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78	Abstract	the world, but little subtropical region occurred in central present day soil of and woodlands, co measured past ch δ^{13} C values in so 14 C activity in so Present day relati δ^{13} C signatures in 0.78; <i>p</i> <0.01). De continuous replac this period, the re across sites) with ecosy stems. Dur central Argentina. grasslands, shrub changes in vegets changes in climat a shift in the C4 s	ogeographic changes are well understood in many parts of e is known about long-term v egetation dynamics in s. Here we investigate shifts in C ₃ /C ₄ plants abundance al Argentina during the past few millennia. We determined organic matter δ^{13} C signatures of grasslands, shrublands ontaining different mixtures of C ₃ and C ₄ plants. We hanges in the relative cover of C ₃ /C ₄ plants by comparing bil profiles with present day δ^{13} C signatures. We analyzed il depths that showed major changes in v egetation. ive cover of C ₃ /C ₄ plants determines whole ecosystems' integrated as litter and superficial soil organic matter (R ² = eeper soils show a consistent shift in δ^{13} C, indicating a ement of C ₄ by C ₃ plants since 3,870 (±210) YBP. During lative abundance of C ₃ plants increased 32% (av erage significant changes being observed in all studied ing the late Holocene C ₃ plants became dominant in We identified increases in the relative C ₃ /C ₄ cover in blands and woodlands, suggesting a physiological basis for ation. The replacement of C ₄ by C ₃ plants coincided with e towards colder and wetter conditions and could represent species optimum range.
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REGULAR ARTICLE

Evidence of shift in C₄ species range in central Argentina during the late Holocene

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7 Natalia Pérez-Harguindeguy · Valeria Falczuk · Larry L. Tieszen ·

8 Marcelo Cabido

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Abstract Millennial-scale biogeographic changes are 12well understood in many parts of the world, but little 13is known about long-term vegetation dynamics in 14subtropical regions. Here we investigate shifts in 15C₃/C₄ plants abundance occurred in central Argentina 16during the past few millennia. We determined present 17day soil organic matter δ^{13} C signatures of grasslands, 18shrublands and woodlands, containing different mix-19tures of C₃ and C₄ plants. We measured past changes 20in the relative cover of C_3/C_4 plants by comparing 21

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Present Address: L. C. R. Silva Land Air and Water Resources Department, University of California, Davis, CA 95616, USA δ^{13} C values in soil profiles with present day δ^{13} C 22signatures. We analyzed ¹⁴C activity in soil depths 23that showed major changes in vegetation. Present day 24relative cover of C_3/C_4 plants determines whole 25ecosystems' $\delta^{13}C$ signatures integrated as litter and 26superficial soil organic matter ($R^2=0.78$; p<0.01). 27Deeper soils show a consistent shift in δ^{13} C, 28indicating a continuous replacement of C₄ by C₃ 29plants since 3,870 (±210) YBP. During this period, the 30 relative abundance of C₃ plants increased 32% 31(average across sites) with significant changes being 32observed in all studied ecosystems. During the late 33 Holocene C₃ plants became dominant in central 34Argentina. We identified increases in the relative C_3/C_4 35cover in grasslands, shrublands and woodlands, sug-36 gesting a physiological basis for changes in vegetation. 37The replacement of C₄ by C₃ plants coincided with 38 changes in climate towards colder and wetter con-39ditions and could represent a shift in the C₄ species 40optimum range. 41

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Introduction

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Paleovegetation reconstructions have shown that45millennial-scale climate variability modulates the46distribution of ecosystems, affecting global terrestrial47biogeography. It is now clear from the palynological48

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record that synchronic changes in plant species range 49occurred in the recent geological past as a response to 5051climate change (Allen et al. 2010; Gajewski 2008; Williams et al. 2004). The vast majority of palyno-52logical studies, however, have been conducted in cold 5354(mostly northern) regions and little is known about climatically driven vegetation dynamics in tropical 55and subtropical ecosystems. In South America, for 5657example, palynological studies have been concentrated in cold zones of high altitudes, as unstable sedimenta-58tion, fast decay rates and exceedingly complex flora 5960 hinder the characterization of fossil pollen in warmer regions (Birks and Birks 2000; Flenley 1985; Marchant 61et al. 2002; Thouret et al. 1997). In areas where 6263 palynological investigation is possible, extrapolations from single isolated pollen profiles are typically used 64to infer regional shifts in vegetation (Behling and Pillar 652007; Ledru et al. 1996; Mancini 2009) and, for this 66 reason, we still lack a detailed understanding of how 67 past climate-vegetation interactions occurred. 68

69 Recent studies have successfully identified past vegetation changes in tropical and subtropical regions 70by analyzing soil organic carbon isotope ratios 7172(Dümig et al. 2008; Silva et al. 2008, 2010a). 73 Although more limited than palynological records for the purpose of describing changes in populations 7475of individual species or genera, carbon isotopes can 76be used to trace paleo events at the ecosystem level. This technique is particularly useful to describe local-7778 scale ecotonal shifts in predominant vegetation, where C₃- and C₄-dominated systems coexist. Generally, 79when light is not a limiting factor, plants with C₄ 80 81 metabolism dominate warm environments, while C₃ 82 plants predominate in cool climates (Sage 2004). The corollary is that C₄ outcompete C₃ plants in the 83 tropics and subtropics, while the converse is observed 84 85 in colder regions (Bond 2008). Fluctuations in climate, however, are expected to directly influence 86 the balance between C3 and C4 plants, imposing 87 88 alternate dominance of either metabolic pathway, explaining their co-occurrence through various scales 89 90 of space and time.

91 When contrasting metabolic pathways also repre-92 sent contrasting life forms (e.g. trees *versus* grasses), 93 indirect effects of changes in climate may play a 94 fundamental role on vegetation dynamics. For exam-95 ple, frequent and intense fires expected during 96 warmer and drier periods would favor C_4 grasses at 97 the expense of C_3 woody plants in local scales (Behling et al. 2004; Hoffmann et al. 2003). On the 98other hand, high atmospheric CO₂ levels globally, 99could favor C₃ species because of their lower 100 photosynthetic efficiency when compared with C_4 101plants (Epstein et al. 1997; Luo et al. 2006; Sage et al. 1022010). Recent studies have reported the expansion of 103C₃- (forests) over C₄-dominated ecosystems (savannas 104and grasslands) during the late Holocene (Behling et 105al. 2005, 2004; Dümig et al. 2008; Silva et al. 2008). 106Unfortunately, these studies were conducted where 107 distinct metabolisms represent differences in life form 108and for this reason they were not able to disentangle 109direct and indirect effects of climate fluctuations on 110 vegetation change. 111

