melt. Likewise, the Eu(II) negative anomaly can be attributed to its compatibility with feldspars and silicate species (similar ionic potential for Eu(II) and alkaline-earth species). So, differences between the fractionation of the light REE (LREE) with respect to HREE may be related to the presence of other silicate species such as pyroxenes and accessory phases. The behavior observed in Figure 4 is similar to that reported for volcanic materials from previous eruptions of the PCCVC [15].

Surface area (BET) of the pyroclastic material (PM) is $2.23 \text{ m}^2 \text{ g}^{-1}$, while the ST sample presented a relatively lower value ($1.73 \text{ m}^2 \text{ g}^{-1}$). The difference can be attributed to the presence of crystalline phases [16].

The micro-Raman spectroscopy is a useful technique to define the mineralogy of composite materials, becoming a sensor for identifying promptly the species, isolated or embedded in the matrix [17]. Figures 5(a)–5(d) correspond to the Raman spectra of samples collected from various sample spots. The presence of a complex composition mixture is revealed: alkali aluminosilicate, hematite, and magnetite associated with siliceous phases and carbonaceous material dispersed in the glassy matrix. Hence, the Raman spectra cover a wide range of Na, K, Mg, and Ca silicates [18]. Experimental Raman data of Figure 5(a), with lines at about 325 , 380 , 530 , 665 , 888 , and 999 cm^{-1} , corroborate

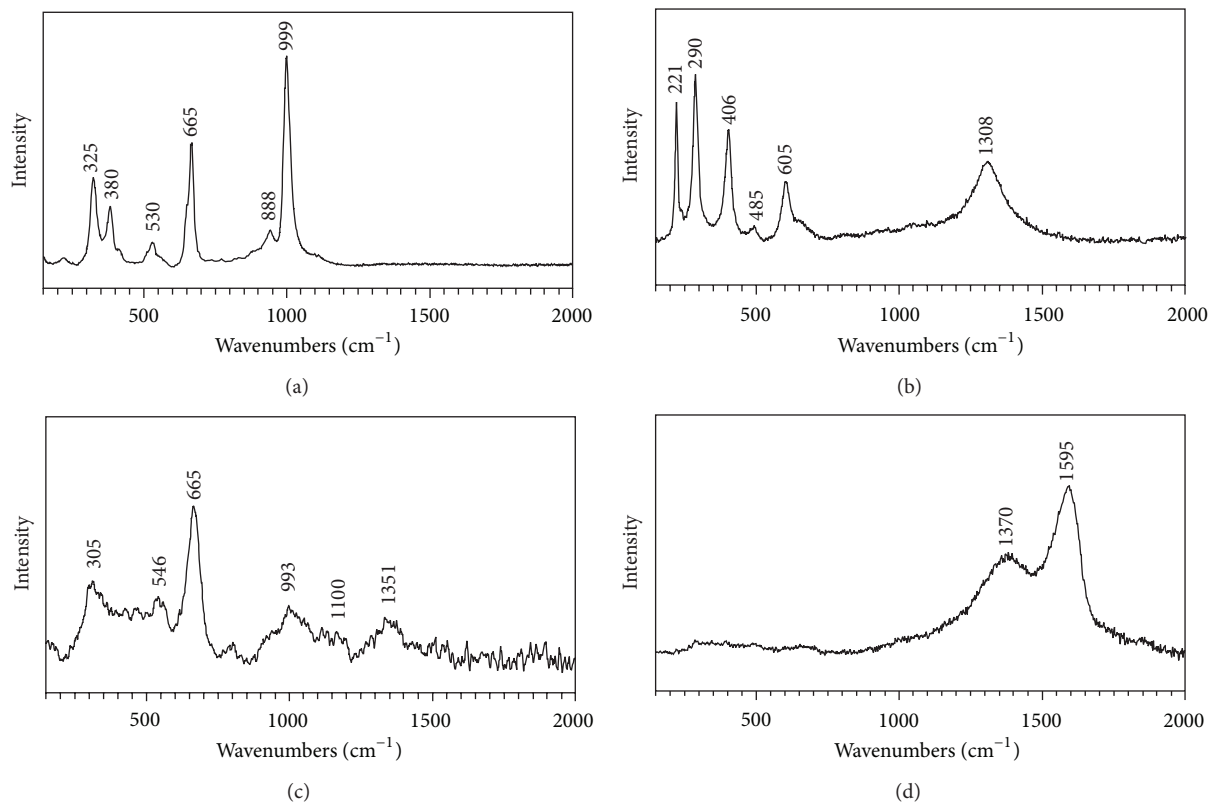


FIGURE 5: Micro-Raman spectra registered in different spots (see text).