Here we further investigate this issue, asking 112whether Holocene climate change has had a signifi-113cant direct effect on terrestrial ecosystems, promoting 114an overall competitive advantage of C3 over C4 115plants, independently of differences in species life 116form. To answer this question we sampled a broad 117region in central Argentina, which corresponds to 118 the present day southernmost limit of C₄ grasses 119distribution within the Americas (Ehleringer et al. 1202005). We investigated ecosystems of different 121structures (grassland, shrubland and woodland) 122containing various proportions of C₃ and C₄ plants. 123We quantitatively described shifts on C_3/C_4 vegeta-124tion cover occurred in the past using isotopic 125analysis of organic carbon and discussed the results 126in relation to previously reported climate and 127vegetation reconstructions. 128

Material and methods 129

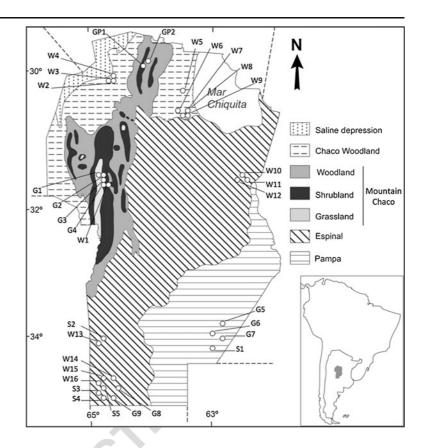
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Study region

The present study was conducted in central Argentina, 131within the limits of Córdoba province (Fig. 1). Both 132 C_4 (grasses) and C_3 (herbaceous and/or woody) plants 133coexist in the five major regional phytogeographical 134units: Pampa, Espinal, Chaco, Monte and Pastizal de 135altura (Cabrera 1976). These distinct vegetation types 136represent a gradient that encompasses grasslands to 137 the eastern lowlands (~100 m ASL), woodlands 138towards west at intermediate altitudes (300-500 m), 139xerophytic open shrublands in the semi-arid western 140 highlands (800–1,200 m) and grasslands in higher 141 altitudes (above the tree line, 1,600 m) (Fig. 1). 142

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Fig. 1 The vegetation of Cordoba as represented by the most important phytogeographical provinces. Circles show the approximated location of the study sites, where W; S and G and GP represent woodland; shrubland, grassland, and grassland with palms respectively



According to the Köppen-Geiger classification, the 143regional climate is temperate/warm-temperate to 144 subtropical, in average ranging from 10°C to 24°C 145between winter and summer (summer-time highs ~30° 146C and winter lows ~4°C), with a marked rainfall 147gradient from east (annual rainfall >1,000 mm) to 148west (annual rainfall <400 mm), with 70% of the 149rainfall occurring from November to March (Cabido 150et al. 2008). 151

152 Vegetation survey

153The regional landscape has been severely altered by human activity throughout the past centuries, but 154relicts of undisturbed vegetation can still be found. 155Using our own unpublished and other authors' 156published recent floristic surveys (Cabido et al. 1571997, 1993, 2008; Diaz et al. 1994, 1999, 2001; 158Perez-Harguindeguy et al. 2000; Pucheta et al. 1998; 159Zak and Cabido 2002), we selected 32 well-preserved 160161sites where the vegetation is representative of the most important regional ecosystems. We classified the 162distinct vegetation types in four major categories: 163

woodlands (16 sites); shrublands (5 sites); grasslands 164(9 sites) and grasslands with palms (2 sites) (W, S, G 165and GP sites indicated in Fig. 3). The number of sites 166 representing each of these vegetation categories 167 varied according to their floristic complexity. Wood-168lands comprise both xerophytic and mesophytic 169ecosystems with high diversity of woody and non-170woody plants, dominated by Aspidosperma and 171species of Prosopis, while shrublands only represent 172xerophytic ecosystems dominated by species of 173Geoffroea, Condalia or Maytenus genera. Grasslands 174encompass both C3- and C4-dominated treeless 175vegetation. In a couple of grassland sites the presence 176of palms of the species Trithrinax campestris was 177remarkable and for this reason these sites were 178classified as grassland with palms. 179

In previous studies we used leaf area index (LAI) 180 measurements to characterize C_3/C_4 relative cover 181 across vegetation gradients (see for example: Silva et al. 2008, 2010a, b, c). These studies were conducted, 183 however, in regions where different strata represent 184 distinct metabolic pathways (e.g. herbaceous layer is 185 dominated by C_4 grasses, while shrubs and trees are 186

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 C_3). In such conditions, LAI measurements (which do 187 not differentiate between species) are sufficient to 188describe changes in the relative cover of C₃ and C₄ 189plants. On the contrary, in the region studied here 190perennial C₃ and C₄ herbaceous plants can be equally 191192abundant (Zuloaga et al. 2008) and predominant metabolic pathways can only be assessed by detailed 193floristic surveys. For this reason, we performed a 194 195complete census, including all herbaceous, shrub and tree species at each studied site. Our survey was 196conducted during the growing season (Summer) of 1972010, when maximum productivity is typically 198attained and all potential species are present. The 199relative contribution of each species for the total 200vegetation cover was estimated (visual estimation-201 projection of canopy cover in 400 m² plots-3 plots 202per site) and from these estimates the relative cover of 203 C_3 and C_4 species was calculated (Cabido et al. 1997, 2042008; Renison et al. 2006). Plants that have the C₄ 205photosynthetic pathway were distinguished from 206those that have the C₃ pathway by examination of 207 208the Kranz anatomy in cross-sections of fresh and herbarium specimens and from the literature (Sánchez 209 210and Arriaga 1990; Smith and Epstein 1971). A list of 211the surveyed species, including growth habit and metabolic pathways, is presented in Appendix 1. 212More details about the regional vascular flora can be 213214found at Zuloaga et al. (2008), but see also their on line updates at: www.darwin.edu.ar. 215

216 Isotopic signatures and past vegetation changes

To test whether and how much C₃ and C₄ relative 217vegetation cover has changed we relied on two well-218known natural processes: (i) during photosynthesis C₃ 219plants discriminate more against CO2 molecules that 220 contain the stable isotope ¹³C than do C₄ plants. For 221222this reason different proportions of C_3/C_4 contribution to the total biomass yield differences in the mean 223carbon isotopic ratios (δ^{13} C) in the soil organic 224225matter (SOM) (Ehleringer et al. 2000; Marshall et al. 2007; Silva et al. 2010a, b, 2008; Smith and Epstein 2261971; Von Fischer et al. 2008); (ii) SOM in soil 227profiles represents a chronological sequence of 228vegetation signature, with past vegetation recorded 229at deeper levels (Dümig et al. 2008; Nordt et al. 2302312007; Sanaiotti et al. 2002; Silva et al. 2010a, 2008; Victoria et al. 1995). Based on these we character-232ized past shifts in vegetation structure by, firstly, 233

determining the current vegetation cover at each 234 study site (described in the previous section), 235 secondly, by determining the carbon isotopic signatures (δ^{13} C) in the litter and superficial SOM and, 237 finally, by examining the δ^{13} C of SOM at different 238 depths in soil profiles. 239

We used 3 to 5 soil profiles at each site to sample: 240fresh litter (any decaying organic matter found at the 241uppermost soil layer), superficial (0-1 cm depth) and 242deep soil (10, 20, 30, 50 and 100 cm depth). Prior to 243analysis litter was dried for two days at 70°C and 244coarsely ground, while superficial and deep soils were 245dried at room temperature after which fine roots were 246removed by sieving through a 0.8 mm mesh and not 247analyzed. The carbon isotope ratios of litter and soil 248samples were determined at the Laboratory of Stable 249Isotope Ecology (LSIETE) at the University of Miami. 250Samples (10 mg) were loaded in tin cups, which were 251placed in an automated elemental analyzer connected to 252a continuous flow isotope ratio mass spectrometer. ¹³C 253abundances were expressed as: 254

$$\delta^{13}C$$
 (‰) = (R_{sample}/RPDB-1) × 1000

where R_{sample} and RPDB represent the ${}^{13}C/{}^{12}C$ ratios 256 and PeeDee standard respectively. The precision of 257 analysis was $\pm 0.1\%$ (\pm SD). 258

We used regression analysis to describe the effect 259 of present day C_3/C_4 relative vegetation cover on litter 260 and superficial SOM δ^{13} C values (signatures). Based 261 on the equation that best described this relationship 262 (Fig. 2) we estimated past C_3/C_4 relative cover, by 263

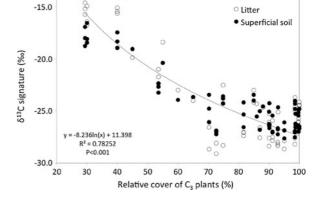


Fig. 2 Relationship between the relative vegetation cover of C_3 plants and the isotopic composition (δ^{13} C) of litter and superficial soil organic matter (0–1 cm depth), across all study sites

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examining SOM δ^{13} C signatures at deeper SOM in 264soil profiles. We estimated past vegetation cover for 265each study site and soil depth (1, 10, 20, 30, 50 266and 100 cm). By calculating percent differences 267between the C_3/C_4 relative cover estimated in 268 superficial (0-1 cm) and in deep soils, where major 269 shifts in vegetation were observed (50 cm), we 270determined the total vegetation change. We also 271measured the slope (angular coefficient) of least 272273square regressions, performed between estimates of C_3/C_4 cover and their respective soil depth at each 274study site. All soil depths were used for this 275276calculation, here named degree of vegetation change. We used two-way analysis of variance (ANOVA) to 277test the effect of ecosystem, altitude and their 278279interaction on both metrics of vegetation change (total percent and degree of vegetation change) 280 across sites. Analysis of variance and regression 281282 analysis were performed according to standard methods, using the software JMP Version 9 (JMP 2832009). For brevity, only statistically significant 284285results are presented.

286 Carbon dating

We analyzed ¹⁴C activity of carbon in the SOM of the 287soil profiles that showed major shifts in C_3/C_4 288vegetation cover at each ecosystem (grassland, shrub-289land and woodland: G4, G8, GP1, S1, W7 and W13; 290all at 50 cm depth; Fig. 1). Because roots from the 291modern day vegetation can exude recent ¹⁴C into the 292 293 deeper soil matrix, dates acquired by analyzing SOM ¹⁴C may be underestimated. For this reason, SOM ¹⁴C 294measurements should be considered as a proxy for 295determining the minimum possible date of past 296vegetation changes (Dümig et al. 2008; Martinelli et 297 al. 1996; Silva et al. 2010a, 2008; Trumbore 2000; 298Victoria et al. 1995). Analyses of ¹⁴C activity were 299 performed using accelerator mass spectrometry at the 300 Beta Analytic Radiocarbon Dating Laboratory (Miami, 301302USA). The results were then calibrated to represent actual calendar years of the mean age of SOM. The 303calibration was performed using the newest (2004) 304 305calibration database (Reimer et al. 2004). Calibrated dates expressed as years before present (YBP) were 306consistent among sites/ecosystems and here we 307 308 present the average date and standard deviation of measurements that represent major vegetation shifts 309 310across sites.

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Results

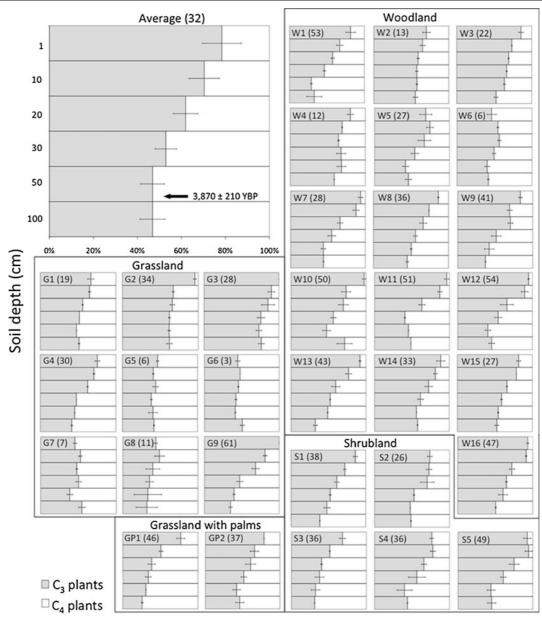
Current vegetation cover and δ^{13} C signatures

Differences in the present day relative contribution of 313 C₃ and C₄ plants to total vegetation cover explain 314 variations in whole ecosystems' $\delta^{13}C$ signatures, 315integrated as litter and superficial soil organic matter 316 $(R^2=0.78; p<0.01)$ (Fig. 2). Similar relationships 317have been found in previous studies conducted in 318different ecosystems (Lloyd et al. 2008; Silva et al. 3192010a, 2008) and represent an expected gradient 320 determined by variations in the amount of C₃- and 321 C₄-originated biomass and their mixture. Due to 322 intrinsic differences in discrimination, C₃ plants 323 δ^{13} C signature (ranging from -20 to -35%) are 324 unmistakably distinguishable from C₄ plants δ^{13} C 325signature (ranging from -9 to -15%) (Ehleringer et 32**Q1** al. 1993, 1997). The δ^{13} C values found here ranged 327from -15‰ in ecosystems dominated by C₄ grasses 328 to -29% in ecosystems dominated by C₃ plants 329 (either herbaceous or woody plants). This supports 330 our estimates of vegetation cover as being represen-331tative of the total in-situ biomass production, attrib-332utable to either C3 or C4 metabolic pathways. Both 333 litter and superficial SOM δ^{13} C significantly reflected 334 the current vegetation cover and, thus, could be used 335as reference to trace past changes in vegetation 336 recorded in the SOM at deeper soil layers. 337