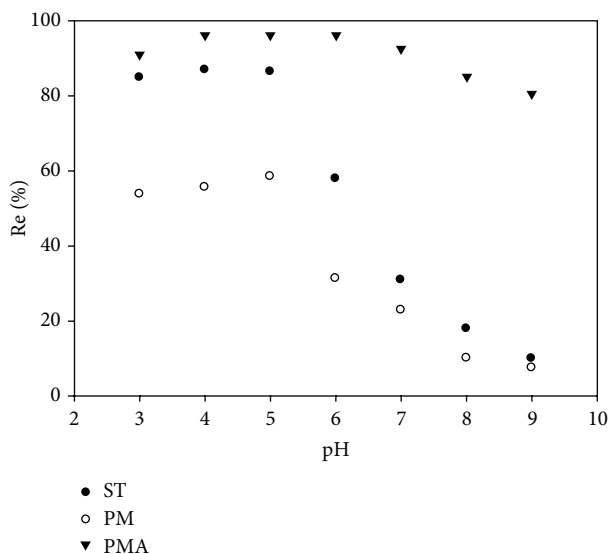


FIGURE 6: Arsenic removal by using volcanic pyroclastic mixture (PM), scoria-type (ST) material, and activated pyroclastic mixture (PMA), at different pH.

the presence of clinopyroxene and plagioclase. The first species is characterized by three strong Raman signals at 322, 667, and 997 cm^{-1} , whereas plagioclase (e.g., anorthite)

presents typical strong signal at 510 cm^{-1} and weak signals in the 1000 cm^{-1} region [18]. Hematite in natural samples or in any iron oxide mixtures is reliably identified by lines at ~ 225 , 291, 411, 611, and 1321 cm^{-1} , whereas magnetite shows two weak bands at ~ 300 and ~ 530 cm^{-1} and an intense signal at ~ 665 cm^{-1} which is the most conspicuous and diagnostically useful [17–19].

Spectrum of Figure 5(b) shows bands at 221, 290, 406, 605, and 1308 cm^{-1} attributed with hematite phase, whereas a shoulder at 662 cm^{-1} could be assigned to magnetite.

Likewise, the spectrum of Figure 5(c) is more complex. However, the magnetite presence can be corroborated from the Raman line at 665 cm^{-1} . Other lines at low frequencies (in the range 305–546 cm^{-1} , surely associated with the partial oxidation of magnetite) and between 990 and 1400 cm^{-1} (993, 1100, 1351 cm^{-1}) usually observed in some plagioclases [18] reveal the strong interaction between phases with different composition and crystal structure. Differences between spectra of Figures 5(b) and 5(c) are surely related to the alteration during the evolution of the eruption process.

Spectrum of Figure 5(d) is characterized by two intense signals centered at 1370 and 1595 cm^{-1} , which can be assigned to active modes of carbon particles (ordered and disordered graphite species, resp.) [20–22]. These seem to be associated to the natural chaoite, mineral species formed by heating graphite at high pressure associated with zircon, rutile, pseudobrookite, and magnetite, among other species.

3.2. *Potentiality in the As Sorption Process.* Although the volcanic ash constitutes a significant environmental hazard [23], investigations have been carried out in order to analyze its potential use as raw material in the preparation of adsorbents, cement, and so forth. In this context, the use of ferric impregnated volcanic ash has been proved in the arsenate adsorption process [12, 24].