Past vegetation change

Using the equation presented on Fig. 2 and isotopic 339measurements performed at different depths of soil 340profiles, we calculated changes in the relative C_3 and 341C₄ vegetation cover at each study site. Our results 342 show a consistent signal of continuous vegetation 343 change, with C₄ being replaced by C₃ plants. 344 Measurements of ¹⁴C activity show that this process 345 began approximately 3,870 (±210) years before 346 present and, since then, the relative abundance of C_3 347plants increased 32% in average (Fig. 3). All wood-348lands and shrublands, currently dominated by C₃ 349 plants (> than 70% of the total vegetation cover), had 350 a greater C_4 cover in the past. In some cases C_4 351grasses were the predominant vegetation, while C₃ 352plants represented less than 20% of the total vegeta-353tion cover (see sites W1, W9, W13; Fig. 3). The least 354noticeable shift observed in woodlands corresponds to 355

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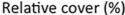


Fig. 3 Changes in the relative cover of C_3 and C_4 plants, estimated from $\delta^{13}C$ values measured in the soil organic matter (SOM) at different depths of soil profiles. The conversion of $\delta^{13}C$ values into percent cover followed the relationship found between present day vegetation cover and $\delta^{13}C$ signatures measured in the litter and superficial SOM (Fig. 2). Grey and

a 6% change towards a greater C_3 cover in the present (W6; Fig. 3). In other woodland sites C_3 cover increased at least 12%, often reaching values greater than 40%. Similarly, the replacement of C_4 by C_3

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white bars represent average values of C_3 and C_4 cover at each study site (3 to 5 profiles per site) and error bars correspond to one standard deviation. Numbers in brackets show total percent change in vegetation cover since 3,870 (±210) years before present. Site locations are shown on Fig. 1

plants in shrublands ranged from 26 to 49% in the 360 past few thousand years (Fig. 3). In grasslands, 361 vegetation cover remained nearly unaltered in a 362 couple of sites (see G6, G7; Fig. 3), but in most 363

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364cases we observed shifts in C_3/C_4 cover of the same365magnitude (in some cases higher) than those observed366in shrub and woodland sites (see for example G9=36761%; Fig. 3).

It is important to note that these results only 368 represent records of past vegetation at each given site. 369 However, the fact that expressive changes in vegeta-370 tion occurred in most sites and always in the same 371direction (C₃ replacing C₄ plants), suggests a regional 372scale shifts in vegetation. Two-way ANOVA models 373indicate that changes in vegetation, measured as 374percent differences between past and recent C3/C4 375376 cover, occurred independently of ecosystem (grassland, shrubland and woodland), altitude (sites ranged 377from 100 to 2,012 m asl) or their interaction. 378 379 Likewise, the degree of vegetation change across sites, measured as the angular coefficient of least 380 square regressions between C_3/C_4 cover and soil 381 depths, were not affected by ecosystem, altitude or 382their interaction (Appendix 2). 383

384 Discussion

 $385 \quad \delta^{13}C$ signatures and past vegetation change

We found a highly significant relationship between the 386relative cover of C_3 (herbaceous and woody) and C_4 387 (grasses) with δ^{13} C values measured in the organic 388matter across sites. It has been well documented that 389differences in the total contribution of C3- and C4-390originated biomass to whole ecosystems production 391 yield distinguishable carbon isotopic signatures (Dümig 392et al. 2008; Ehleringer et al. 2000; Lloyd et al. 2008; 393Sanaiotti et al. 2002; Silva et al. 2010a, b, 2008; Victoria 394 et al. 1995; Von Fischer et al. 2008). Similarly, here we 395 show that differences in vegetation cover, with respect 396397 to metabolic pathways, explain variations in soil organic matter δ^{13} C signatures. Litter carbon isotope ratios 398 varied with C_3/C_4 cover, remaining nearly unaltered 399after decomposition. The spectrum of $\delta^{13}C$ ratios 400 reported here, varying up to 15‰ depending on the 401vegetation cover, is consistent with those described in 402 different C3- and C4-dominated ecosystems (Lloyd et al 4032008, Silva et al 2008, Von Fischer et al. 2008) and 404 represent reliable signatures that can be used to trace 405 shifts in vegetation through time. 406

407Although processes other than changes in vegeta-408tion may alter isotopic ratios before and after litter

deposition, major paleoecological events can be 409 identified with good confidence in soil profiles, 410 because organic matter δ^{13} C ratios tend to be stable 411 long after decomposition, conserving vegetation sig-412 natures throughout millennia (Silva et al. 2008; Von 413 Fischer et al. 2008). Changes in the isotopic compo-414 sition of atmospheric CO₂ could have affected values 415of δ^{13} C prior to deposition, but with the exception of 416past century, δ^{13} C variations in atmospheric CO₂ have 417been small (<1.0% within the last 10,000 years; 418Flückiger et al. 2002). Variations in δ^{13} C due to 419changes in photosynthetic discrimination are also 420 possible, but would not have produced similar 421 changes across sites/ecosystems and even if so, such 422 variations would have been more than one order of 423 magnitude smaller than the range observed here 424 (Martin-Benito et al. 2010; Nock et al. 2010; Silva 425et al. 2010c, 2009). 426

Differences in root depth between trees and grasses 427 could have contributed to soil organic matter isotopic 428 enrichment with depth. However, the presumably 429 deeper rooting habit of trees could have depleted (C₃ 430signal), but not enriched (C₄ signal), SOM δ^{13} C. 431Moreover, differences in root depth between C₃ and 432 C_4 plants typically yield small changes in $\delta^{13}C$ and 433isotopic offsets that exceed 3‰ are usually interpreted 434 as a shift in vegetation (Krull et al. 2002; Von Fischer 435et al. 2008). Additionally, because changes in vegeta-436tion were observed in different ecosystems (grassland, 437 shrubland and woodland), differences in root depth 438cannot explain changes in $\delta^{13}C$ observed here. Post 439 depositional differential degradation of chemical com-440 pounds and fractionation associated with microbial 441 activity could also have affected soil organic matter 442isotopic ratios (Marshall et al. 2007). However, this 443 influence is typically observed only within the top 444 20 cm of the soil profile, correlated with total soil 445 organic carbon content (Ehleringer et al. 2000; Krull et 446 al. 2002). We observed changes in δ^{13} C values at much 447 deeper layers of the soil profile that were not correlated 448 with soil carbon content. Total soil carbon in the study 449 region range from 2 to 26 gkg^{-1} according to 450vegetation type and altitudinal/climatic variations 451(Alvarez and Lavado 1998), yet we observed consis-452 tent variations in δ^{13} C across sites, which did not 453change significantly in different ecosystems or alti-454tudes. For these reasons, we conclude that changes in 455 δ^{13} C values observed with soil depth in the present 456study were caused by past changes in vegetation. 457

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458 A possible northward shift in the C_4 species range