The As content in the PM, PT, and ST studied materials (12.3, 15.7, and 12.4 ppm, resp.) is comparable to that observed in the loess of Argentinean aquifers [13]. However, it was observed that the As was not leached, according to the EPA international method [25]. These aspects were a good signal to prove the capability of the raw material without chemical modification to retain the toxic. On this basis, tentative assays for As removal were done.

The literature reports of adsorption data for ferric-impregnated volcanic ash showed that an increase of ~5% as Fe₂O₃ (by treatment of ferric solution (FeCl₃ 20 g L⁻¹)) greatly elevates the As(V) removal ability by the formation of iron (hydro)oxides [12]. This modification also affects the range of adsorption capability, through an electrostatic attraction as well as surface complexation processes between the As species (H₂AsO₄⁻, HAsO₄⁻²) and iron oxide surface sites [16, 26]. In our work conditions, the increase of iron content in the activated sample (PMA) can be correlated with a great increase of the surface area (73.69 m² g⁻¹).

Figure 6 gives the average values of As adsorbed in the PM, ST, and PMA materials (5.58, 8.92, and 10.40% Fe₂O₃, resp.), at the conditions indicated in the experimental section versus pH. The comparison of the results reveals that the As retention increases with acidity. Maximums of 58.5%, 87%, and 96.27% for PM, ST, and PMA were observed at pH 5, 4, and in the 4–6 range, respectively. These acidic conditions can be correlated with those reported in the literature, working with volcanic ash chemically modified with iron [12]. The pH effect can be attributed to the increase of adsorption sites by surface protons (formation of Fe-O-H and eventually Al-O-H previous to the dealumination process), facilitating the interaction with the (H₂AsO₄⁻ or HAsO₄⁻²) predominant species. This behavior at the surface level shows that arsenate has a direct chemical bond with iron species, with a specific sorption mechanism, following a surface complexation model (inner sphere complexes), as it is widely reported [16].

4. Conclusions

The components of the volcanic ash emitted by PCCVC eruption were identified by means of spectroscopic and microscopic techniques, which revealed a mixture of alkali silicates (predominantly glass) and microcrystals of iron oxides (hematite, magnetite), phases of plagioclase type, pyroxene type, and carbonous material. The pyroclastic- and the scoria-type materials, with bulk iron contents of 5.58 and 8.92% Fe₂O₃, have not shown arsenic leaching, retaining the arsenic in solution. On the other hand, the “in situ” chemical modification with a small proportion of iron oxide (~5% as Fe₂O₃) led to a useful adsorbent for arsenic removal from

aqueous solution, increasing also the effective range of pH. Comparatively, the good performance can be attributed to the increase of the adsorption sites through the formation of Fe-O-H groups. On the basis of physicochemical characterization, the studied ash seems to be an interesting raw material for the arsenic removal, transforming the volcanic waste in a suitable, inexpensive, and abundant adsorbent.

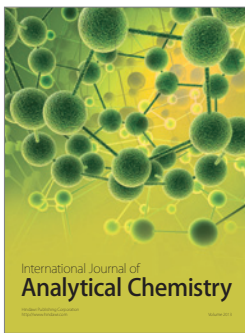
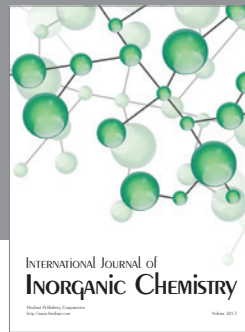
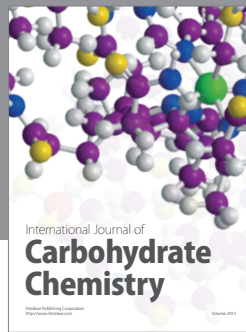
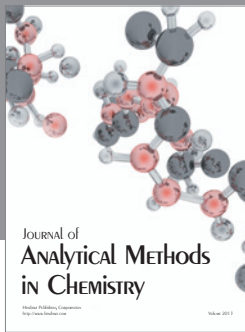
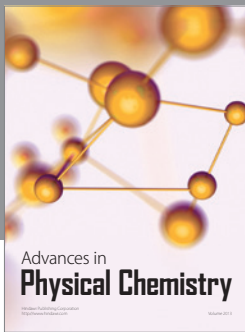
Acknowledgments

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