459We detected a convergent signal of soil organic matter isotopic signatures, showing a continuous increase in 460the relative abundance of C₃ plants since 3,870 461 462 (±210) years ago. Previous palynological and isotopic studies from southern South America have identified 463the expansion of C3-dominated (forests) over C4-464465 dominated ecosystems (savannas and grasslands) during the same period (Behling et al. 2005, 2004; 466 Dümig et al. 2008; Mancini 2009; Silva et al. 2008; 467Wille and Schabitz 2009). While these studies have 468 described local shifts in forest-savanna or forest-469grassland ecotones, the results presented here show 470that vegetation changed at a regional scale and 471independently of ecosystem structure. In the past, C₄ 472grasses greatly contributed to the total biomass of 473sites now dominated by distinct groups of C₃ plants. 474Because changes in the predominant metabolic 475pathway occurred in woodlands, shrublands and 476477 grasslands, they were probably not related to intrinsic differences in plant life forms but rather had a 478physiological basis. 479

480 Changes in ecosystem structure, from dense to open vegetation, can be site-specific and determined 481by local environmental gradient or disturbance history 482 (Crisci et al. 1991; Da Silva and Bates 2002). For 483 example, it has been shown that spatial and/or 484temporal changes in vegetation structure may be 485attributed to variations in edaphic properties (Haridasan 486 2008), water availability (Bush et al. 2004; Furley et al. 4871992; Silva et al. 2009), fire regime (Accatino et al. 4882010; Hoffmann et al. 2003; Moreira 2000) and 489 grazing intensity (Diaz et al. 2001). We suspect that 490all these factors have contributed to configure the 491 492current structure of the vegetation at each studied site. However, a simultaneous shift in the predominant 493metabolic pathway across sites, suggests that a major 494495force has promoted changes in plant communities during the late Holocene. 496

497 We believe that changes in climate were responsible for the shift in vegetation reported here. World-498 wide, the extant C₄ genera occupy a wider range of 499drier and warmer habitats than their C3 counterparts 500501because the C_4 pathway represents a pre-adaptation to hot and arid conditions (Osborne and Freckleton 5022009). Reconstructions of paleoclimate have shown 503504a significant drop in global temperatures during the past few thousand years (Fig. 4). This climate cooling 505



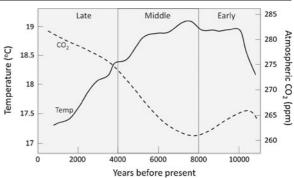


Fig. 4 Temperature and CO_2 records for the Holocene time period. Temperature record according to alkenone measurements (core GeoB 7139–2) of north-central Chile (De Pol-Holz et al. 2006) and CO_2 record described from ice core analyses (EPICA ice cores; Flückiger et al. 2002)

followed the retreat of glaciers in the northern 506hemisphere during the Holocene climatic optimum, 507which also led to substantial water input in tropical 508and subtropical systems (Markgraf et al. 1992; 509Vimeux et al. 2009; Wanner et al. 2008). Fossil 510pollen, charcoal and isotopic evidence from different 511South American biomes indicate that C₃ plants 512became more abundant during the late Holocene, 513coinciding with changes in climate towards colder 514and wetter conditions (Behling et al. 2005, 2004, 515Mancini 2009, Silva et al. 2008). 516

Presently, the composition of C_3 and C_4 species in 517 the study region follow temperature parameters more 518 strongly than rainfall (Cavagnaro 1988, Cabido 51Q21997). Along altitudinal gradients C_4 appear in 520 greater number than C_3 species in locations where 521

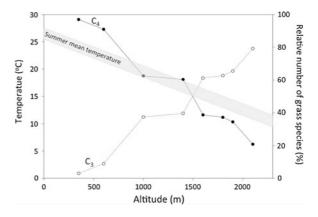


Fig. 5 Changes in the present day relative species composition of C_3 and C_4 grasses across an altitudinal gradient in central Argentina, as defined by summer (January) average temperatures (adapted from Cabido et al. 1997)

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summers are warmer than 18°C, while in colder areas near future (Sage 2004; Sage et al. 2010; Wand et al. the number of C₃ species represent more than 60% of 1999). This could represent a return to C₄-dominated the total grass diversity (Fig. 5). Although the number systems in the latitude studied here, as observed of species is not necessarily linked to the total cover during warmer periods. It is important to note of either C₃ or C₄ plants, it seems fair to assume that however that the velocity of climate change during the Holocene allowed time for species to migrate in colder temperatures (average annual temperatures dropped from >18.5°C to <17.5°C; Fig. 4) increased locked step with their optimal range and the the competitive advantage and, consequently, the consequences of the much faster rates of recent relative abundance of C3 plants at the latitude studied here. Increases in atmospheric CO₂ concentrations could also have favored C3 at the expenses of C4 plants (Epstein et al. 1997; Huang et al. 2007; Luo et al. 2006; Sage et al. 2010). Atmospheric CO₂ levels have increased monotonically (~20 ppm in total) since the mid Holocene (Fig. 4). The reasons are debatable, but decreases in total terrestrial biomass in response to colder temperatures (Indermühle et al. 1999) and a drop in oceanic carbonate ion concentration (Broecker et al. 2001) are possible causes behind this natural CO₂ enrichment. However, even though C₃ plants are worldwide. more responsive to elevations in CO₂ than C₄ plants, enrichments of this magnitude should not have altered the competitive balance between these two metabolic pathways (Wand et al. 1999). Therefore, changes in temperature are the most likely explanation for a regional scale vegetation displacement.

Despite increasingly higher atmospheric CO₂ 565566levels, warmer global temperatures are expected to promote the expansion of C₄ species range in the 567

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Appendix 1 602

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t1.1Table 1 Identity, life form and metabolic pathway of all plant species present in the studied sites. Plants with C4 photosynthetic pathway were distinguished from those that have the C₃ pathway by examination of the Kranz anatomy in crosssections of fresh and herbarium specimens and from the literature

Family/species	Life form	Metabolic pathway
Acanthaceae		
Dicliptera squarrosa Nees	herb	C3
Justicia squarrosa Griseb.	herb	C3
Stenandrium dulce (Cav.) Nees	herb	C3
Amaranthaceae		
Amaranthus hybridus L. ssp. hybridus	herb	C3
Guilleminea densa (Willd. ex Roem. & Schult.) Moq.	herb	C3
Iresine diffusa Humb. & Bonpl. ex Willd. var. diffusa	herb	C3
Pfaffia gnaphaliodes (L. f.) Mart.	herb	C3

global warming on plant communities remain un-576known. There is still great uncertainty in predicting 577 the future of C3- and C4-dominated ecosystems 578globally (Bond 2008). So far attempts to predict 579climate change impacts on terrestrial ecosystems 580have been limited to correlations between the current 581climate and species distribution (Pearson and Dawson 5822003). The search for long-term convergent patterns as 583those described here, could lead to an improved picture 584of climate-vegetation interactions, increasing our abil-585ity to predict the future impacts of climate change 586

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Family/species	Life form	Metabolic pathwa
Amaryllidaceae		
Zephyranthes longistyla Pax	herb	C3
Anacardiaceae		
Lithraea molleoides (Vell.) Engl.	tree	C3
Schinus bumeloides I.M. Johnst.	shrub	C3
Schinus fasciculatus (Griseb.) I. M. Johnst. var. fasciculatus	shrub	C3
Apiaceae		
Ammi visnaga (L.) Lam.	herb	C3
Bowlesia incana Ruiz & Pav.	herb	C3
Conium maculatum L.	herb	C3
Eryngium agavifolium Griseb.	herb	C3
Eryngium horridum Malme	herb	C3
Eryngium nudicaule Lam.	herb	C3
Apocynaceae		
Amblyopetalum coccineum (Griseb.) Malme	herb	C3
Aspidosperma quebracho-blanco Schltdl.	tree	C3
Ditassa buchellii Hook. & Arn.	epiphyte	C3
Metastelma tubatum Griseb.	epiphyte	C3
Morreria odorata (Hook. & Arn.) Lindl.	epiphyte	C3
Arecaceae	epipityte	05
	nolm	C3
Trithinax campestris (Burmeist.) Drude & Griseb.	palm	03
Asteraceae	11	<u></u>
Ambrosia tenuifolia Spreng.	herb	C3
Baccharis coridifolia DC.	herb	C3
Baccharis pingraea DC.	herb	C3
Baccharis salicifolia (Ruiz & Pav.) Pers.	herb	C3
Baccharis stenophylla Ariza	herb	C3
Baccharis ulicina Hook. & Arn.	herb	C3
Bidens andicola var. decomposita Kuntze	herb	C3
Bidens pilosa L. var. pilosa	herb	C3
Bidens subalternans DC.	herb	C3
Carduus acanthoides L.	herb	C3
Carduus thoermeri Weinm.	herb	C3
Chaptalia nutans (L.) Pol.	herb	C3
Cirsium vulgare (Savi) Ten.	herb	C3
Conyza bonaeriensis (L.) Cronquist	herb	C3
Conyza primulifolia (Lam.) Cuatrec. & Lourteig	herb	C3
Eupatorium inulifolium Kunth	herb	C3
Eupatorium patens D. Don ex Hook. & Arn.	herb	C3
Eupatorium viscidum Hook. & Arn.	herb	C3
Galinsoga parviflora Cav.	herb	C3
Gamochaeta falcata (Lam.) Cabrera	herb	C3
Gamochaeta sp.	herb	C3
Gaaphalium gaudichardianum DC.	herb	C3
Heterosperma ovatifolium Cav.	herb	C3

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Plant Soil

t1.58

Table 1 (continued)

Family/species	Life form	Metabolic pathway
Heterotheca subaxillaris (Lam.) Britton & Rusby	herb	C3
Hypochaeris caespitosa Cabrera	herb	C3
Lactuca serriola L.	herb	C3
Schkuria pinnata (Lam.) Kuntze ex Thell.	herb	C3
Senecio ceratophylloides Griseb.	herb	C3
Senecio pampeanus Cabrera	herb	C3
Sonchus oleraceus L.	herb	C3
Synedrellopsis grisebachii Hieron. & Kuntze	herb	C3
Tagetes minuta L.	herb	C3
Taraxacum officinale G. Weber ex F. H. Wigg.	herb	C3
Verbesina encelioides (Cav.) Benth. & Hook. f. ex A. Gray	herb	C3
Vernonia incana Less.	herb	C3
Vernonia nudiflora Less.	herb	C3
Xanthium spinosum L. var. spinosum	herb	C3
Basellaceae		
Anredera cordifolia (Ten.) Steenis	epiphyte	C3
Bignoniaceae		
Amphilophium cynanchoides (DC.) L. G. Lohmann	epiphyte	C3
Dolichandra cynanchoides Cham.	epiphyte	C3
Brassicaceae	115	
Descurainia argentina O.E. Schulz	herb	C3
Exhalimolobos weddellii (E. Fourn.) Al-Shehbaz & C. D. Bailey	herb	C3
Lepidium bonariense L.	shrub	C3
Bromeliaceae		
Tillandsia aizoides Mez	epiphyte	C3
Tillandsia bryoides Griseb. ex Baker	epiphyte	C3
Tillandsia capillaris Ruiz & Pav.	epiphyte	C3
Tillandsia duratii Vis. var. duratii	epiphyte	C3
Tillandsia vinhioides Ker Gawl. var. xiphioides	epiphyte	C3
Calyceraceae	opipiijio	0.5
Boopis anthemoides Juss.	herb	C3
Campanulaceae	nero	05
Wahlenbergia linarioides (Lam.) A. DC.	herb	C3
Capparaceae	nero	05
Capparia caae Capparis atamisquea Kuntze	shrub	C3
Caryophyllaceae	siiruo	0.5
	an in lasta	C2
Cardionema ramosissima (Weinm.) A. Nelson & J. F. Macbr.	epiphyte	C3
Polycarpon tetraphyllum (L.) L.	herb	C3
Silene antirrhina L. var. antirrhina	herb	C3
Celastraceae		C 2
Moya spinosa Griseb.	shrub	C3
Celtidaceae		
Celtis ehrenbergiana (Klotzch) Liebm.	tree	C3
Chenopodiaceae		

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Family/s	pecies	Life form	Metabolic pathway
Allenrol	fea patagonica (Moq.) Kuntze	herb	C3
Atriplex	lampa (Moq.) D. Dietr.	herb	C3
Chenopo	odiaceae sp.	herb	C3
Chenopo	odium album L.	herb	C3
Salsola	kali L.	herb	C3
Sarcoco	rnia ambigua (Michx.) M.A. Alonso & M.B. Crespo	herb	C3
Comme	inaceae		
Comme	ina erecta L.	herb	C3
Convolv	ulaceae		
Cressa t	ruxillensis Kunth	herb	C3
Dichond	ra microcalyx (Hallier f.) Fabris	herb	C3
Evolvul	us sericeus Sw.	herb	C3
Ipomoea	hieronymi (Kuntze) O'Donell	epiphyte	C3
Ipomoea	nil (L.) Roth	epiphyte	C3
Ipomoea	sp.	epiphyte	C3
Cucurbi	taceae		
Sicyos r	nalvifolius Griseb.	epiphyte	C3
Cyperac	eae		
Bulbost	/lis juncoides (Vahl) Kük. ex Herter var. juncoides	sedge	C3
Carex so	proria Kunth ssp. sororia	sedge	C3
Cyperac	ee sp.	sedge	C3
	aggregatus (Willd.) Endl. var. aggregatus	sedge	C3
	entrerianus Boeck.	sedge	C3
	hermaphroditus (Jacq.) Standl.	sedge	C3
	incomtus Kunth	sedge	C3
Ephedra		-	
-	americana Humb. & Bonpl. ex Willd.	epiphyte	C3
-	triandra Tul. emend. J. H. Hunz.	epiphyte	C3
Euphorb			
-	a communis Müll. Arg.	herb	C3
	achnostachyus Baill.	herb	C3
	ia acerensis Boiss.	herb	C3
-	ia dentata Michx.	herb	C3
-	ia lorentzii Müll. Arg.	herb	C3
-	ia serpens Kunth	herb	C3
-	eraniifolia Klotzch ex Baill.	herb	C3
	ieronymii Pax & K. Hoffm.	herb	C3
Fabacea	-		00
	uroma Gillies ex Hook & Arn.	shrub	C3
	caven (Molina) Molina	shrub	C3
	gilliesii Steud.	shrub	C3
	praecox Griseb.	shrub	C3
-	a bicolor (Poir.) DC.	epiphyte	C3
	m praecox (Ruiz & Pav. ex Hook.) Harms glaucum Ssp. (Cav.) Burkart & Carter	tree	C3

Plant Soil

Table 1 (continued)

t1.148

Far	mily/species	Life form	Metabolic pathway
Co	logamia broussonetii (Balb.) DC.	epiphyte	C3
Co	ursetia hassleri Chodat	herb	C3
De	smanthus sp.	herb	C3
De	smodium uncinatum (Jacq.) DC.	herb	C3
Fał	baceae sp.	herb	C3
Ge	offraea decorticans (Gillies ex Hook. & Arn.) Burkart	tree	C3
Gle	editsia triacanthos L.	tree	C3
Me	edicago lupulina L.	herb	C3
Mi	mosa detinens Benth.	shrub	C3
Mi	mozyganthus carinatus (Griseb.) Burkart	shrub	C3
Pro	osopis alba Griseb.	tree	C3
Pro	osopis caldenia Burkart	tree	C3
Pro	osopis flexuosa DC.	tree	C3
Pro	osopis nigra (Griseb.) Hieron. var. nigra	tree	C3
Pro	osopis pugionata Burkart	tree	C3
Pro	osopis torquata (Cav. ex Lag.) DC.	tree	C3
Rh	ynchosia senna Griseb. ex Hook.	epiphyte	C3
	nna aphylla (Cav.) H. S. Irwin & Barneby	herb	C3
	raniaceae		
Ge	ranium dissectum L.	herb	C3
Jun	ncaceae		
Jun	neus acutus L.	sedge	C3
Jun	ncus imbricatus Laharpe	sedge	C3
	ncaceae ncus acutus L. ncus imbricatus Laharpe ncus sp miaceae	sedge	C3
	miaceae	e	
Cli	nopodium gilliesii (Benth.) Kuntze	shrub	C3
	ptis floribunda (Briq.) Briq. ex Micheli	herb	C3
	ptis mutabilis (Rich.) Briq.	herb	C3
	onurus japonicus Houtt.	herb	C3
	thraceae		
-	phea glutinosa Cham. & Schltdl.	herb	C3
	imia salicifolia (Kunth) Link	herb	C3
	lpighiaceae		
	rdobia argentea (Griseb.) Nied.	epiphyte	C3
	nusia guaranitica (A. StHil.) A. Juss.	epiphyte	C3
	comaria usillo Hook. & Arn.	shrub	C3
	lvaceae	511140	
	apovicasea flavescens	herb	C3
	alvortadeu naveseens alvastrum coromandelianum (L.) Garcke	herb	C3
	odiolastrum malvifolium (Griseb.) K. Schum.	herb	C3
	eudoabutilon pedunculatum (R. E. Fr.) Krapov.	herb	C3
	la dictyocarpa Griseb. ex K. Schum.	herb	C3
	la rhombifolia L.	herb	C3
		11010	05

t1.193	Table 1 (continued)		
	Family/species	Life form	Metabolic pathway
t1.194	Moraceae		
t1.195	Morus alba L.	tree	C3
t1.196	Olacaceae		
t1.197	Ximenia americana L.	shrub	C3
t1.198	Onagraceae		62
t1.199	Oenothera affinis Cambess.	herb	C3
t1.200 t1.201	Oenothera sp. Orchidaceae	herb	C3
t1.201 t1.202	Cyclopogon elatus (Sw.) Schltr.	herb	C3
t1.202 t1.203	Habenaria sp.	herb	C3
t1.203 t1.204	Oxalidaceae	liero	
t1.201	Oxalis conorrhiza Jacq.	herb	C3
t1.206	Passifloraceae		
t1.207	Passiflora mooreana Hook. f.	epiphyte	C3
t1.208	Phytolaccaceae		
t1.209	Petiveria alliacea L.	herb	C3
t1.210	Rivinia humilis L.	herb	C3
t1.211	Plantaginaceae		
t1.212	Plantago tomentosa Lam.	herb	C3
t1.213	Poaceae		
t1.214	Agrostis montevidensis Spreng. ex Nees f. montevidensis	grass	C3
t1.215	Aristida laevis (Nees) Kunth	grass	C4
t1.216	Aristida mendocina Phil.	grass	C4
t1.217	Bothriochloa barbinodis (Lag.) Herter	grass	C4
t1.218	Bothriochloa laguroides (DC.) Herter ssp. laguroides	grass	C4
t1.219 t1.220	Bothriochloa springfieldii (Gould) Parodi Bromus catharticus Vahl var. catharticus	grass	C4 C3
t1.220 t1.221	Cenchrus myosuroides Kunth var. myosuroides	grass	C3 C4
t1.221 t1.222	Cenchrus spinifex Cav.	grass grass	C4 C4
t1.223	Chascolytrum subaristatum (Lam.) Desv.	grass	C3
t1.224	Chloris halophila Parodi var. halophila	grass	C4
t1.225	Chloris sp.	grass	C4
t1.226	Cortaderia selloana (Schult. & Schult. f.) Asch. & Graebn.	grass	C3
t1.227	Cynodon dactylon (L.) Pers. var. dactylon	grass	C4
t1.228	Deyeuxia hieronymi (Hack.) Türpe	grass	C3
t1.229	Digitaria californica (Benth.) Henrard var. californica	grass	C4
t1.230	Digitaria sacchariflora (Nees) Henrard	grass	C4
t1.231	Digitaria sanguinalis (L.) Scop.	grass	C4
t1.232	Distichlis scoparia (Kunth) Arechav. var. scoparia	grass	C4
t1.233	Distichlis spicata (L.) Greene var. spicata	grass	C4
t1.234	Eleusine tristachya (Lam.) Lam.	grass	C4
t1.235	Eragrostis curvula (Schrad.) Nees	grass	C4
t1.236	Eragrostis lugens Nees	grass	C4
t1.237	Eragrostis mexicana (Hornem.) Link ssp. mexicana	grass	C4

Plant Soil

Family/species	Life form	Metabolic pathway
Eustachys retusa (Lag.) Kunth	grass	C4
Festuca hieronymi Hack. var. hieronymi	grass	C3
Gouinia paraguayensis (Kuntze) Parodi	grass	C4
Hordeum stenostachys Godr.	grass	C3
Jarava plumosa (Spreng.) S. W. L. Jacobs & J. Everett	grass	C3
Jarava pseudoichu (Caro) F. Rojas	grass	C3
Leptochloa chloridiformis (Hack.) Parodi	grass	C4
Lolium perenne L.	grass	C3
Muhlenbergia peruviana (P. Beauv.) Steud.	grass	C4
Nassella hyalina (Nees) Barkworth	grass	C3
Nassella tenuissima (Trin.) Barkworth	grass	C3
Nassella trichotoma (Nees) Hack. ex Arechav.	grass	C3
Neobouteloua lophostachya (Griseb.) Gould	grass	C4
Pappophorum pappiferum (Lam.) Kuntze	grass	C4
Paspalum dilatatum (Poir.) ssp. dilatatum	grass	C4
Paspalum malacophyllum Trin.	grass	C4
Paspalum notatum Flüggé	grass	C4
Paspalum quadrifarium Lam.	grass	C4
Piptochaetium sp.	grass	C3
Poa sp	grass	C3
Setaria lachnea (Nees) Kunth	grass	C4
Setaria oblongata (Griseb.) Parodi	grass	C4
Setaria pampeana Parodi ex Nicora	grass	C4
Setaria parviflora (Poir.) Kerguélen	grass	C4
Sorghum halepense (L.) Pers.	grass	C3
Spartina spartinae (Trin.) Merr. ex Hitchc.	grass	C3
Sporobolus indicus (L.) R. Br.	grass	C4
Sporobolus pyramidatus (Lam.) Hitchc.	grass	C4
Tragus berteroniamus Schult.	grass	C4
Trichloris crinita (Lag.) Parodi	grass	C4
Polygonaceae		
Rumex acetosa L.	herb	C3
Portulacaceae		
Portulaca oleraceae L.	herb	C3
Talinum fruticosum (L.) Juss.	herb	C3
Talinum paniculatum (Jacq.) Gaertn.	herb	C3
Talinum polygaloides Gillies ex Arn.	herb	C3
Ranunculaceae		
Clematis montevidense Spreng.	epiphyte	C3
Rhamnaceae		
Condalia buxifolia Reissek	shrub	C3
Condalia microphylla Cav.	shrub	C3
Condalia montana A. Cast.	shrub	C3
Ziziphus mistol Griseb.	tree	C3

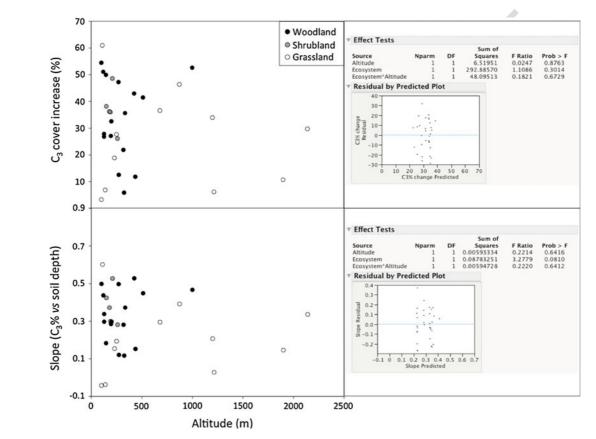
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Family/species	Life form	Metabolic pathway
Rosaceae		
Lachemilla pignata (Ruiz & Pav.) Rothm.	herb	C3
Rubiaceae		
Borreria eryngioides var. ostenii (Standl.) E.L. Cabral & Bacigalupo	herb	C3
Galium richardianum (Gillies ex Hook. & Arn.) Endl. ex Walp.	herb	C3
Mitracarpus megapotamicus (Spreng.) Kuntze	herb	C3
Rutaceae		
Zanthoxylum coco Gillies ex Hook. f. et Arn.	tree	C3
Santalaceae	1 1	C 2
Jodina rhombifolia (Hook. & Arn.) Reissek	shrub	C3
Sapindaceae	an in her to	C2
Cardiospermum halicacabum L. Urvillea chacoënsis Hunz.	epiphyte	C3
Selaginellaceae	epiphyte	C3
Selaginella peruviana (Milde) Hieron.	fern	C3
Selaginella sellowii Hieron.	fern	C3
Simaroubaceae	icin	05
Castela coccinea Griseb.	tree	C3
Solanaceae		03
Capsicum chacoënse Hunz.	herb	C3
	shrub	C3
Cestrum parqui L'Hér. Datura ferox L. Grabowskia aff. duplicata Arn. Lycium chilense Miers ex Bertero Lycium ciliatum Schltdl.	herb	C3
Grabowskia aff. duplicata Arn.	shrub	C3
Lycium chilense Miers ex Bertero	shrub	C3
Lycium ciliatum Schltdl.	shrub	C3
Lycium elongatum Miers.	shrub	C3
Lycium gilliesianum Miers	shrub	C3
Nierembergia aristata D. Don	herb	C3
Petunia axillaris (Lam.) Britton, Stern & Poggenb.	herb	C3
Salpichroa origanifolia (Lam.) Baill.	herb	C3
Solanum angustifidum Bitter	herb	C3
Solanum argentinum Bitter & Lillo	herb	C3
Solanum chenopodioides Lam.	herb	C3
Solanum elaeagnifolium Cav.	herb	C3
Solanum sisymbriifolium Lam.	herb	C3
Solanum stuckertii Bitter	herb	C3
Vassobia breviflora (Sendtn.) Hunz.	shrub	C3
Urticaceae		
Parietaria debilis G. Forst.	herb	C3
Urtica sp.	herb	C3
Verbenaceae		
Aloysia gratissima (Gillies & Hook. ex Hook.) Tronc.	shrub	C3
Glandularia peruviana (L.) Small	herb	C3

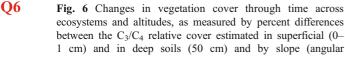
Plant Soil

t1.328	Table 1 (continued)			
	Family/species	Life form	Metabolic pathway	
t1.329	Lippia turbinata Griseb. f. turbinata	shrub	C3	
t1.330	Phyla canescens (Kunth) Greene	herb	C3	
t1.331	Verbena bonariensis L.	herb	C3	
t1.332	Zygophyllaceae			
t1.333	Larrea divaricata Cav.	shrub	C3	
t1.334	Porlieria mycrophylla (Baill.) Descole, O'Donell & Lourteig	shrub	C3	

604 Appendix 2



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606 References

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